

Tectonic-Geophysical Mapping of Israel and the Eastern Mediterranean: Implications for Hydrocarbon Prospecting

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ABSTRACT

The eastern Mediterranean is a tectonically complex region evolving in the long term located in the midst of the progressive Afro-Eurasian collision. Despite years of investigation, its geological-geophysical structure is not completely known. At the same time, the recent discovery of large gas deposits has attracted the attention of many researchers to this region. For instance, the latest U. S. Geological Survey estimates using conventional assessment methodology suggest that there are on the order of 1.7 billion barrels of recoverable oil and more than 4 trillion m³ of recoverable gas in the Levant Basin [1]. This highlights the need for analysis of the paleogeographical conditions that can yield deep paleotectonic criteria for oil and gas discovery in this region. For this purpose, isopach maps of the Middle-Upper Jurassic and Lower Cretaceous were generated from detailed examinations of numerous well sections and the most significant outcroppings in the eastern Mediterranean. The maps confirm an earlier model of continental accretion [2]. In particular, abrupt changes in the trend and thickness of the Early Mesozoic formations coincide with the terrane boundaries. These compiled isopach maps also pinpoint significant distinctions between the Arabian and Sinai plates on the one hand and the Syrian arc on the other. A new tectonic map of the eastern Mediterranean is presented that first of all integrates geophysical satellite-derived gravity and airborne magnetic fields, as well as tectonic-structural, paleogeographical and facial analyses. The results have clear implications for hydrocarbon prospecting in this region.

Keywords: Eastern Mediterranean, Isopach Maps, Hydrocarbon Potential, Satellite Gravity, Airborne Magnetics

1. Introduction

The eastern Mediterranean is a tectonically complex region evolving in the long term located in the midst of the progressive Afro-Eurasian collision. Its geological-geophysical structure has been studied for years, but is still not completely known. The eastern Mediterranean represents a classic area for the emergence of plate tectonics [2-7]. Recent discoveries of significant hydrocarbon deposits in this region (e.g., [8-10]) have heightened the need for a robust tectonic-structural description to assist in prospecting for hydrocarbon structures. This article reports on a combined geophysical (potential geophysical fields and seismic) and geological (formational, facial, tectonic-structural and paleogeographical) investigation conducted for this purpose.

In the eastern Mediterranean there are three main types

of crust: (a) platforms with a Precambrian crust, (b) regions with an oceanic crust, and (c) a folded belt of the transition type with a thinned continental crust [4]. However, the exact boundaries of these zones are not reliably known.

Integrated geophysical data on crust thicknesses [2,4, 11,12] indicate that the folded belt is composed of several terranes (Cyprus, Eratosthenes, the southern part of the Levant Basin (Pleshet), Galilee-Lebanon, Judea-Samaria and Negev) (**Figure 1A**). **Figure 1B** shows a velocity-depth distribution of compressional seismic velocities for these terranes (after [4] and [13]). It is obvious that these graphs exhibit different behaviors.

Alongside geophysical investigations, the Mesozoic formations were also mapped. The presence of a large number of deep wells in this region makes it possible to

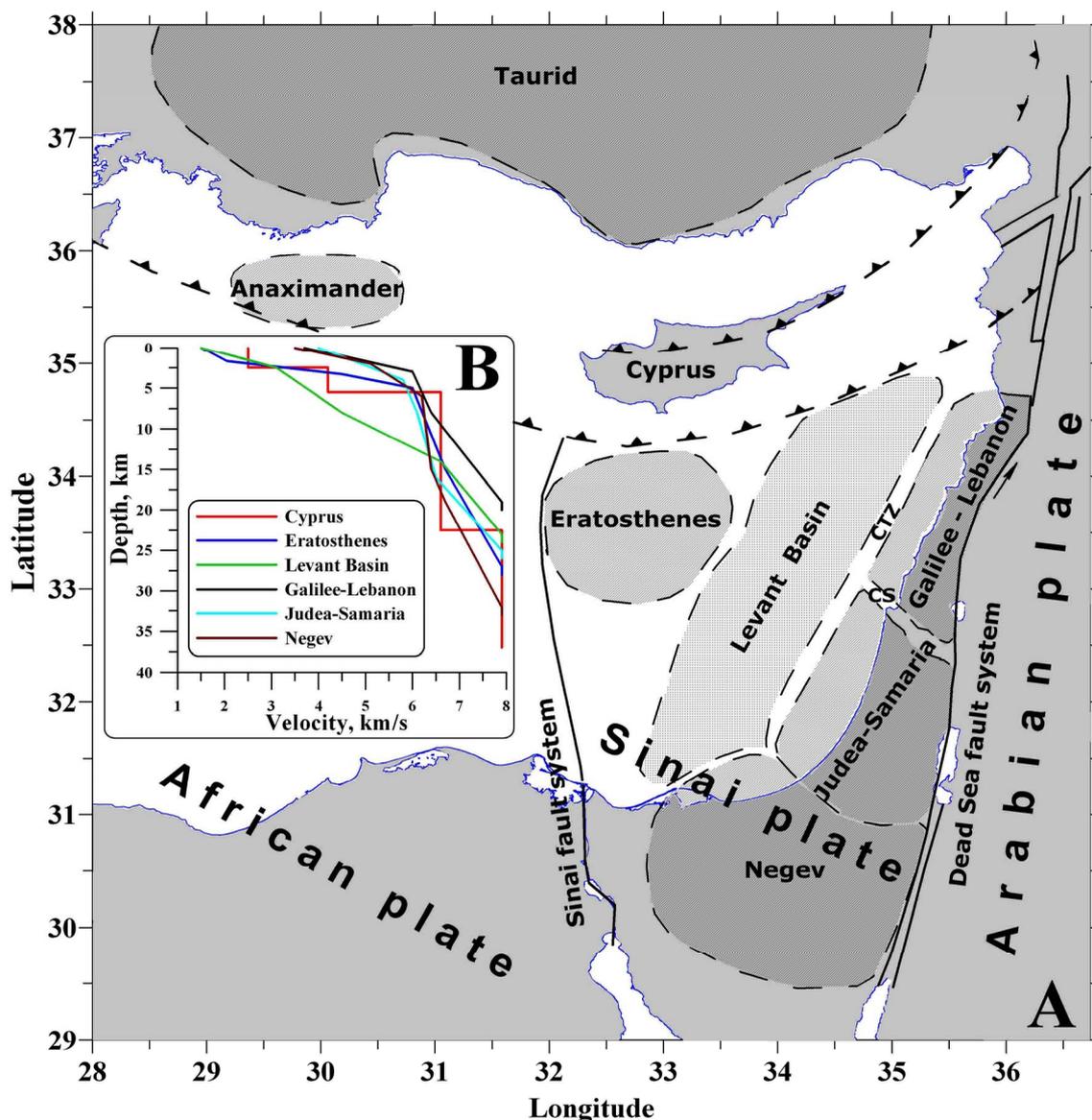


Figure 1. Tectonic scheme of the eastern Mediterranean (A) with generalized seismic sections (B). (A) after [11] and (B) after [4] and [12]. CS, Carmel structure; CTZ, crustal transition zone.

carry out geological mapping of the terranes and their boundaries. For this purpose a series of maps of the Early Mesozoic sedimentary rocks was constructed (here, the two most significant paleogeological maps are presented). These maps confirm an earlier model of continental accretion [2]. They show in particular that abrupt changes in the trend and thickness of the Early Mesozoic formations coincide with the terrane boundaries [14].

In this region numerous seismostratigraphic investigations have been carried out (e.g., [9,10,15-21]). These works have mainly been oriented toward the study of prospective hydrocarbons structures.

2. A Brief Geological Overview

According to modern geological-geophysical constructions, the eastern Mediterranean region covers four geotectonic plates: the African, Sinai, Arabian and Aegean-Anatolian (**Figure 1**). The first two plates consist of heterogeneous blocks of continental and oceanic crust and each of these plates is circumscribed by deep faults. According to Robertson *et al.* [22], the oceanic crust of the eastern Mediterranean is a remnant of the Neotethys. The Sinai plate [22] is bounded from the west and the east by a system of Late Cenozoic deep faults [3,24]. The eastern

part of the eastern Mediterranean is formed by the Dead Sea Fault System (**Figure 1**) along which about 100 km of left-lateral strike-slip motion has developed [25]. Within this fault system numerous horsts and grabens have emerged. These grabens usually represent pull-apart basins [26] filled by terrigenous evaporite molasses and basic rocks.

Between the oceanic and continental crust is the Syrian fold belt (Syrian arc) (e.g., [5]). This belt is the eastern part of the larger Alpine-Mediterranean fold belt and basin system [6,27]. The structural inhomogeneity of the Syrian arc was revealed by analysis of subsurface Paleogene-Upper Cretaceous sediments [28,29]. However, the tectonic characteristics of the deeper structural stages have been studied to a lesser extent. Geophysically it was shown [2] that a thinned continental crust composes the Syrian arc structure. Tectonically, this belt is made up of a sequence of extensional independent crystal units [11].

Precambrian, Paleozoic, Mesozoic and Cenozoic associations form the geological section of the African and Arabian platforms within the eastern Mediterranean. The Precambrian is known to contain metamorphic and igneous rocks, and in the upper part of Precambrian there is molassa which passes to the Cambrian base. Cambrian rocks are extensive in the Sinai, but are not found in the borehole sections of the Negev desert. In the Negev the Permian deposits overlie the Precambrian arkosic sandstone. In the Heletz block there is no molasses and the Permian deposits are in direct contact with the crystalline schist.

More completely than in the Sinai, Lower Paleozoic sequences are present to the east of the eastern Mediterranean coast in Syria, Jordan and Iran. Analysis of geophysical data testifies to the absence of Lower Paleozoic strata in the Levant Basin and in the Pleshet terrane [12,17,18]. Devonian and Carbon deposits have not been detected in the borehole sections of Israel. The Carbon and Permian continental associations were studied in Sinai [30]. Deep boreholes drilled in the Negev, the coastal plain of Israel, the Judean mountains and western Jordan have revealed marine Permian deposits [31,32]. Thus, the Precambrian and Paleozoic associations of the Arabian plate and southern Sinai differ significantly from the folded zone of the Syrian arc.

The Mesozoic and Paleogene associations make up the carbonate platform of the eastern Mediterranean (including reefs and other organogenic constructions). These carbonate associations also include small terrigenous complexes and traps. The Miocene and Quaternary associations are represented by terrigenous molassa with Early Messinian evaporates, sometimes generating salt

domes and diapirs. These associations are developed mainly in the pull-apart basins within the Dead Sea transform zone and Levant Basin [33]. Besides sedimentary deposits, in the eastern Mediterranean there are widely developed Mesozoic and Late Cenozoic magmatic rocks [29,34,35].

3. Materials and Methods

The tectonic-geophysical investigation of such a complex region requires careful analysis of numerous materials and data sources such as deep boreholes sections and major outcroppings, geological and geophysical maps, seismic regional constructions and seismostratigraphic examinations, as well as the results from potential geophysical fields.

Here, we utilized and examined more than 200 deep boreholes (isopach constructions were made from more than 120 boreholes) and about 10 of the most significant outcroppings (for instance, such as Jebel Maghara, Makhtesh Ramon and Mt. Hermon) (**Figure 2**). Data on borehole sections and outcroppings were taken from numerous publications and reports [8,17,19,21,30-32,36-49].

The geological materials included the following data: geological maps of Israel [29], Jordan [50], Egypt [46] and photomaps of Israel and adjacent areas [51]. The following materials on the dating of Mesozoic and Cenozoic magmatic associations were employed: [34,35,47,52,53].

The geophysical data used here encompassed a wide spectrum of geophysical fields. Of primary importance are the geophysical maps that included gravity and magnetic maps of Israel and the Levant [54-56], satellite gravity data [57], a map of the Moho discontinuity for the eastern Mediterranean [58], a map of the Curie discontinuity for Israel [58], the geothermal gradient map of Israel [59] and maps of earthquakes epicenters [60,61]. In addition we drew on the results of regional analyses of seismic, gravity, magnetic and thermal data in this region [2-4,11,13,20,62,63], seismostratigraphic examinations [10, 11,15-21], and detailed geophysical-geological studies in separate areas [14,64-74].

For the first time in this region we applied a regional integrated geological-geophysical zonation of deep-seated complexes. A facial-paleogeographical analysis and thickness analysis were conducted mainly for the localization of geological structures of the Early Mesozoic. This analysis was also employed to study both sedimentary and magmatic complexes. Paleobiogeographical examinations were utilized to assess initial positions of the terranes [75]. This integrated methodology was accompanied by a broad examination of previous geological-geophysical research in the region.

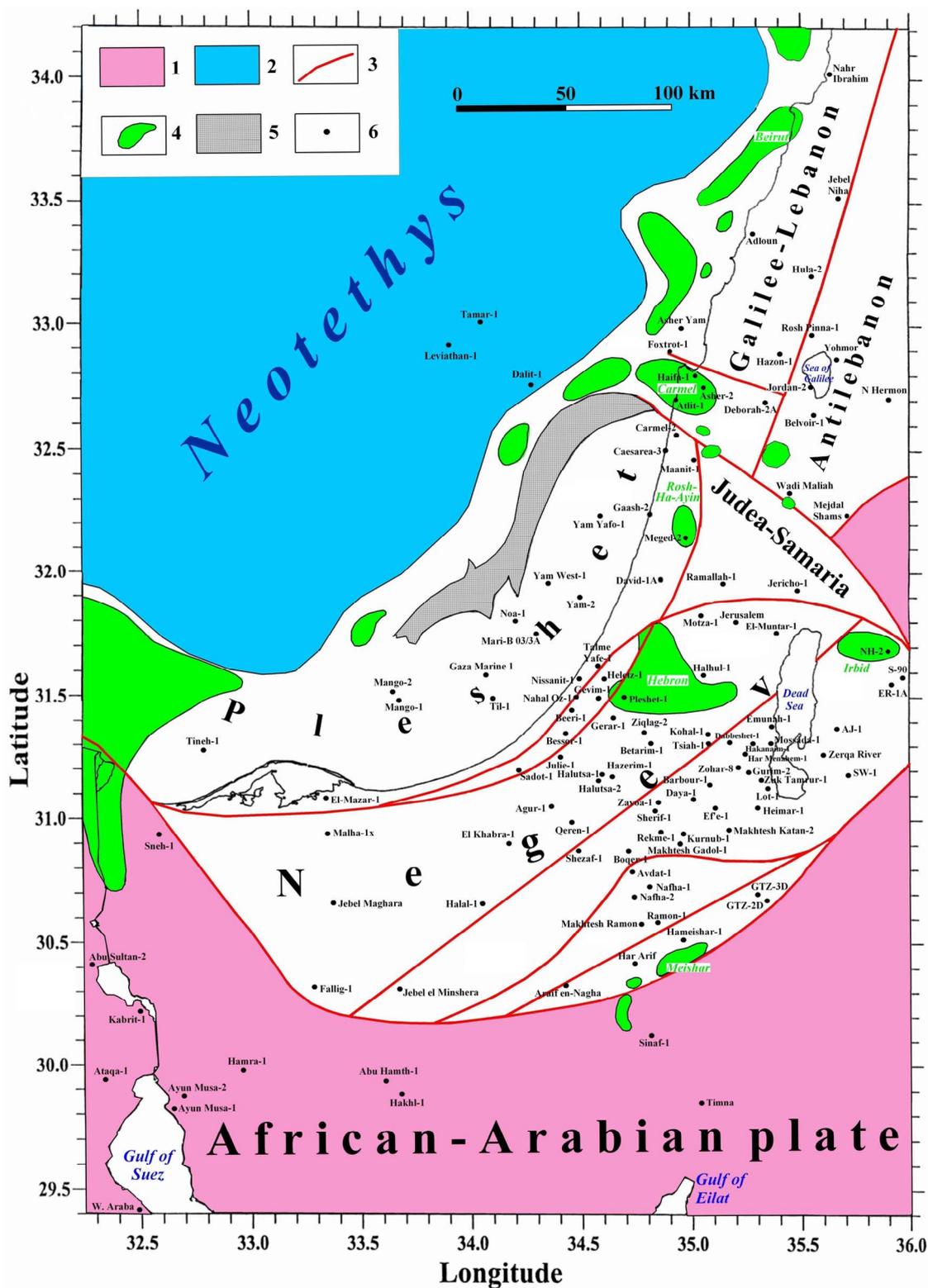


Figure 2. Map of Lower Mesozoic wells and outcrop locations in the eastern Mediterranean with some geological-geophysical zonation. (1) African-Arabian Precambrian plate, (2) Neotethys, (3) main faults, (4) contours of significant magnetic anomalies, (5) marginal erosional zone (Jonah Ridge), (6) deep boreholes and distinctive outcroppings.

4. Analysis of Geophysical Fields

Previous divisions of the area under study for the above-mentioned terranes have been made on the basis of comprehensive analyses of seismic, gravity and magnetic data (e.g., [2,4,11,62]). The first Curie map of Israel [58] also correlates with the position of the terranes (e.g., [11]). The sources of the two most significant gravity-magnetic anomalies in Israel – Hebron and Carmel – were classified as phenomena associated with tectonically weakened zones between the terranes [11,68,69].

The satellite gravity data were obtained from the World Gravity DB as retracked from Geosat and ERS-1 altimetry [57]. A highly positive factor is that these observations were made with regular global 1-minute grids [56] and the error of gravity data computation was estimated at 2-3 mGals. The compiled gravity map (**Figure 3**) shows the intricate gravity pattern of the investigated area (isoline interval is 5 mGal; “zero” isoline is dashed and bolded). This figure indisputably proves that in specific cases it is useful to display the gravity field map without any reduction. The selected positive and negative gravity anomalies (**Figure 3**) clearly reflect the main structural-geotectonic units of the region. The geomorphologic names of the corresponding gravity anomalies were taken from [11,18,32,76,77].

A map of the magnetic field horizontal gradients (**Figure 4**) was computed (a sliding interval of 2×2 km was applied) on the basis of an airborne total magnetic field analysis [54]. Some peculiarities of the horizontal magnetic gradient behavior indicate the presence of regional tectonic faults which mark crystal units (terranes and Precambrian platform) in northern Israel. The Cretaceous and Neogene-Quaternary local volcanic eruptions also follow these faults (**Figure 4**). Magnetoactive rocks producing high gradients generally correspond to surface and subsurface trap complexes, the thicknesses of which depend on the disjunctive tectonics in the region [53,71,74]. This map indicates that the sharpest gradients (125-750 nT/km) correspond to the Precambrian plate. The minimal gradients (0-12.5 nT/km) correspond to northern part of the Negev, Pleshet, Judea-Samaria and the northern part of the Galilee-Lebanon terranes. The location of the Neogene-Quaternary local volcanic eruptions was used as an additional indicator for the detection of significant faults.

In order to define the position of tectonic boundaries of the main structures in the eastern Mediterranean, an analysis of the regional magnetic anomalies (on the basis of [54] and [56]) was carried out. The main magnetic features it reveals are the high-amplitude anomalies of Hebron [54,68,78], Carmel and Beirut [62], Irbid [78]

and Meishar [74] (see **Figure 2**). Furthermore a low-amplitude anomaly was contoured on this map in Rosh-Ha-Ayin [64] and several unnamed low-amplitude magnetic anomalies associated with the boundaries of terranes. This map shows the positions of the deep boreholes and the most distinctive outcroppings of the Mesozoic sediments utilized for the construction of the series of isopach maps (see below).

5. Construction of Isopach Maps: Their Relation to Tectonic and Geophysical Patterns

Isopach maps of the Middle-Upper Jurassic and Lower Cretaceous (**Figures 5 & 6**, respectively) were constructed on the basis of a detailed examination of numerous well sections and the most significant outcroppings (a list of references is presented in Materials and Methods). The isopach maps were compiled on classical principles described, for instance, in [79]. A thorough analysis of these maps indicates that practically all the aforementioned terranes (**Figure 1**) are reflected in the Middle- Upper Jurassic and Lower Cretaceous isopach maps.

The compiled maps also capture significant distinctions between the Arabian and Sinai plates on the one hand and the Syrian arc on the other. Mesozoic associations in the ancient African-Arabian plate are developed to a lesser extent than in the Syrian arc and sometimes are lacking. From a facial point of view [80,81] thin terrigenous marine and continental associations are presented here mainly comparatively. The isopach map of the Middle-Upper Jurassic (**Figure 5**) indicates that the Cenozoic folded belt of the Syrian arc was formed within the complex system of paleotectonic blocks. These blocks correspond completely to the terranes identified earlier. The isopach map permits, for instance, to define terrane boundaries and to single out some block structures within these terranes. This map indicates that the axes of the terrane and platform paleostructures are located discordantly to each other. The axes of the platform Early Mesozoic eastern Desert [46] and Jebel Druze depression [31] are located at an angle of 90° to the axes of Syrian Arc paleostructures (**Figure 5**). Analyses of facies and thicknesses of Jurassic deposits show that the Late Cenozoic Dead Sea Transform in the southern part of the region dissects the Anti-Lebanon, Judea-Samaria and Negev terranes and displays their western and eastern blocks for 100 km.

The distribution of the Lower Cretaceous thicknesses (**Figure 6**) testifies to the lack of influence of the terranes' geotectonic boundaries on sedimentation. Rather, it is caused by the influence of erosion relief forms and denudation processes on the post-accretional sedimenta-

tion. As can clearly be seen from the map, the relief of the denudation surface is inclined from SE to NW at 1600 m to the studied region (Figure 6). The depth of erosional cutting reaches 500 m, and in the transition zone of the Tethys Ocean erosional processes have shaped the canyons to a depth exceeding 1000 m. This is an important criterion for regional oil-and-gas prospecting (an example is the Heletz oil deposit – the first oil field discovered in the eastern Mediterranean [82]).

Besides thickness analyses, a facial examination was also carried out. Given the series of bauxite iron-pisolith redbeds of the Jurassic weathering crust, the Negev and Pleshet terranes as well as the northern and southern Palmyride and Heletz blocks are joined to a common tectonic-paleogeographic structure. At the same time the

Galilee-Lebanon terranes and Anti-Lebanon as well as Aleppo belong to another tectonic-paleogeographical zone. The facial analysis highlights that the terranes were isolated structures until the end of Jurassic and the beginning of the Early Cretaceous. Evidently, their movements triggered the formation of dyke complexes with an estimated age of 140 m.y. [47]. The terrane accretion was estimated to be the Lower Cretaceous (Hauterive) [75] after which post-collision traps began to form in this region [74]. Besides these traps, terrigenous and at times carbonate associations were generated.

Thus, the data on the sedimentary thickness distribution provide an unambiguous confirmation of the terrain model of the evolution of the eastern Mediterranean derived from the regional geophysical data analysis [2].

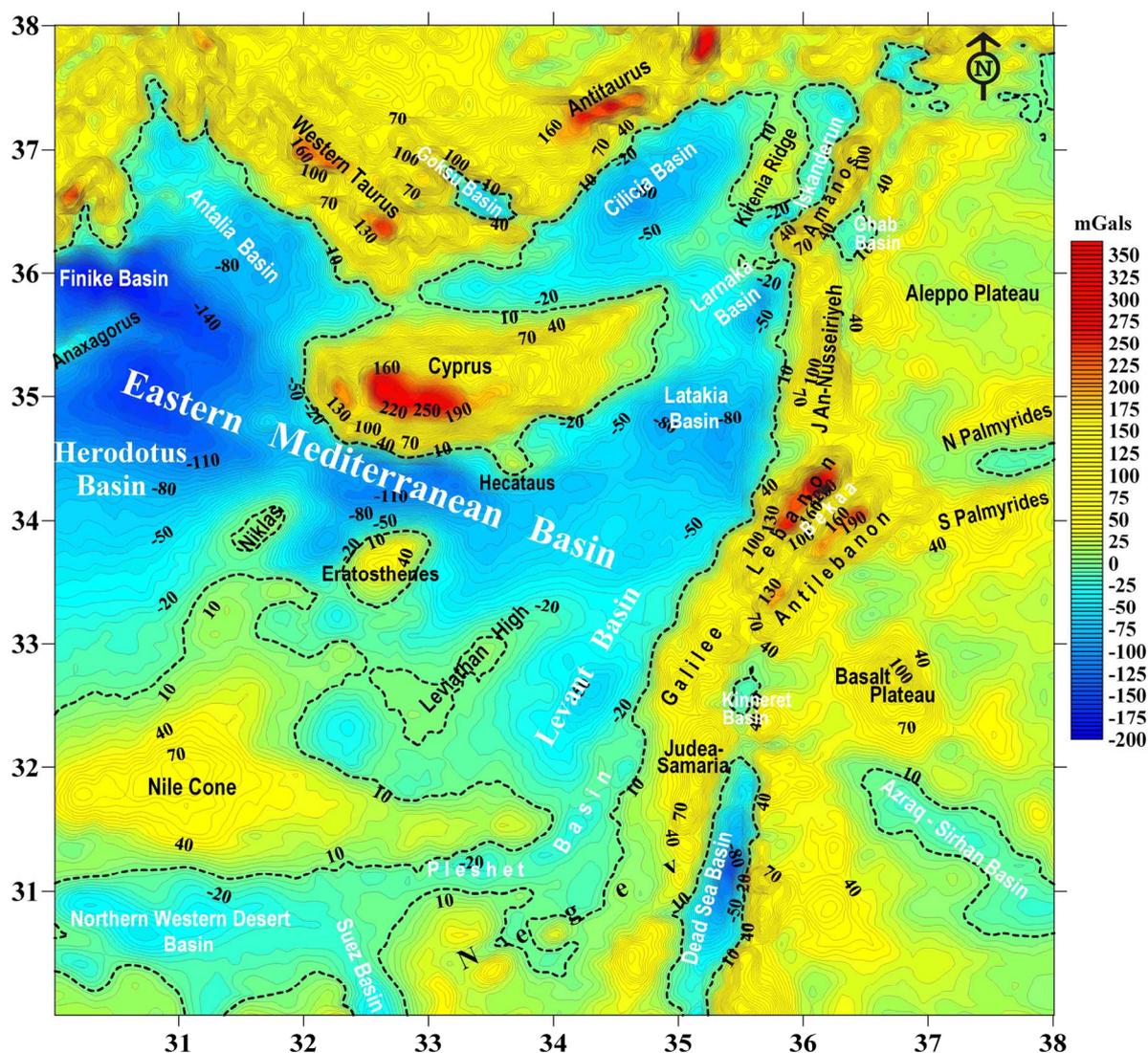


Figure 3. Map of significant gravity anomalies of the eastern Mediterranean and adjacent areas (isolines are given in mGals).

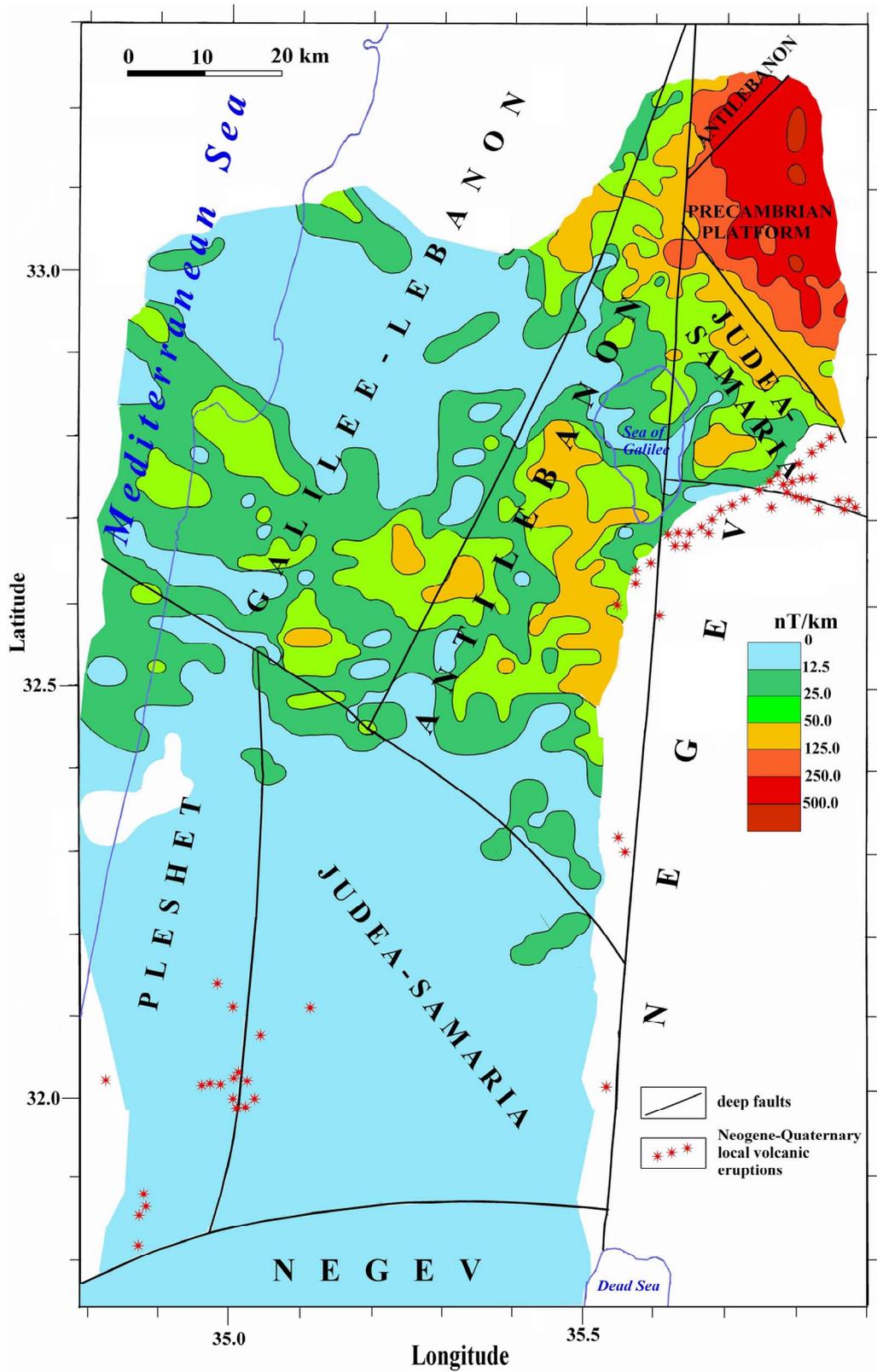


Figure 4. Magnetic gradient map of Israel and some adjacent areas with elements of tectonic analysis.

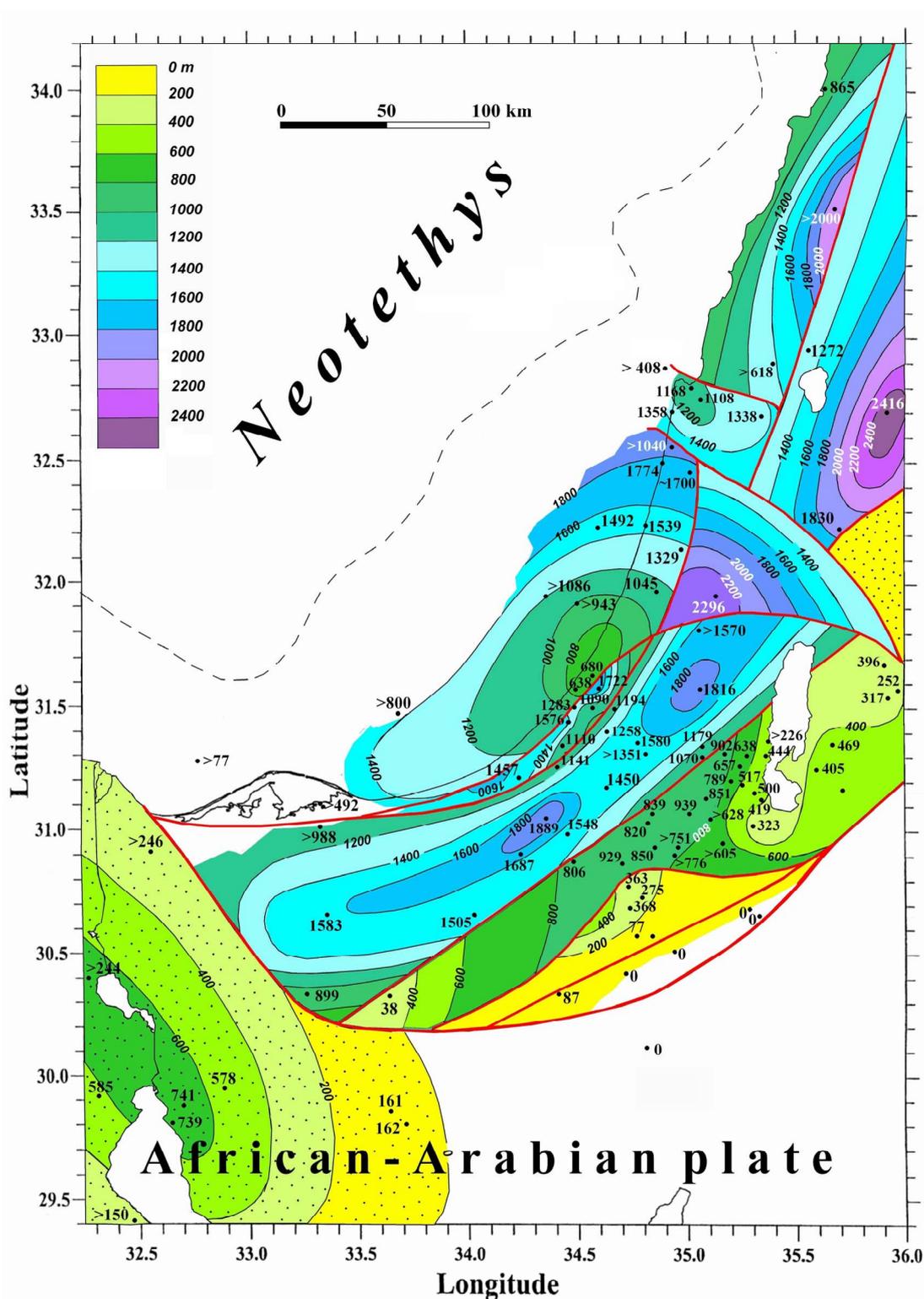


Figure 5. Isopach map of the Middle-Upper Jurassic for the Levant color isopach column is presented in the upper left corner. Black lines with numbers in italics represent isopachs. Numbers next to solid black dots designate determined Bajocian – Upper Jurassic thicknesses. Dotted areas show terrigenous (predominantly continental and shallow marine) deposits (other captions are the same as in Figure 3).

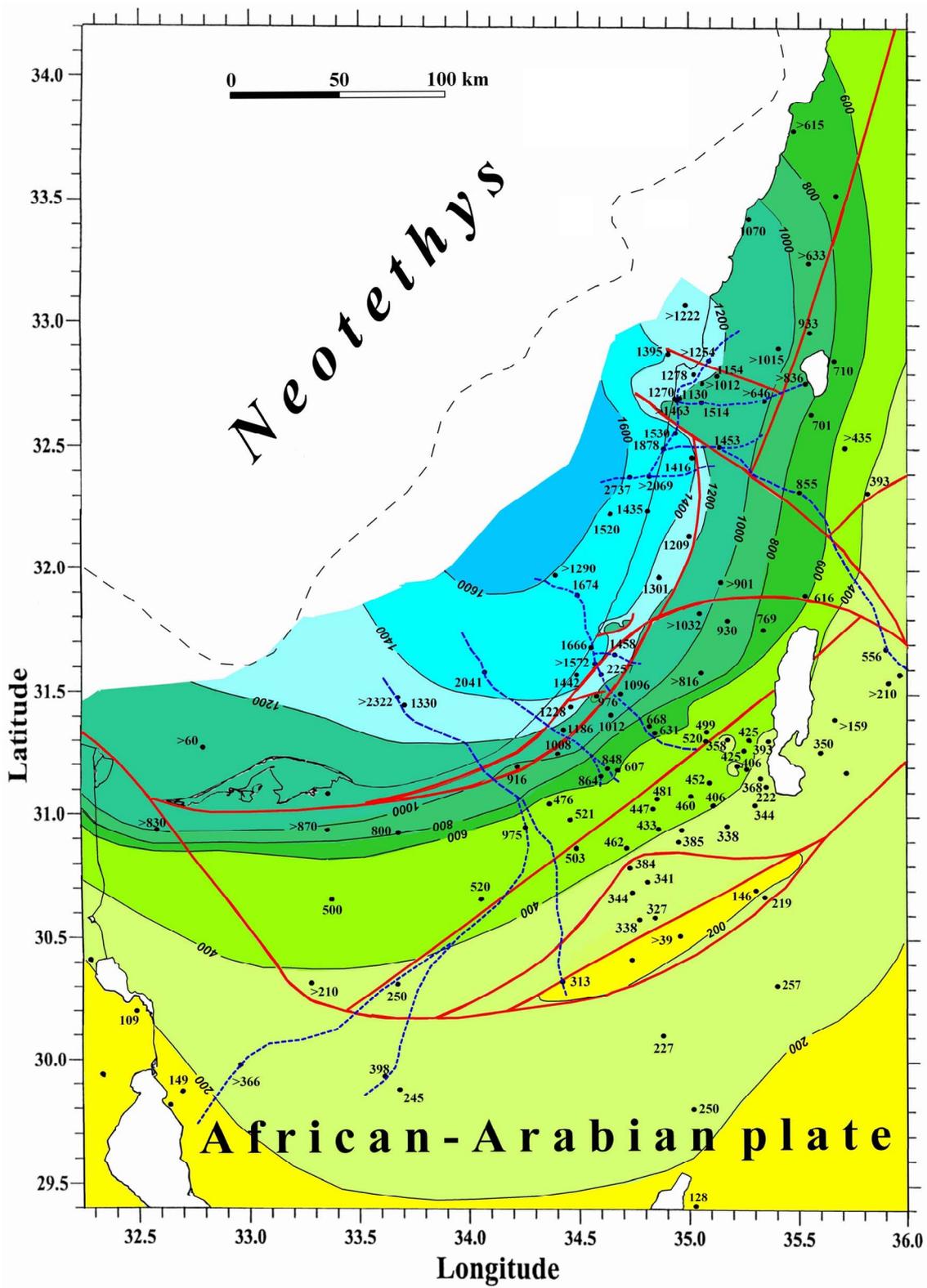


Figure 6. Isopach map of the Lower Cretaceous for the Levant Dashed blue lines designate the position of the ancient erosional system (other captions are the same as in Figures 3 and 4).

6. Application of Paleotectonic Zonation in the Heletz-Ashdod Oil Field (Central Israel)

Deep paleotectonic mapping (zonation) supported by gravity-magnetic examination was used to investigate the hydrocarbon fields of the central plain of Israel (Figures 7-9). In the detail drilled (about 100 boreholes) of the Heletz-Ashdod oil and gas area, paleogeographic and paleotectonic mapping based on the facies analysis was applied and the Jurassic and Cretaceous horizons were studied. Figure 7 shows the Early Cretaceous erosion section at a depth of up to 1000–1300 m, splitting the Jurassic thicknesses down to the Sherif formation of the Late Bajocian. The post-accretional Gevar formation of the Late Hauterivian – Barremian is well known as regards alluvial, deltaic and deep marine channel facies [83].

The structural differentiation of this erosional zone appears both in differences in the integrated thicknesses of Jurassic rocks (Figure 7) and in thicknesses of their separate horizons (Figure 8). The examined geological structures are mainly inherited ones. The Heletz oil field is a triangular depression within which thicknesses of the Upper Jurassic and Lower Cretaceous are sharply increased. The Lower Cretaceous thicknesses in this area offshore Israel reach maximal eastern Mediterranean values in the Heletz-33 (2359 m), Heletz-31 (2257 m) and Heletz-32 (2290 m) boreholes [42]. The Ashdod gas field may be described as an anomalous uplifted zone where thicknesses of the Jurassic and Lower Cretaceous horizons abruptly decrease.

Gravity and magnetic anomalies in the area under study are significantly influenced by some regional anomalies. Therefore, for better visualization, the gravity and magnetic maps are overlaid and also show the geological structural elements (Figure 9). In this integrated map the local structural elements of the oil and gas Heletz-Ashdod area stand out well. The clearest delineations are the Heletz triangle, Yinnon High, the Beeri depression and Gevim High (Figure 9). Note that the Talme Yafe arch, within which the Ashdod High is located, is characterized by a heterogeneous structure that calls for further detailed geological-geophysical investigation for prospective oil and gas structure detection in this area.

7. Combined Analysis of Geophysical and Tectonic Indicators as a Basis for Hydrocarbon Prospecting in the Eastern Mediterranean

The analysis of the regional gravity field (Figure 2) was combined with an integrated examination of the magnetic field [4,84], a thermal field analysis [4,63,66], the Moho

discontinuity map of the eastern Mediterranean and the Curie map of Israel [58], maps of earthquake epicenters (e.g., [50,61]), and seismic data analyses (e.g., [4,11]) to construct a modified tectonic map of the Eastern Mediterranean (Figure 10).

The satellite gravity map displays the terranes, the main and intraplate tectonic faults, the African-Arabian Precambrian basement with the continental crust, areas occupied by oceanic crust, Mesozoic terranes and the Alpine tectonic belt (Figure 10). To improve readability, the gravity map was smoothed and the isoline interval was increased to 20 mGals.

This map (Figure 10) was developed on the basis of regional tectonic reconstructions of the eastern Mediterranean [2-4,11-14,22,75,85] and the present investigation. This map depicts the main plate tectonic structures of this region: the African, Arabian, Sinai and Aegean-Anatolian plates. The latter plate is part of the Alpine mobile belt. Associations of the Precambrian continental crust, oceanic crust and Mesozoic accretional complexes compose the African, Arabian and Sinai plates. It should be noted that the earlier (e.g., [86]) Pelusium Line (some 60 km off the coast of Israel) in the northern part mainly coincides with the boundary between the oceanic crust and the Galilee-Lebanon terrane. Thus, the Pelusium Line cannot be a marker of the northern part of the Sinai plate (Figure 10). At the same time it agrees with the conceptual model proposed by Ben-Avraham [3] where the northern boundary of the Sinai plate collides with the Aegean-Anatolian (Turkish) plate.

In the region under study, several hundred deep boreholes were drilled. Among these roughly 40 known borehole hydrocarbon reserves have been discovered [8, 19,36,40,42,45]. Interestingly, borehole Tamar-1 (where a deposit was discovered with total reserves of more than 200 billion m³ of gas condensate) and the promising prospective borehole Leviathan [9] (the most recent estimates, in January of 2011 report about 450 billion m³ of gas condensate), are located in the vicinity of a positive gravity anomaly (Figure 10). The last U. S. Geological Survey estimate indicates the presence of 1.7 billion barrels of recoverable oil and more than 4 trillion m³ of recoverable gas in the Levant Basin Province [1].

Integrated methodology of the regional-structural, facial and paleo-geodynamical analysis can serve to estimate the overall hydrocarbon potential in the eastern Mediterranean. For the first time three structural-prognosis geotectonic zones for hydrocarbon prospecting were revealed: 1) a province located on the ancient Nubian-Arabian Precambrian platform, 2) a province of a terrane belt with thinned Precambrian crust, and 3) a province of basins with oceanic crust.

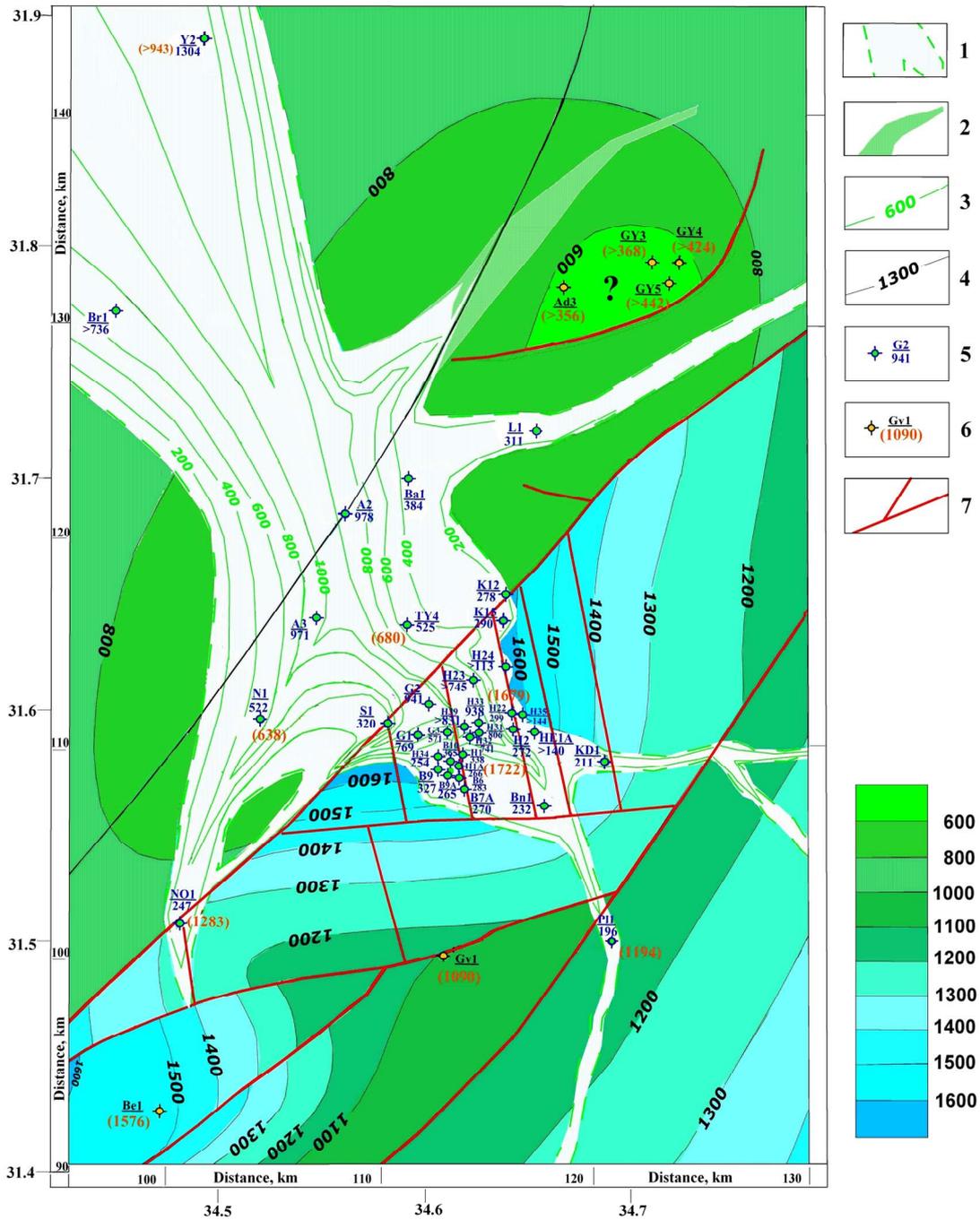


Figure 7. Paleogeomorphological map of GevarAm erosional stage (Hauterivian – Barremian) of the Heletz-Ashdod oil and gas field and adjacent areas. (1) Gevar Am deltaic, alluvial and marine canyon zone, (2) zone of posterior Gevar Am denudation, (3) isopachs of the Gevar Am formation (in m), (4) isopachs of the Upper – Middle Jurassic (Halutza – Upper Inmar formations) (in m), (5) boreholes penetrating Gevar Am formation (numerator – borehole name and number, denominator – thickness in m), (6) boreholes penetrating Jurassic rocks (numerator – borehole name and number, denominator – thickness in m), (7) faults. Borehole abbreviations: A – Ashkelon, Ad – Ashdod, B – Brur, Ba – Barnea, Be – Beeri, Bn – Binyamin, Br – Bravo, BW – Beeri West, G – Gevar Am, Gi – Givati, Gv – Gevim, GY – Gan Yavne, H – Heletz, Hd – Hodiya, Ho – Hoga, HE – Heletz East, K – Kokhav, Kr – Karmon, KD – King David, L – lior, MY – Massuot Yizhak, N – Nissanit, Ng – Negba, NA – Nir Am, NO – Nahal Oz, Pl – Pleshet, Rv – Revaha, S – Shimon, Sa – Saad, Sh – Shuva, SA – Sedot Akiva, TY – Talme Yafo, Y – Yam, Ya – Yakhini, Yi – Yinnon Color columns show the Middle – Upper Jurassic isopach intervals (in m).

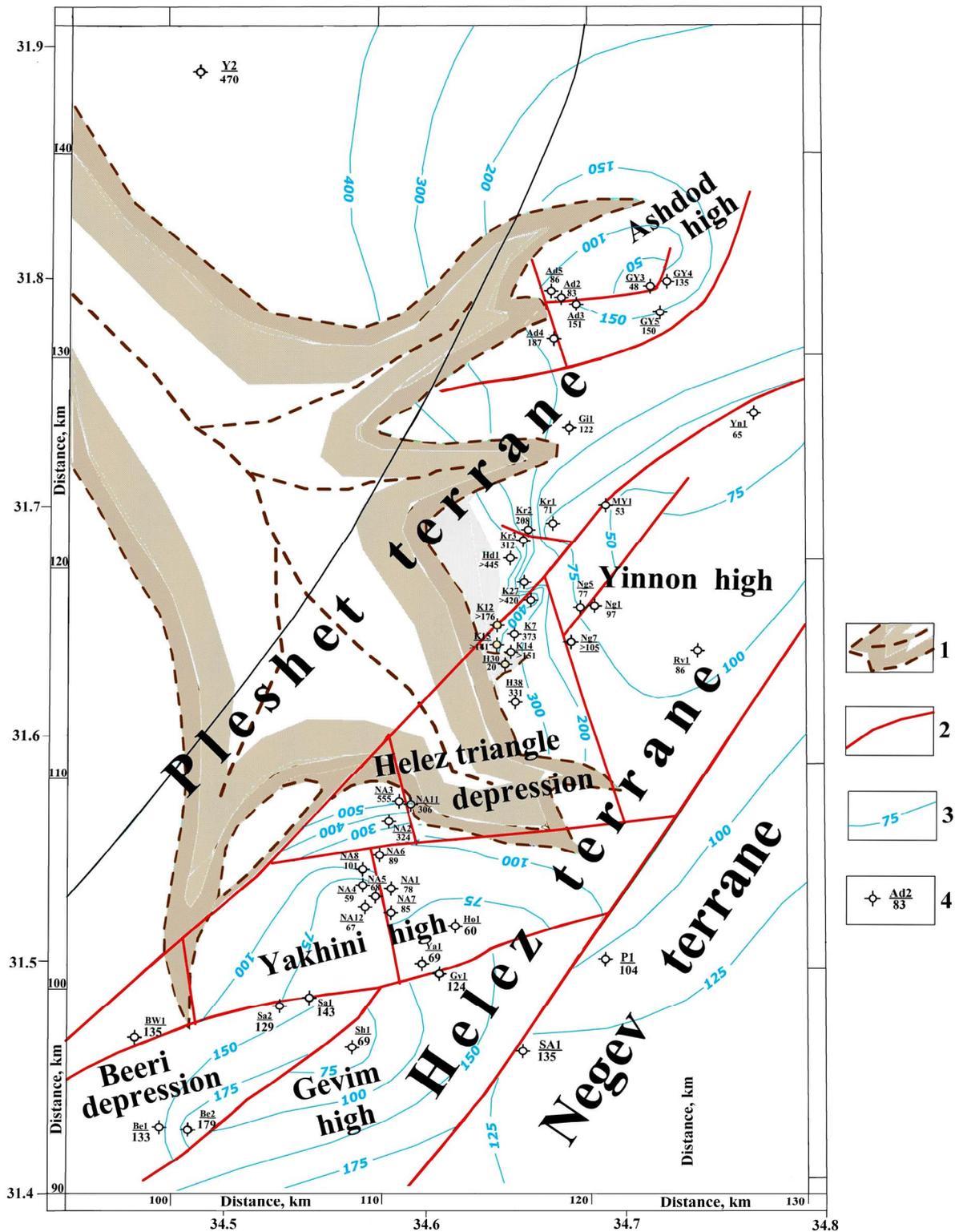


Figure 8. Late Hauterivian – Barremian structural-paleogeographic map of the Heletz-Ashdod oil-gas field and adjacent areas. (1) The Early Cretaceous Gevar Am erosional zone, (2) isopachs of the Middle Oxfordian (Beer Sheva formation) (in m), (3) boreholes penetrating the Beer Sheva formation (numerator – borehole name and number (see caption to Figure 7), denominator – thickness in m), (4) faults.

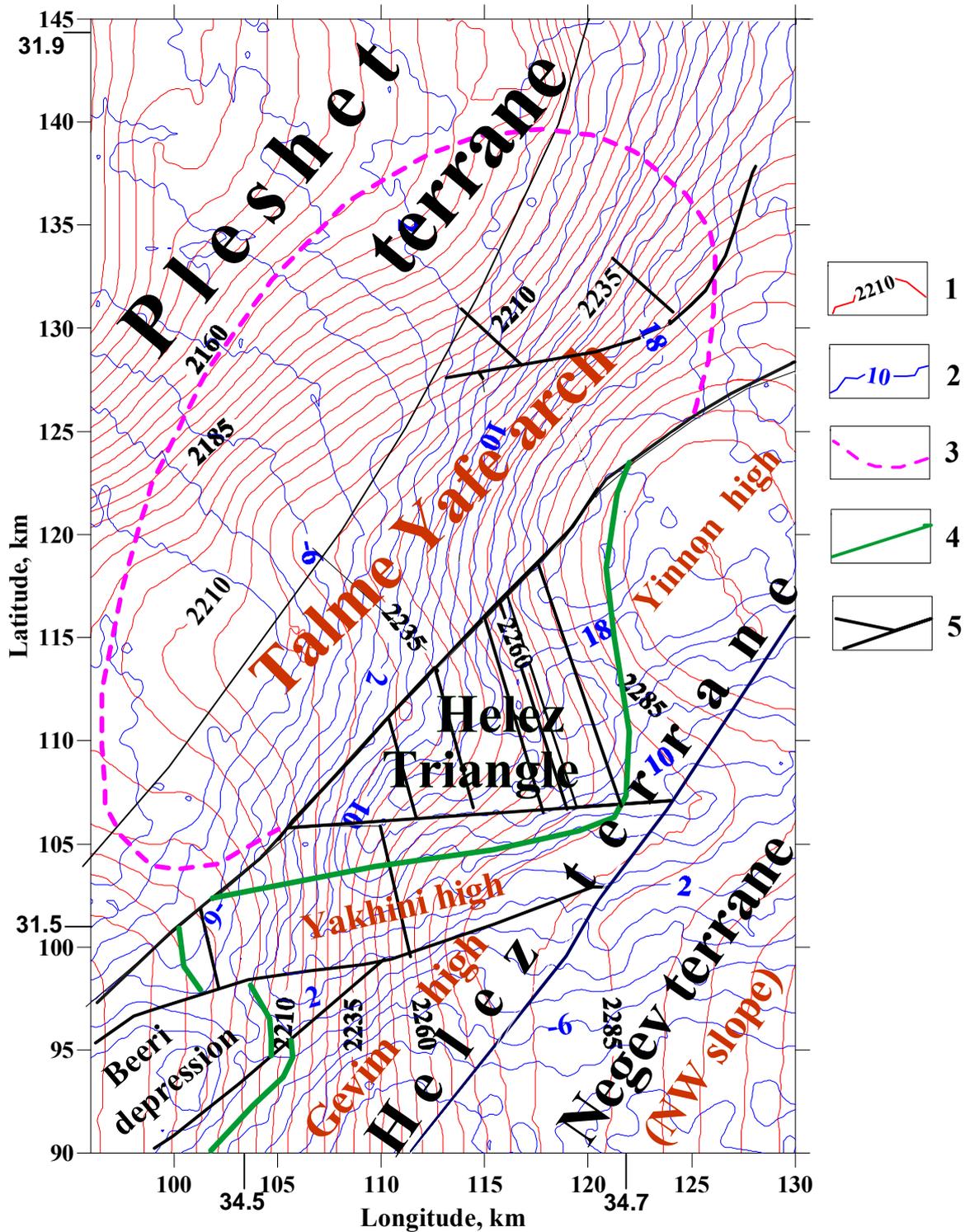


Figure 9. Combined gravity-magnetic map overlaid on the Late Jurassic-Early Cretaceous paleotectonic structure of the Heletz-Ashdod oil and gas field and adjacent areas. (1) isolines of Bouguer anomalies (in mGal), (2) isolines of magnetic anomalies (in nT), (3) isopach 800 m of the Middle - Upper Jurassic rocks in the Pleshet terrane, (4) isopach 1300 m of the Middle - Upper Jurassic rocks in the Heletz terrane, (5) faults. The gravity and magnetic maps were constructed using the Gravity-Magnetic Database compiled by the Geophysical Institute of Israel.

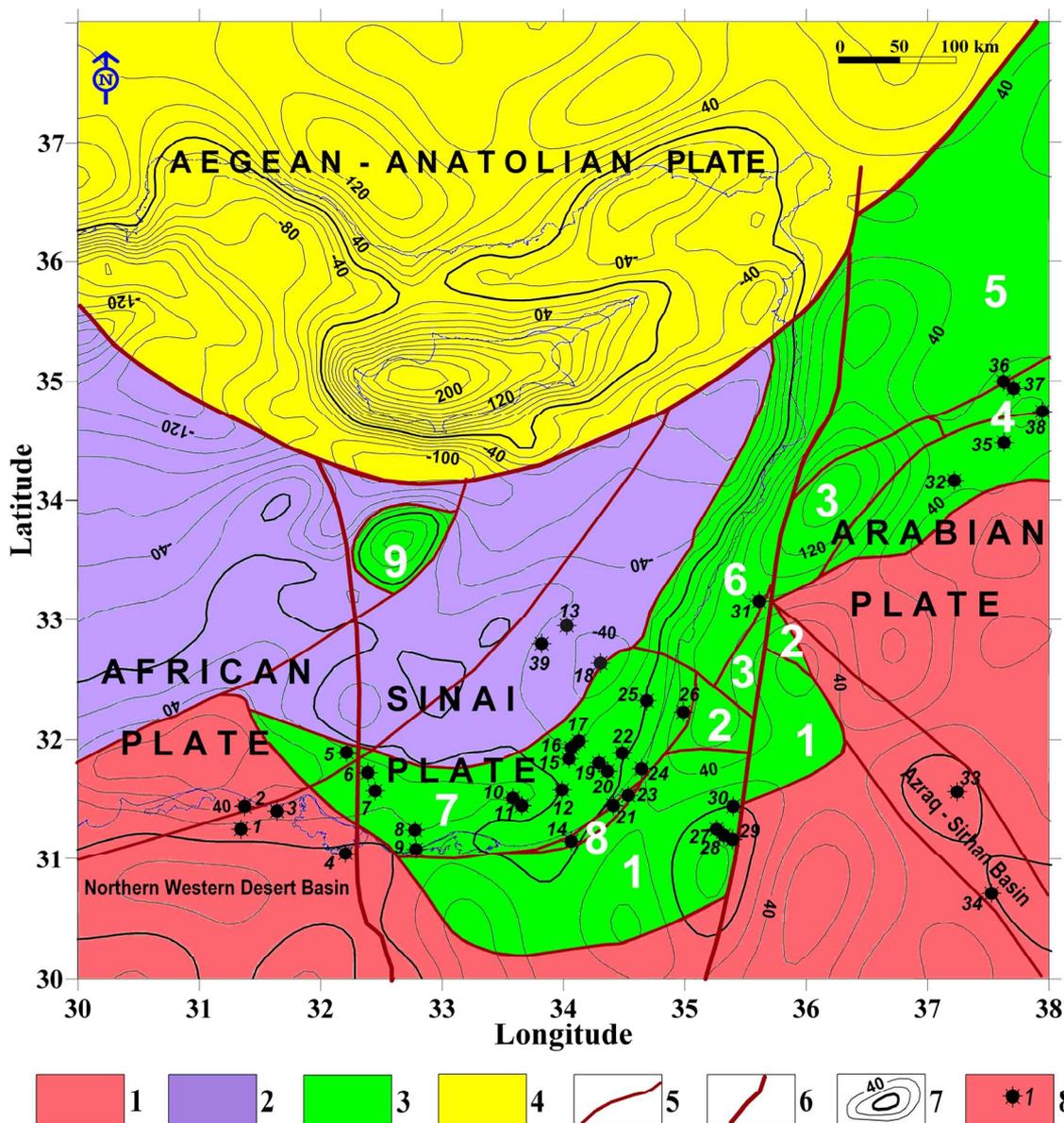


Figure 10. Tectonic map of the eastern Mediterranean overlaid on the smoothed gravity map. Terranes position (white numbers): 1 – Negev, 2 – Judea-Samaria, 3 – Antilebanon, 4 – Palmyride, 5 – Aleppo, 6 – Galilee-Lebanon, 7 – Pleshet, 8 – Heletz, and 9 – Eratosthenes. (1) Precambrian plates with continental crust, (2) oceanic crust, (3) Mesozoic terranes, (4) Alpine tectonic belt, (5) main tectonic faults, (6) intraplate tectonic faults, (7) gravity isolines, mGal, (8) boreholes with discovered hydrocarbon reserves (black numbers): 1 – Khelala-1, 2 – El-Qara-1, 3 – Wastani-1, 4 – Qnantara-1, 5 – Temsah-1, 6 – Kersh-1, 7 – Port Fouad-1, 8 – Tineh-1, 9 – Bougaz-1, 10 – Mango-1, 11 – Mango-2, 12 – Gaza Marine, 13 – Tamar-1, 14 – Sadot-1, 15 – Noa South-1, 16 – Noa-1, 17 – Or-1, 18 – Dalit-1, 19 – Mari-1, 20 – Nir-1, 21 – Shiqma-1, 22 – Yam-2, 23 – Heletz, 24 – Ashdod, 25 – Yam-Yafo-1, 26 – Meged-2, 27 – Zahar-8, 28 – Gurim-1, 29 – Zuk-Tamrur-4, 30 – Emunah-1, 31 – Hula, 32 – Qaruateine-1, 33 – Hamza, 34 – Wadi Sirhan-4, 35 – Cheriffe-1, 36 & 37 – boreholes with unknown names, 38 – Al Shaier-1, 39 – Leviathan-1.

The hydrocarbon basins of the first province are narrow, intraplate marginal graben-like troughs. Their formation was initiated in the Late Cretaceous era. However, the main stage of their development and hydrocarbon migration are associated with new Neotectonics epochs

(30 m.y.a. – to present) and are connected to the Red Sea system riftogenesis. There are known oil deposits (October and other oil fields in the Suez Basin) [87], ozozerites (Dead Sea Basin) [88] and small gas deposits (Hula Basin) [89]. The prospects for oil and gas discov-

ery are associated with the presence of a thick series of the Early Messinian salt.

In the second province, three structural floors have developed in a sedimentary cover: (a) Late Precambrian molasses, (b) the Pre-accretional carbonate platform of the Late Permian – Early Mesozoic, and (c) the Post-accretional Late Mesozoic-Cenozoic predominantly terrigenous complex. Oil and gas deposits were found in floors (b) and (c). The carbonate platform (b) has thickness up to 5-6 km and hydrocarbon searching is oriented here to discovering reef traps. The present investigation helps define a search strategy for these objects based on a combined historical-geodynamical and paleogeographical examination. Taking into account that the thick Mesozoic carbonate series were formed in terranes of the peripheral zones of the Tethys Ocean, the presence of reef complexes and the backreef sediments in this region is obvious. A predominantly terrigenous complex (c) mostly developed in the coastal plain of the eastern Mediterranean and directly in the shelf zone (its thickness reaches 5-6 km). Hydrocarbon collectors in this area are the Late Cenozoic terrigenous rocks associated with Late Messinian salt. The Early Cretaceous erosional channel complexes are of special interest. The long-term Heletz oil deposit (central Israel) (see Section 6 and **Figures 7-9**) belongs to just this type. In the Early Cretaceous paleogeographical map (**Figure 6**) for the first time a complex system of ancient erosional channels is shown. Therefore, this map could be used for hydrocarbon investigation in this region. The presence of a majority of hydrocarbon deposits and a geotectonic basis in the Pleshet terrane (**Figure 10**) is confirmed by the present work. In this terrane terrigene-carbonaceous rocks and salt predominantly occur. Such a facial type is known in many oil and gas provinces worldwide.

The third province consists of many basins, but our analysis focuses solely on the largest Levant basin (**Figures 1,2,10**). The tectonic-geophysical analysis was used to designate these basin boundaries with respect to the zone of terrane belt and other eastern Mediterranean basins of this province. The western part of this basin is bounded by the Leviathan high and eastern part – by the continental-ocean transition zone (**Figure 10**). Note that in the central part of the Levant basin a Moho discontinuity lowering-in was detected previously (see **Figure 3** in [58]). The Early Mesozoic-Cenozoic facial characteristics of the sedimentary cover are similar to the corresponding cover of the Pleshet terrane. The available data indicate that structural zones have developed here of the oceanic type as found in the Gulf of Mexico, the Pricaspian basin and the southern Caspian basin. Ben-Avraham [3] claims that there are intensive geodynamic movements at

the boundary of the Sinai and Aegean-Anatolian plates allow us to suggest that in the Levant basin numerous prospective local structures and swells may occur. Hence, it is likely that the high-potential hydrocarbon deposits in Tamar-1 and Leviathan (**Figure 10**) discovered recently in this region are not isolated features.

8. Conclusions

This article presented the results of combined structural isopach map compilations and geophysical field analyses of the deep floors of the eastern Mediterranean region, thus constituting the first deep geological-geophysical mapping ever produced. For the construction of the isopach maps of the Middle-Upper Jurassic and Lower Cretaceous a regional structural analysis method was applied (previously employed mainly for mapping subsurface structures in this region). The same analysis for zoning the Heletz-Ashdod oil field was successfully applied. It revealed that the Syrian arc is an accretional belt of the Mesozoic terranes, which was formed in the Early Cretaceous. It was shown that in the Early Cretaceous post-accretional stage numerous erosional channels, which may be used for searching for hydrocarbon collectors, were formed. An integrated methodology of regional-structural, paleogeographical and paleogeodynamical analysis was used to estimate the overall hydrocarbon potential of the eastern Mediterranean. Three structural-prognosis geotectonic zones for hydrocarbon prospecting were delineated: 1) a province located on the ancient Nubian-Arabian Precambrian platform, 2) a province of a terrane belt with a thinned Precambrian crust, and 3) a province of basins with oceanic crust.

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