

Measurements of Natural Radioactivity in Some Granite Samples Using Alpha Spectrometric Analysis

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ABSTRACT

Alpha spectrometry using pulse height analysis has been used for the determination of uranium concentrations in different environmental samples. The concentration of ²³⁸U was measured by both destructive and non-destructive techniques with a detection limit of less than 1.8 mBq/kg. However, because of the extremely low ²³⁴U concentrations in environmental samples, it was necessary to use a destructive technique to separate U from the sample matrices as well as remove interfering elements from the sample solution to determine ²³⁸U/²³⁴U ratio. In this study, the uranium was separated from the environmental samples using anion exchangers in (Dowex 1 × 8 Cl⁻ form) and purified via co-precipitation with Lanthanum fluorides (LaF₃) and the alpha source prepared by electrodeposition. The results obtained were validated using some certified reference samples.

Keywords: Uranium Separation; α -Spectrometry; Alpha Pulse Height Analysis Isotopic Ratio; Granite Samples

1. Introduction

Uranium is a widely distributed lithophile metallic element. It may be present as a significant constituent in some minerals (e.g. uraninite, brannerite and carnotite) or as an accessory element in others (e.g. zircon, apatite, allanite and monazite). The natural uranium consists of three radioisotopes; ²³⁸U, ²³⁵U and ²³⁴U with atomic abundances of approximately 99.275%, 0.72%, and 0.0055% respectively. All three isotopes comprise the natural uranium and have the same geochemical behavior [1]. The uranium concentration and U-isotopic ratios are usually detected and determined in various environmental samples by different non-destructive and destructive techniques. The non-destructive techniques are mostly achieved by γ -spectrometry (e.g. NaI- and HPGe-detectors). They are carried out on the bulk samples without the need for complicated and time consuming radiochemical separations methods [2,3].

Moreover, the destructive techniques are carried out through several analytical methods (e.g. α -spectrometry, fluorimetry, kinetic phosphorescence, neutron activation analysis, etc). Among these techniques, α -spectrometry is

the most common one that measures radioisotopes and can detect low uranium concentrations (below ng l⁻¹). Its detection limit is typically 100 to 1000 times lower than γ -spectrometry [4]. This technique is mostly used for detection and analysis of U as well as Th radioisotopes, particularly in the environmental samples, such as natural waters, which are characterized with low radioactivity concentration levels [3,5].

The procedure for alpha spectrometry is carried out through several steps including sample preparation, radiochemical separation, preparation of a thin alpha source, such as via electro-deposition or co-precipitation, and α -counting employing high-resolution pulse height analysis [6-9]. Sample preparation aims to convert the sample into a thin layered, chemically isolated form that can be placed into the spectrometer and counted with a minimum alpha particle energy interferences and self absorption, and alpha peak broadening due to energy straggling with thick sources. It is often an extensive process and requires several steps including: 1) sample digestion (preliminary treatment); 2) uranium separation and purification 3) alpha source preparation, and 4) alpha counting over extended periods for low detection limits. The uranium separation is usually

carried out using various techniques such as co-precipitation, liquid-liquid extraction, ion exchange and extraction chromatography. Three main methods are commonly used for preparation of the sample on a stainless steel disc (source preparation) namely; 1) direct evaporation from an organic solvent; 2) electro-deposition and 3) coprecipitation with NdF. The later method is more preferred technique but it requires careful preparative steps to eliminate organic material and to adjust the pH of the electrolyte. Several methods and flow charts have been reported in different literatures for sample preparation for U-analysis using α -spectrometry [6,10,11].

2. Experimental Work

2.1. Non-Destructive Analysis

Six granite samples were collected from eastern desert of Egypt and were prepared for destructive and nondestructive technique. For nondestructive analysis, the samples were mechanically pulverized and passed through 0.8 mm mesh sieve. Samples were collected using the standard methods to get composite sample that represents each site. The dried and sieved portion of the samples were transferred to Marinelli beakers of 100 or 1000 ml volume and sealed at least for 4 weeks to reach secular equilibrium between radium and thorium, and their progenies. ^{226}Ra (^{238}U series), ^{228}Ra (^{232}Th series), and K activities were measured using gamma-spectrometry based on hyperpure germanium detectors. The HPGe detector had a relative efficiency of 40% and full width at half maximum (FWHM) of 1.95 keV for ^{60}Co gamma energy line at 1332 keV and operated with Canberra Genie 2000 software for gamma acquisition and analysis. The gamma transmissions used for activity calculations were 351.9 (^{214}Pb), 609.3, 1120.3 and 1764.5 keV (^{214}Bi) for the ^{226}Ra series, 338.4, 911.1 and 968.9 keV (Ac) for the Th-series and 1460.7 keV for ^{40}K presented in gamma spectrum below in **Figure 1**. The gamma-spectrometers were calibrated using both ^{226}Ra point source and potassium chloride standard solutions in the same geometry as the samples [12].

2.2. Destructive Analysis

For destructive analysis, 10 g of ashed soil sample was

spiked with uranium tracer (^{232}U) for chemical yield monitoring. The dried samples were ashed at 550°C for eight hours. The ashed sample was dissolved in 40 ml of 65% HNO_3 , 15 ml of 37% HCl and 10 ml of 40% HF acids. Uranium in the dissolved sample solution was extracted from most of the matrix elements with 25 ml of 0.2 M TOPO/Cyclohexane (Trioctyl-phosphine oxide) and then back-extracted with 25 ml of 1 M $\text{NH}_4\text{F}/0.1$ M HCl solution. The solution is co-precipitated by LaF_3 (25 mg/ml of $\text{La}(\text{NO}_3)_3$ with HF 40%). Then, the solution is centrifuged and the formed precipitate is dissolved in hot boric acid (saturated solution) and HNO_3 . The uranium is re-oxidized to the hexavalent state by adding H_2O_2 . This followed by evaporation of the solution to dryness and the obtained residue is dissolved in 9 M HCl , and then passed through a conditioned anion exchange resin column (15 cm long; its inner diameter is 8 mm) at a flow rate of 1 ml/minute. The used resin is 2 g Dowex 1×8 Cl^- form, 50 - 100 mesh (strongly basic gel type polystyrene resin) with appropriate functional groups. To elute U from the column, 0.5 M of HNO_3 is passed through the column with a flow rate of 1 ml/minute and the eluted U is evaporated to dryness in a crystallizing dish using 1 ml of concentrated HCl . The eluted uranium is transferred into the electrolysis cell from the crystallization dish with 0.4 ml of 4 M HCl , three times by 1 ml of $(\text{NH}_4)_2\text{C}_2\text{O}_4$ (4%) and then once 0.6 ml distilled water. The electrolysis is carried out for 3 hours at 300 mA (0.3 A), and then 1 ml of ammonia solution (NH_4OH with 25% conc.) is added. After one minute, the electrolysis current is cut off. The ammonia increases the OH concentration which prevents re-dissolution of the hydroxide from the cathode surface and then measured by alpha spectrometry [13]. Schematic radiochemical procedure of uranium is shown in **Figure 2**.

2.3. Uncertainty Calculation

The error associated with any particular counting result is determined by the use of the following equation.

$$\sigma = \frac{\sqrt{N}}{t} = \sqrt{\left(\frac{N}{t^2}\right)} = \sqrt{\frac{r}{t}} \quad (1)$$

where, r : is the net count rate.

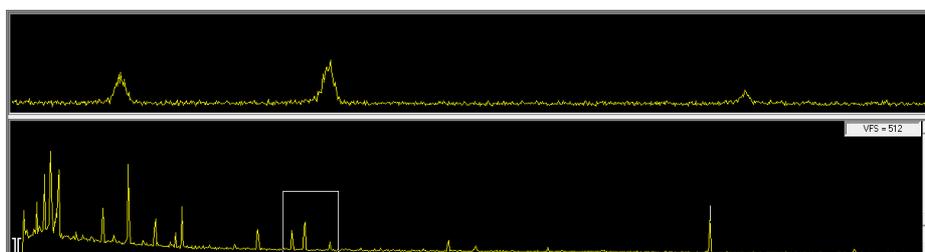


Figure 1. An example of gamma spectra.

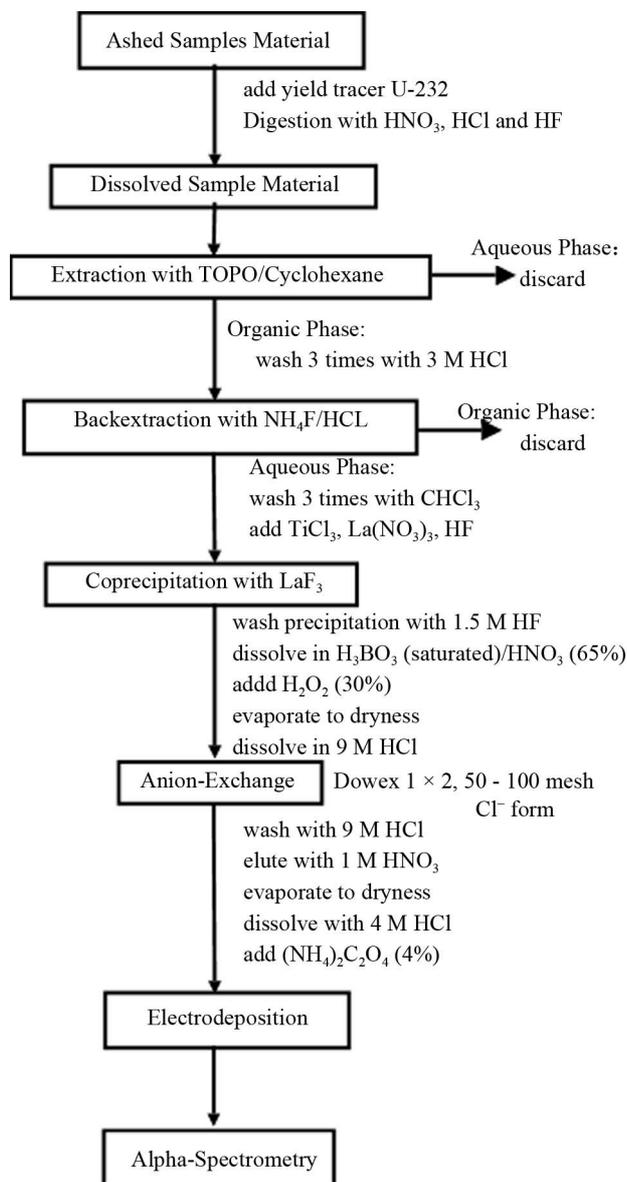


Figure 2. Schematic representation of the radiochemical procedure of uranium.

In this case we are interested in subtracting one count from another (gross counts minus background counts) and determining the resulting % error of the NCPS (Net Count Per Second) based on the standard deviation σ value. Counting instruments typically have a confidence interval of 95%. Thus Equation (2) is written as:

$$\sigma_n = \sqrt{\frac{r_{0Y}}{t_0} + \frac{r_{sY}}{t_s}} \quad (2)$$

where,

$-r_{0Y}, r_{sY}$: are the net count rate at the gamma line (Y) for the background and the sample respectively,

$-t_0, t_s$: are the real counting time of the background and the sample respectively.

2.4. Validation of the Method

The precision and accuracy of the method were determined by analyzing reference materials: soil IAEA-326, IAEA-375 and sediment IAEA-300. The precision achieved was 6.7% for U isotopes. Typical lower limits of detection for the alpha measurements were 1.6 mBq/kg for ^{238}U and 1.8 mBq/kg for ^{234}U . Blank samples and reagent blanks were processed and measured at the beginning of the analysis to batch to trace any cross contamination which might occur during the analysis steps. The data obtained shows good accuracy without any sign of cross contamination.

2.5. Apparatus

The alpha spectrometry system employed 450 mm² silicon surface barrier detectors, (ORTEC model 576 A) with 450 mm², USA. The silicon surface barrier detector was characterized by high resolution performance, low background, excellent stability and high permissible counting rates. The detector resolution was about 25 keV for ^{241}Am and the detector efficiency was approximately 23% with no significant variation in the range interval 2.5 - 8.8 MeV. It was determined using the following equation:

$$\eta = \frac{N}{A \cdot t_c}$$

where, η is the detector efficiency, N is the counts of the alpha peak, A is the activity of the radionuclide and t_c is the counting time.

The system was vacuum controlled (anti-recoil) and controlled with ORTEC software for calculation of the radionuclide activity. The counting time used for measurements was 4 - 8 days, depending on the sample activity, to achieve a detection limit of about 0.002 Bq per sample. The counting time can be reduced by increasing sample weight, but it was found that it is cost effective because as the sample weight increase the chemicals and acids needed for digestion will increase and the time required for sample digestion will also increase. The measured ^{238}U , ^{235}U and ^{234}U activity concentrations were reported in Bq/kg as shown in **Figure 3**. The chemical yield for the process involved in alpha sample analysis is around 70%. The system energy calibration was performed with a mixed alpha source containing ^{239}Pu ($E_\alpha = 5.1$ MeV), ^{241}Am ($E_\alpha = 5.48$ MeV) and ^{244}Cm ($E_\alpha = 5.8$ MeV) radionuclides. They have the same chemical composition, concentration, geometry as well as counting configuration. The detection limit of the α -spectrometry was about as 0.002 Bq per sample [14,15].

3. Results and Discussion

3.1. Non-Destructive Analysis

The three most common primordial radionuclides inves-

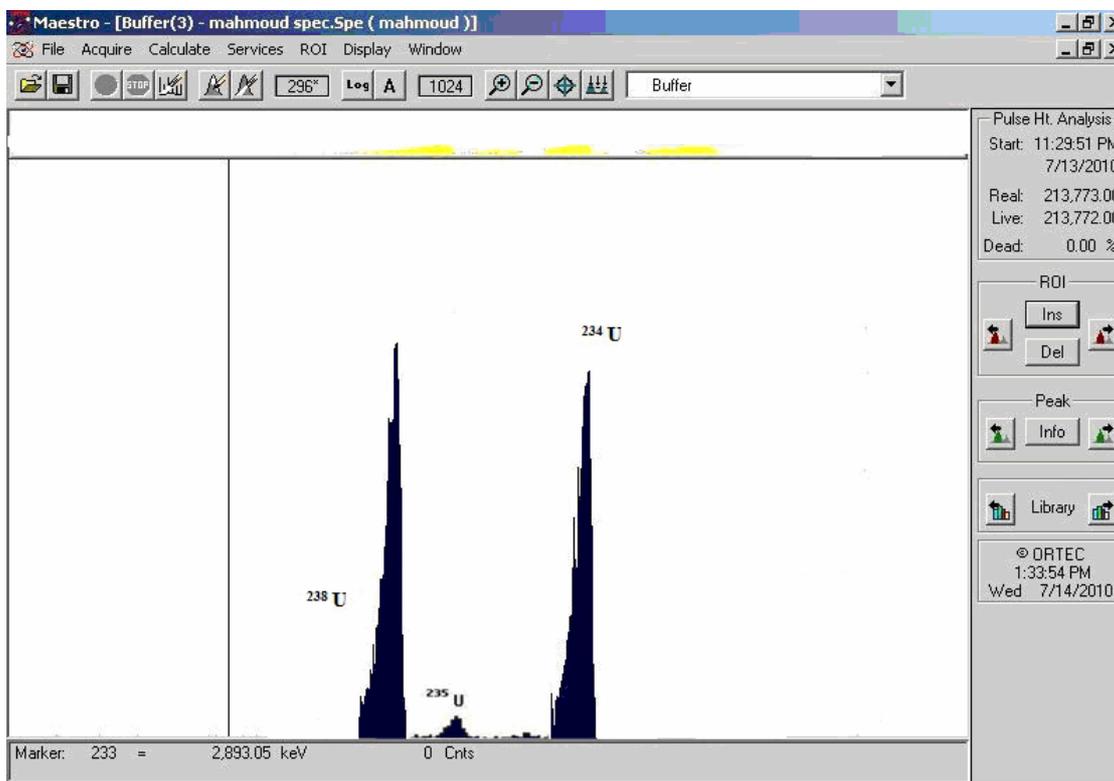


Figure 3. An example of U-isotopes measured by Alpha Spectrometer.

tigated in the study area were ^{40}K , ^{238}U (^{226}Ra) and ^{232}Th . The tabulated activity for the naturally occurring radionuclides ^{238}U (^{226}Ra) and ^{232}Th are the average of the activities of most abundant photo peaks of the decay products of the uranium series (351, 609, 1120, 1764 keV) and thorium series (238, 583, 911 keV). The specific activities of ^{238}U , ^{232}Th and ^{40}K for the collected samples were shown in **Table 1**. ^{238}U , ^{232}Th and ^{40}K concentrations ranged from 44.9 to 149 Bq/kg, from 4.8 to 79.6 Bq/kg and from 19.7 to 334.1 Bq/kg respectively. The high concentrations of ^{238}U activity might be due to its geological formation.

3.2. Destructive Analysis

The concentrations of ^{234}U , ^{235}U , and ^{238}U as well as activity ratios of $^{234}\text{U}/^{238}\text{U}$, and $^{235}\text{U}/^{238}\text{U}$ measured were shown in **Table 2**. The specific activities of ^{238}U ranged from 13.2 to 66.1 Bq/kg with an average of 37.2 Bq/kg, while the specific activities of ^{234}U ranged from 13.8 to 62.6 Bq/kg with an average of 36.7 Bq/kg. The ratios of $^{234}\text{U}/^{238}\text{U}$ ranged from 0.94 to 1.05 Bq/kg with an average of 0.98 Bq/kg. A very good correlation ($r^2 = 0.998$) exists between ^{234}U and ^{238}U as shown in **Figure 4**.

The specific activities of ^{235}U were below the detection limit of the system. The isotope ratio involving the minor isotope namely ^{234}U , can be obtained with a reasonable accuracy of about 5%, which is promising especially

Table 1. The activity concentration in (Bq/kg) for the analyzed samples measured by HpGe Detector.

Sample	^{226}Ra (^{238}U -series)	^{228}Ra (^{232}Th -series)	K-40
S1	91.1 ± 3.9	52.0 ± 1.6	288.0 ± 6.0
S2	77.2 ± 3.1	4.8 ± 0.2	19.7 ± 1.5
S3	129.9 ± 5.1	79.6 ± 2.1	334.1 ± 6.9
S4	149.0 ± 5.3	52.0 ± 1.7	288.2 ± 6.1
S5	139.3 ± 5.0	75.6 ± 1.9	360.8 ± 3.9
S6	44.9 ± 3.9	25.3 ± 1.5	152.8 ± 5.3

Table 2. The activity concentration in (Bq/kg) for the analyzed samples measured by α -spectrometry in reference and granite samples.

Sample	^{238}U	^{234}U	$^{234}\text{U}/^{238}\text{U}$
S1	40.3 ± 1.1	38.4 ± 1.1	0.95
S3	58.6 ± 1.4	60.2 ± 2.7	1.02
S6	13.2 ± 0.6	13.8 ± 0.4	0.95
IAEA-326	32.0 ± 1.4	31.4 ± 3.4	0.98
IAEA-375	13.2 ± 0.4	13.9 ± 0.4	1.05
IAEA-300	66.1 ± 4.8	62.6 ± 5.1	0.94

with lower sample volume.

The validations of the method were tested using IAEA reference materials IAEA-326 (soil), IAEA-375 (soil), and IAEA-300 (sediment) samples. The values obtained were in good agreements with the reference values reported by IAEA indicating robustness of our procedure.

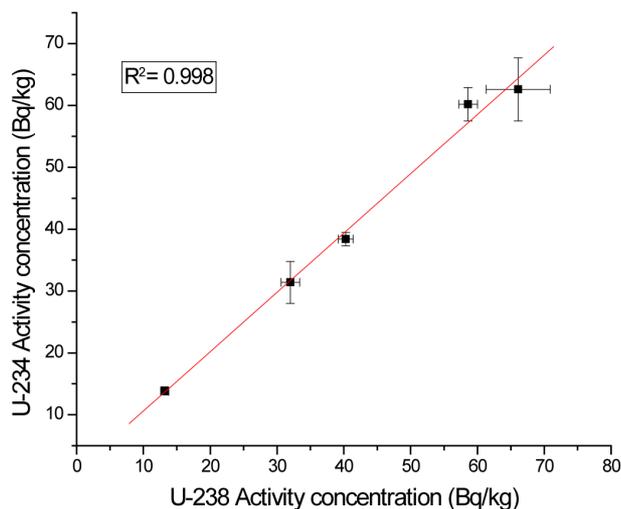


Figure 4. Correlation between U-238 and U-234 measured by alpha spectrometry.

4. Conclusion

A simple analytical technique for the determination of uranium isotopes in soil samples was developed and validated by testing several IAEA reference samples. The results obtained using current procedure indicates that these radioactivity concentrations are of natural origin. The isotopic values of approximately one between ^{234}U to ^{238}U indicate secular equilibrium between these two isotopes in the soil samples.

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