

Hydrodynamics and Recharge of the Aioune Sandstone Aquifer in the Taoudenni Transboundary Basin in Mauritania

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Abstract

In Mauritanian, 4/5 of the territory is subject to an arid climate, where groundwater resources are the main sources of drinking water for populations and livestock. The study area is located in south-eastern Mauritania and in the western part of the Taoudenni Transboundary Basin. A better knowledge of the hydrogeological system to ensure a sustainable and rational management is necessary for this vulnerable water resource. This pioneering study, based on the compilation of hydrogeological, geochemical and isotopic data, aims to characterize the hydrodynamics and recharge of this one layer aquifer system, composed of Aioune infracambrian sandstone. The piezometric surface exhibits relatively the same morphology of the soil topography, and computed effective velocity is 94 m/yr (Northeast area), and the direction of flow is generally from cliff to plains area. The chemical data show water types evolving between two poles namely the HCO₃-Ca/Mg pole (low mineralized), and the Cl-Ca/Na pole; these waters are characterized by a high vulnerability to pollution, especially at shallow depth area (Douerare, Aioune). The dissolution of bicarbonate and evaporate minerals is likely the main source of mineralization. The isotopic tool used to support comprehension of the system functioning, evidences a local indirect recharge water localized at the cliff or Wadi zones. The age of the waters ranges from the present to 1060 year. The recharge rate calculated by ¹⁴C and the chloride masse balance is respectively 0.4 to 5 mm/year. This recharge compared to the water pumping rate is lower and hence this imbalance is likely the cause of water shortages particularly during summer season period.

Keywords

Aioune, Isotopes, Hydrochemistry, Piezometry, Recharge

1. Introduction

In Mauritania, 4/5 of the territory is subject to an arid climate, where groundwater resources are the main sources of drinking water supply for populations and livestock.

The study area is located in southeast Mauritania, and covers the southern part of the Taoudenni Basin in Mauritania. It is characterized by a high concentration of livestock (49%) [1], and a high spatio-temporal variability in rainfall and increasing population pressure on water resources, which cause water shortages and decrease in boreholes yield.

The implementation of sustainable water resources management plan is essential in this context which requires a better knowledge of the system functioning. The present work aims to investigate the hydrodynamic and hydrochemical characteristics of Aioune sandstone aquifer, the most important generalized aquifer in the Taoudenni transboundary basin in Mauritania, and to identify and quantify the recharge at this Aioune sandstone aquifer. Hydrodynamic, hydrochemical and isotopic data used in this study are powerful tools to identify the origin, the age of the water, mineralization processes as well as location of recharge areas in order to establish measures for sustainability of these valuable water resources [2].

2. Study Area Description

2.1. Geography

The study area called “Afolle” is located in the southern part of the Taoudenni Transboundary Basin in Mauritania. It has a triangular shape and covers an area of 21,000 km². In the center of the basin, culminates a cliff of 350 m IGN which is strongly affected by weathering and its borders are covered in the North, East and West by recent sedimentary layers (**Figure 1**).

2.2. Climate

The climate in the study area is generally Sahelian. Rainfall is irregular, and occurs mainly between July and October, with an average of 200 mm at Aioune station (**Figure 2(a)**). 70% of annual rainfall events have an intensity between 10 - 30 mm [3]. The series of rainfall data (1939-2006) is characterized by a deficit cycle (1969-1990), followed by slightly increase relative to annual mean (**Figure 2(b)**). The average temperature is 30°C, with peaks of heat in May and June. Mean annual PICHE evaporation at Aioune zone is 5.32 m [4].

2.3. Hydrography

Apart the Senegal River, there is no perennial flow in Mauritania. In the study area, the hydrographic network is more active in the center, with a coefficient of flow between 1.5% - 10% [4].

The tributaries of the Senegal River in the study area are mainly the “Kolimbine” Wadi to the east of Afolle, and to the West the Wadi “Karakoro”. These

Wadis are currently cut by sandy cover (endoreism), which is often present in the study area [4] (Figure 3).

2.4. Geology and Hydrogeological Context

Several French geologists, studied this zone, from the 19th century to the 1960s (H. HUBERT, R. HUDEAU, R. FURON and C. BENSE) and had identified an anticlinal structure inferred by tectonics, with the North, East and West edges plunging under recent sedimentation layers (Figure 1). The Afolle relief is characterized by a succession of tabular formations affected by numerous diaclasses. The oldest sedimentary formations which outcrop are the sandstones of the

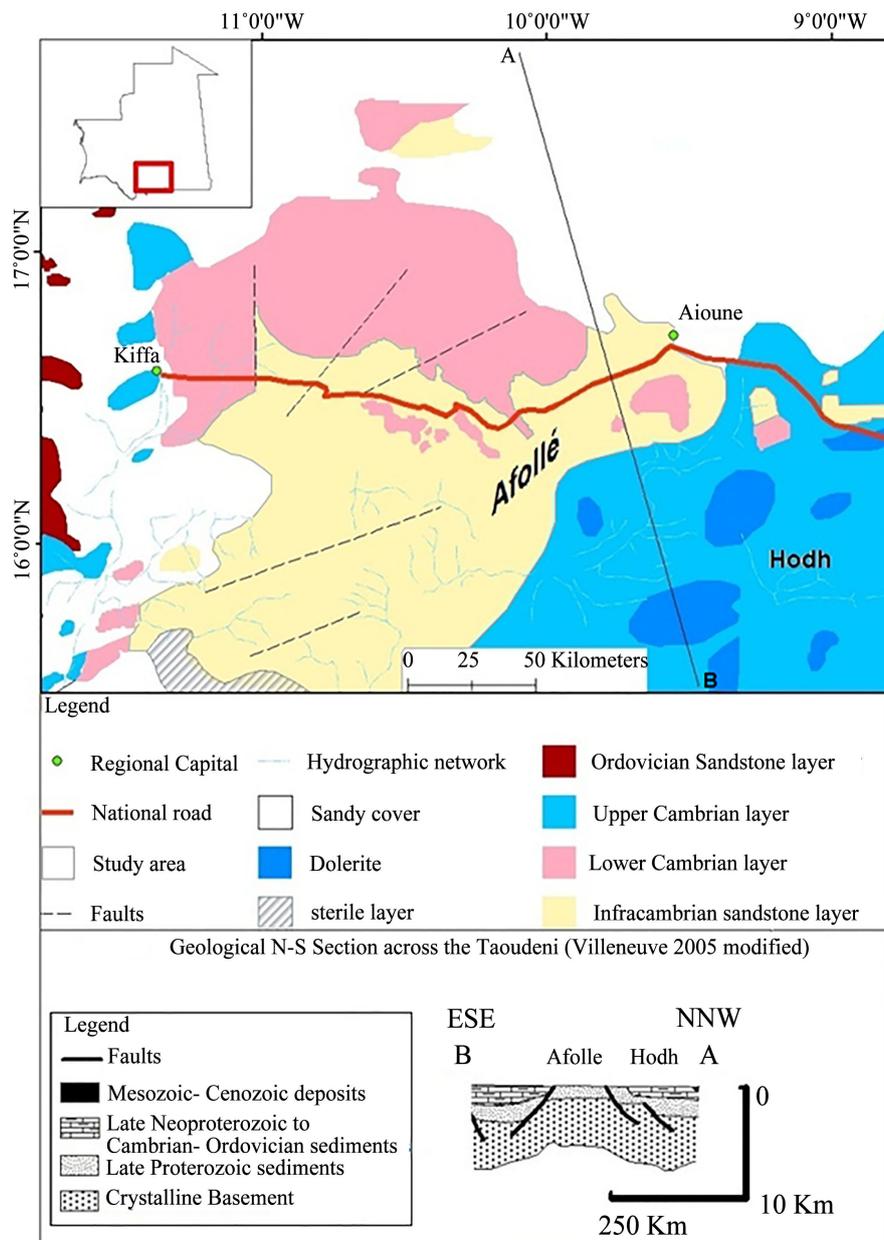
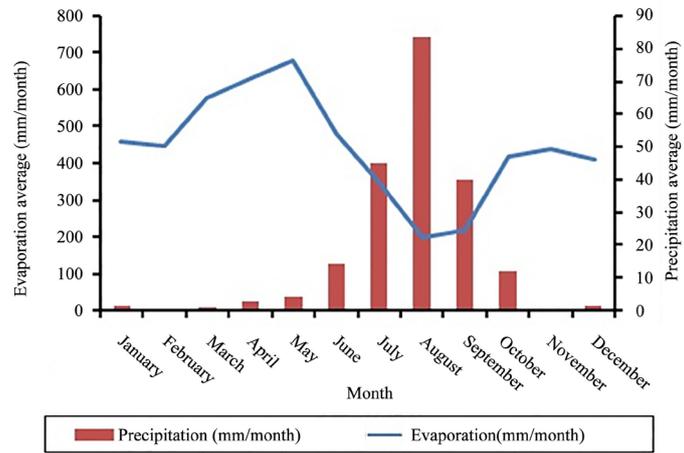
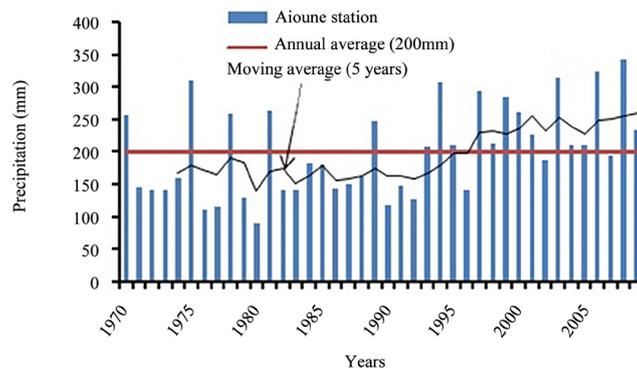


Figure 1. Geology, section cross AB and localization map of the Aïoune aquifer area.



(a)



(b)

Figure 2. (a) Pluviometric and evaporation average per month at aioune station (1970-2009); (b) Pluviometric variation average per year to the Aioune station (1970-2009).

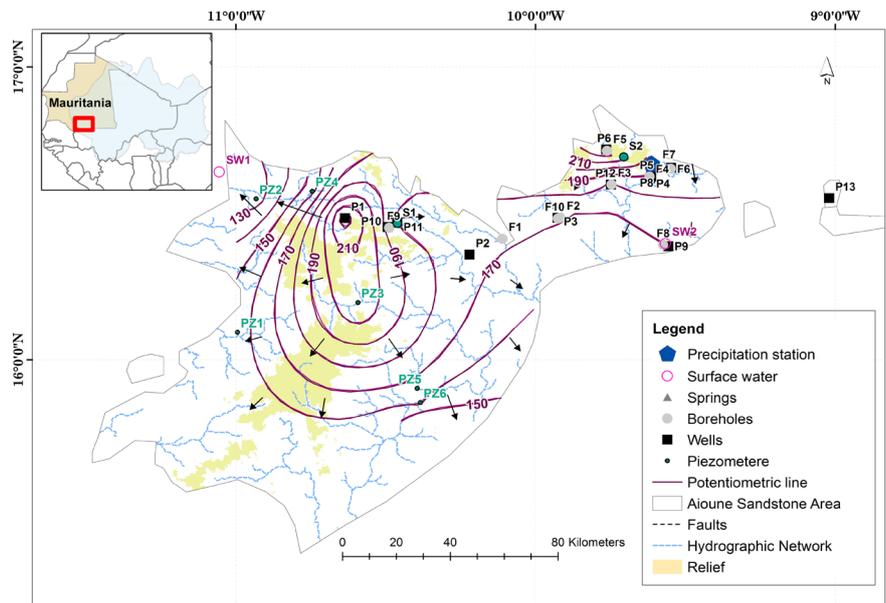


Figure 3. Piezometric head map of aioune groundwater (August 2013).

infracambrian series. This sedimentary series unconformably overlay the granitic basement of “Birrimien”. The infracambrian series begin at the base with basic compact clayey quartzite sandstones of 400 - 1000 m thickness [5], covered by ferruginous sandstones which extends a large area and have oblique stratification, called “Aioune sandstones”. This unit is between 80 and 120 m in thickness, and its oblique stratification characteristic could be due to strong marine currents during the past [6]. At the top, the infracambrian series occur with the sandstone compact quartzite “brain sandstone”. The thickness of the infracambrian series is in the order of 2000 m [7].

In the north, the previous series are covered in angular unconformity [7], by the lower Cambrian outcrop series, which begin at the base with coarse sandstones, interspersed by a granite cobble clay conglomerate of so-called glacial origin, namely Tillite. The thickness of this series in the west is 210 m [5]. To the West and East, the upper Cambrian outcrop is characterized by the presence of the Pelite, jasper and dolomitic limestone minerals.

The study area is characterized by a large extension of the unconsolidated and interstratified sandstone of Aioune [7]. This formation becomes aquifer in area where are favorable conditions occur [4], such as the primary and secondary permeabilities of the Aioune sandstone formation occur. The superficial part of this formation mostly exhibits these characteristics. The thickness of the one layer aquifer of “Grès d’Aioune” is variable, with an average of 15 m [8]. The static level varies between 2 and 35 m below ground surface. The average transmissivity is 6×10^{-4} m²/s. The porosity in the order of 5% and flow rates can reach 15 m³/h [8], and the renewable reserve is estimated at 2.1 million m³/year [7].

3. Methodology

A total of 23 groundwater samples, two springs water samples, and two surface water samples, were collected from the Aioune sandstone aquifer, in December 2011 and September 2013. In addition rainwater samples covering only the rainy events of August 2013 were collected during the course of this study. Physico-chemical parameters (pH, conductivity, temperature) were measured on site. The determination of the concentration of the major elements (28 samples) was carried out by high performance liquid chromatography (HPLC). The stable isotopes were analyzed by the laser spectrometer (28 samples). The tritium content was performed by liquid scintillator after electrolytic enrichment (28 samples). Only ten samples were analyzed for ¹⁴C and ¹³C. The radiocarbon analyses were performed by accelerator mass spectrometry (AMS). All analyses were carried out at the Laboratory of Radio-Analyses and Environment (LRAE) of ENIS of Sfax (Tunisia) except for ¹⁴C and ¹³C, carried out at the isotopic laboratory of the University of Groningen (Holland). The analytical errors on the measurements are $\pm 0.2\%$ for oxygen-18, $\pm 0.6\%$ for deuterium, for Tritium $\pm 0.3\%$ and carbon-13 (Table 1).

Only the chemical balance data of $\pm 5\%$ were considered.

Table 1. Piezometric, chemical and isotopics data of aioune groundwater (P: Wells, F: Boreholes, S, springs, PZ: Piezometer, SW: Surface Water and EP: precipitation water).

Code Number	Date	Piezometric Level		EC	T	pH	TDS	Cl ⁻ (mg.l ⁻¹)	NO ₃ ⁻ (mg.l ⁻¹)	SO ₄ ²⁻ (mg.l ⁻¹)	HO ₃ ⁻ (mg.l ⁻¹)	Na ⁺ (mg.l ⁻¹)	K ⁺ (mg.l ⁻¹)	Mg ²⁺ (mg.l ⁻¹)	Ca ²⁺ (mg.l ⁻¹)	d ¹⁸ O ₂ (‰ vs Smow)	d ² H (‰ vs Smow)	Tritium (TU)	C13 (‰ VPDB pcm)	C14 (‰ BP)	Age	Depth (m)	
		(m)	(m)																				
P1	18/12/2011	226.05		1477	30.9	7.4	859	71.82	341.80	73.53	213.50	35.67	7.08	72.23	93.45	-31.4	-3.4	0.4					40
15	16/08/2013			1609	31	7.34										-30.1	-4.0		-9.1	90.09	actuel		
P3	19/12/2011			3275	32.0	7.1	2810	294.3	1128	272.93	158.6	101.66	22.45	111.28	445.66	-26.1	-2.3	2.5					16
P4	20/12/2011	190.15		906	28.1	7.7	690	113.4	190.1	71.83	42.7	55.06	7.74	28.68	87.57	-32.7	-4.3	5.2					16
22	20/08/2013			1050	29.8	5.56										-27.5	-3.6		-12	104.37	actuel		
P8	20/12/2011			1670	26.7	7.5	1263	175.08	257.32	134.35	292.80	121.62	77.7	40.47	154.37	-28.3	-3.0	3.0					15.2
P9	20/12/2011			1003	27.2	7.5	653	82.14	65.24	90.43	274.50	80.76	16.74	31.10	64.51	-20.4	-2.0	2.1					12
P10	18/12/2011			906	30.5	8	610	51.21	152.60	26.99	219.60	66.64	25.2	25.57	57.53	-33.3	-4.4	1.7					20
F8	20/12/2011	170.37		1860	30.3	7.8	1140	175.22	6.32	361.21	286.70	279.76	0	40.03	67.80	-33.8	-4.5	0.2					25
30	19/08/2013			2030	37	7.16										-36.0	-4.6		-12	69.27	1062		
F10	19/12/2011	166.6		1325	30.9	7.4	819	153.6	27.83	267.65	103.7	141.47	5.61	33.65	82.65	-42.6	-5.3	0.4					75.5
35	19/08/2013			961	28.8	6.68										-41.3	-5.8		-9	59.75	171		
P2	18/12/2011			186	30.6	8	129	7.23	60.66	3.69	12.20	8.79	5.97	4.65	13.12	-31.4	-4.3	1.3					25.3
P5	20/12/2011			450	27.6	7.4	313	47.05	48.48	34.48	73.20	26.39	38.58	9.20	27.28	-30.3	-4.0	1.7					15
P6	20/12/2011			279	29.7	7	201	19.93	36.98	18.68	57.00	15.92	3.66	7.53	22.02	-38.2	-5.5	0.3					16
P7	21/12/2011			158	28.1	7.7	96	14.99	34.64	3.78	12.2	9.123	2.49	4.04	10.57	-40.2	-5.1	0.4					17
P11	18/12/2011	176.85		520	29.9	8	334	12.84	30.27	39.89	189.1	19.97	19.59	20.94	34.23	-36.2	-4.3	0.4					8.65
34	17/08/2013			1512	29.8	7.74										-32.6	-4.0		-9.9	94.49	actuel		
P12	19/12/2011	180.05		591	29.7	7.3	359	44.51	49.35	23.03	170.80	28.42	5.00	13.07	60.28	-38.9	-4.7	1.2					20.95
36	19/08/2013			600	30.8	7.28										-38.0	-4.8		-11	93.01	actuel		
P13	21/12/2011	175.05		422	29.4	7.5	303	9.10	16.93	7.03	213.50	23.59	2.97	13.88	44.28	-39.1	-5.6	0.5					30

The rainfall samples collected at the Aioune station are representative ($((^2\text{H}-(^{18}\text{O} * 8))-10): 0.7$), located on the global meteorological water line, and are in the range of stable isotope values recorded at the nearby GNIP station (Bamako (Mali) and Niamey (Niger) (**Table 2**).

The correct Bamako GNIP data (low deviation from mean ($(^2\text{H}-(^{18}\text{O} * 8))-10$)), and extrapolated from the Vienna GNIP series, were used in tritium tracing.

4. Results and Discussions

4.1. Piezometry

Head data measured from 16 piezometers were used to establish piezometric map of Aioune groundwater. The piezometric surface morphology of the aquifer follows the soil topography, characterizing a low depth water table (water table). Head values range between 238 and 143 m of altitude and the flow direction is generally from the cliff towards the plains (**Figure 4**). The hydraulic gradient calculated for this water table varies between 1 and 5‰, the effective velocity is 94 m/year (**Figure 3**).

The data collected in 2013, at the National Hydraulic Database (SIPPE2) [9], were additionally used and extrapolated to evaluate the volume taken in 2006, which is estimated at 18,000 m³/d (represents 6.5 million m³/year).

4.2. Hydrochemical Characteristics

4.2.1. Surface Water and Rainfall

In December 2011 and August 2013, two samples of surface water were collected from the study area, and all the August 2011 rain events greater than 5 mm, at the Aioune station. The average temperatures are 23°C and 14°C, respectively for surface water and rainwater. The average pH is 8.2 and 7.1, respectively for surface water and rain. Salinity ranges from 318 to 755 mg/l for surface water and 18 mg/l for rainwater. The surface water and rain samples exhibit HCO₃-Ca water type.

4.2.2. Groundwater

1) Salinity and chemical facies

The chemical and physicochemical data from the 2011 campaign samples were assessed to evaluate the chemical quality of the water in the study area and identify the geochemical processes controlling mineralization of the groundwater.

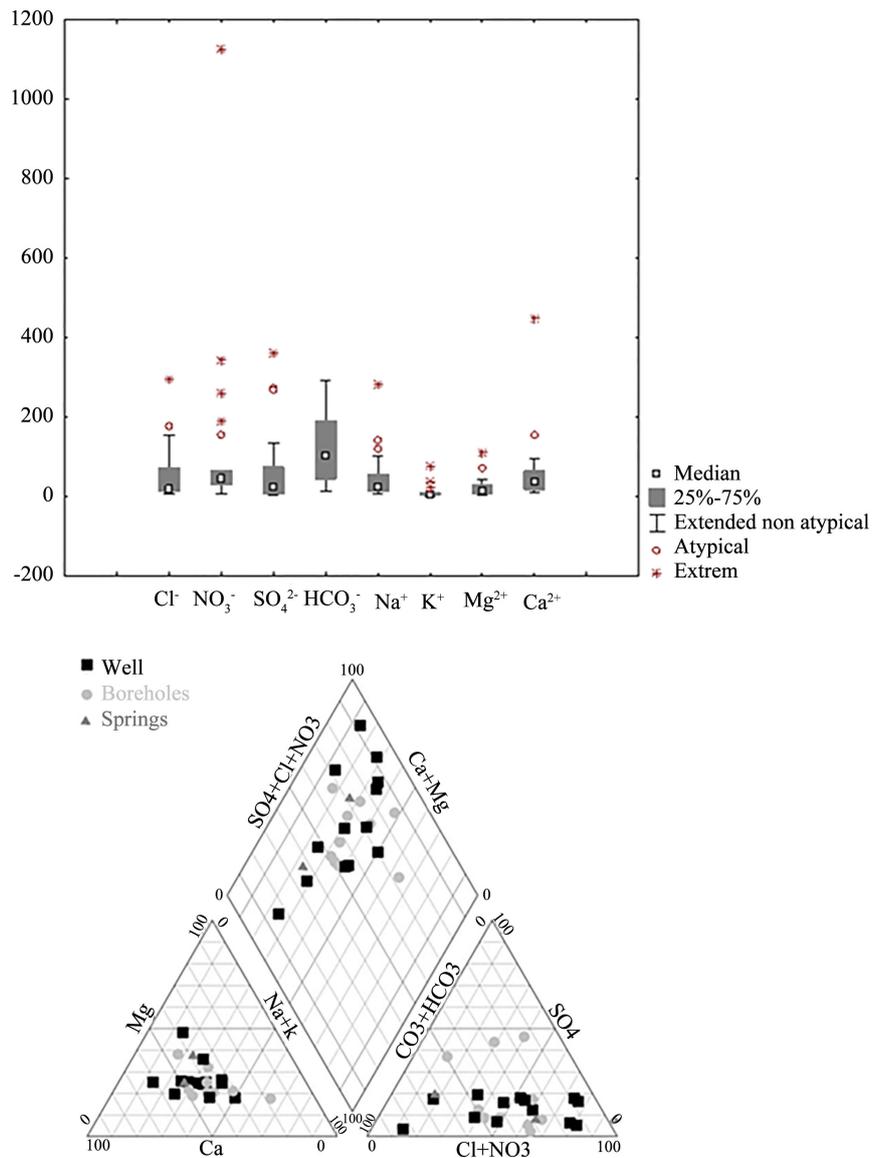
Indeed, depending on the sampling points, the waters were classified into three groups: wells, boreholes and springs (**Figure 1**).

The temperatures range from 33.1°C (Borehole F7) to 27.2°C (P8 Well), with average 30°C, and show some increase in depth, except for El Ghouds borehole (F6).

The pH varies between 6.5 and 8.6, with an average of 7.5. The electric

Table 2. Stable isotopes of precipitation in the study area and near regional stations sample name.

	$\delta^2\text{H}$ (‰ vs Smow)	$^2\text{H_ERR}$	$\delta^{18}\text{O}$ (‰ vs Smow)	$^{18}\text{O_ERR}$	Date	Latitude	Longitude	Altitude	Type	Prof	Locality
EP1	-36.3	0.9	-5.7	0.1	31 au 18/08/2013	16.6102	-7.281	230	Rainfall	0	Aioune
Bamako	-45.40		-6.58	-		12.6932	-7.995	381	Rainfall	0	Bamako
Niamey	-40.10		-5.78	-		13.5200	2.090	220	Rain	0	Niamey

**Figure 4.** (a) Major elements of the Aioune sandstone water; (b) Piper diagram of the analysed Aioune groundwater samples.

conductivity varies from 140 $\mu\text{s}/\text{cm}$ to 3200 $\mu\text{s}/\text{cm}$, and the average is 451 $\mu\text{s}/\text{cm}$. The spatial distribution of the electric conductivity shows higher values in the shallow depth of the alluvial spreading zones and in particular near the major

agglomerations (eg Gounguel, Aioune, etc.).

The general order of abundance of the cations in meq is $\text{Ca}^{2+} > \text{Na}^+ > \text{Mg}^{2+} > \text{K}^+$, with the exception of few water points where $\text{Na} > \text{Ca}$. For anions, the general order is $\text{HCO}_3^- > \text{NO}_3^- > \text{Cl}^- > \text{SO}_4^{2-}$, with high concentrations of Nitrates, reaching 1.1 g/l, mainly in the shallow level, reflects the vulnerability of the water table to the pollution (**Figure 4(a)**).

The representation of the chemical data in the Piper diagram shows waters whose chemical facies evolve mainly between two poles, a HCO_3 -Ca (low mineralized) pole, and a Cl-Ca/Na (generally more mineralized) pole. Some other facies types occur such as, HCO_3 -Ca/Na, HCO_3 -Na/Ca, HCO_3 -Na and SO_4 -Na, reflecting concomittent other chemical processes (**Figure 4(b)**).

2) Geochemical processes

In order to determine the origin of the mineralization of the Aioune sandstone groundwaters samples, some correlations between the main major elements have been studied, notably with Cl which is considered as conservative.

The positive correlation between Ca/Cl vs. Cl (**Figure 5(a)**) generally indicates the contribution of this element to mineralization, and similarly for Na/Cl vs. Cl (**Figure 5(c)**), indicates the significance of the contribution of these elements. These elements generally result from the dissolution of carbonates and evaporites and the basic exchange phenomenon. The K and Mg elements do not show a clear trend (**Figure 5(b)** and **Figure 5(d)**).

The positive correlation between Ca, Mg, HCO_3 and SO_4 (**Figures 5(e)** and **Figures 5(f)**), probably reflects carbonates and gypsum dissolution phenomenon. **Figure 5(g)** shows a good correlation between Na, K and HCO_3 , up to high values ($3 \text{ meq}\cdot\text{l}^{-1}$) of HCO_3 , where there is a stability of the HCO_3 concentration, the increase in Na + K, can be explained by the achievement of the equilibrium of water with respect to the carbonates, therefore the presence of the phenomenon of precipitation-dissolution.

Figure 5(h) indicates the presence of both types of cationic exchange with type I clays through the release in water of Ca^{2+} and or Mg^{2+} ions against fixation Na^+ ion (P 3), and type II, where the Na^+ ions of the complex are released and are replaced by the Ca^{2+} (F8) ions [2]. This cationic exchange mainly affects the mineralized waters, which are close to the alluvial spreading (Gounguel Zone, Douerare etc.).

Saturation index (SI) vs. major elements (**Figure 6**) show waters mostly saturated to supersaturated in carbonates, which explain by the presence of the phenomena of precipitation/dissolution of carbonates (**Figures 6(a)-(c)**). The waters are generally under-saturated in evaporite, which explains by the presence of the phenomenon of dissolution of evaporites (**Figures 6(d)-(f)**) [10].

The high exposure of the aquifer to contamination is probably due to variation in the unsaturated zone thickness (<30 m), characterizing a water shallow.

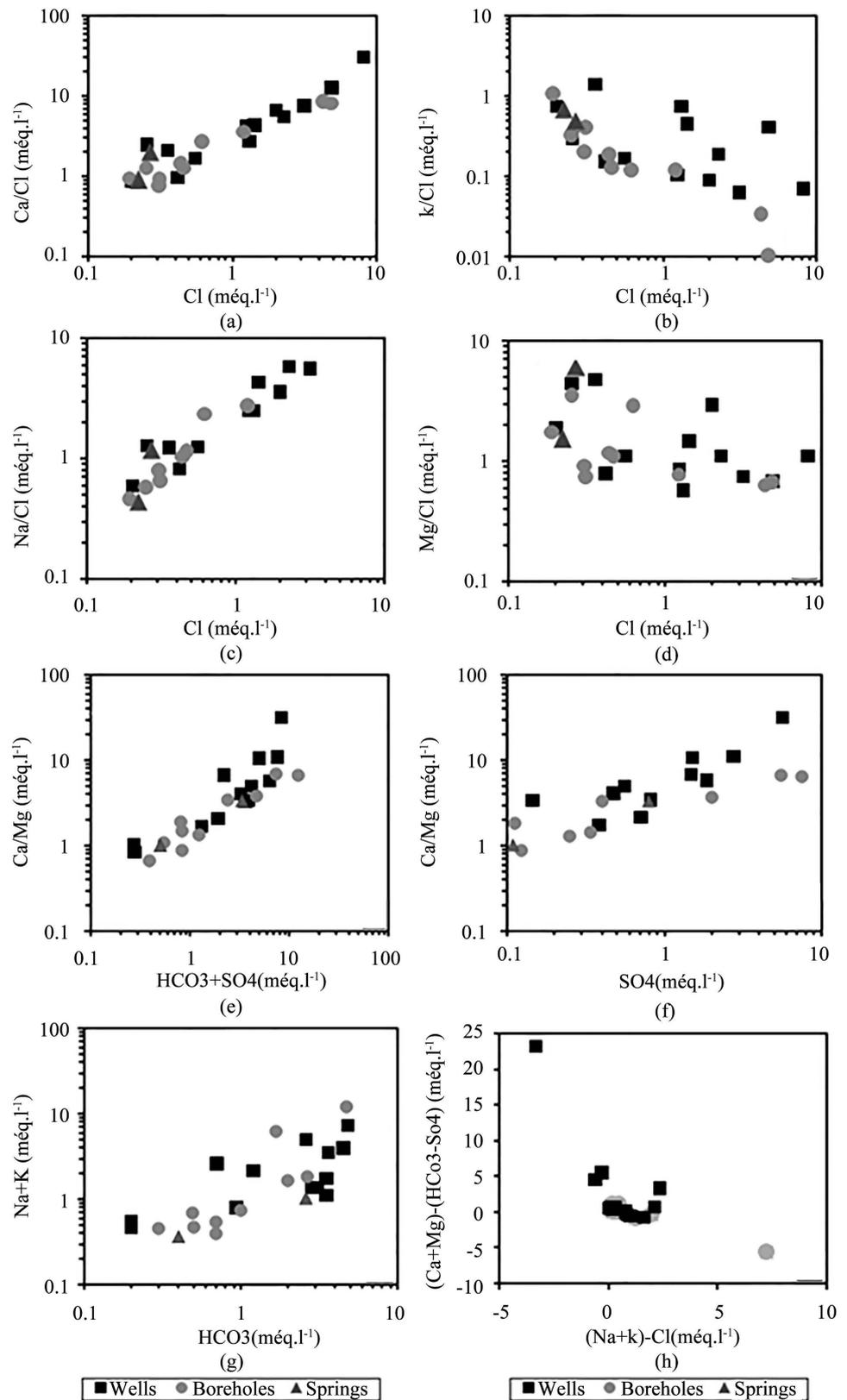


Figure 5. (a) Cl vs. Ca/Cl; (b) Cl vs. K/Cl; (c) Cl vs. Na/Cl; (d) Cl vs. Mg/Cl; (e) $(\text{HCO}_3 + \text{SO}_4)$ vs. $(\text{Ca} + \text{Mg})$; (f) SO_4 vs. $(\text{Ca} + \text{Mg})$; (g) HCO_3 vs. $(\text{Na} + \text{K})$; (h) $((\text{Na} + \text{K}) - \text{Cl})$ vs. $((\text{Ca} + \text{Mg}) - \text{HCO}_3 + \text{SO}_4)$.

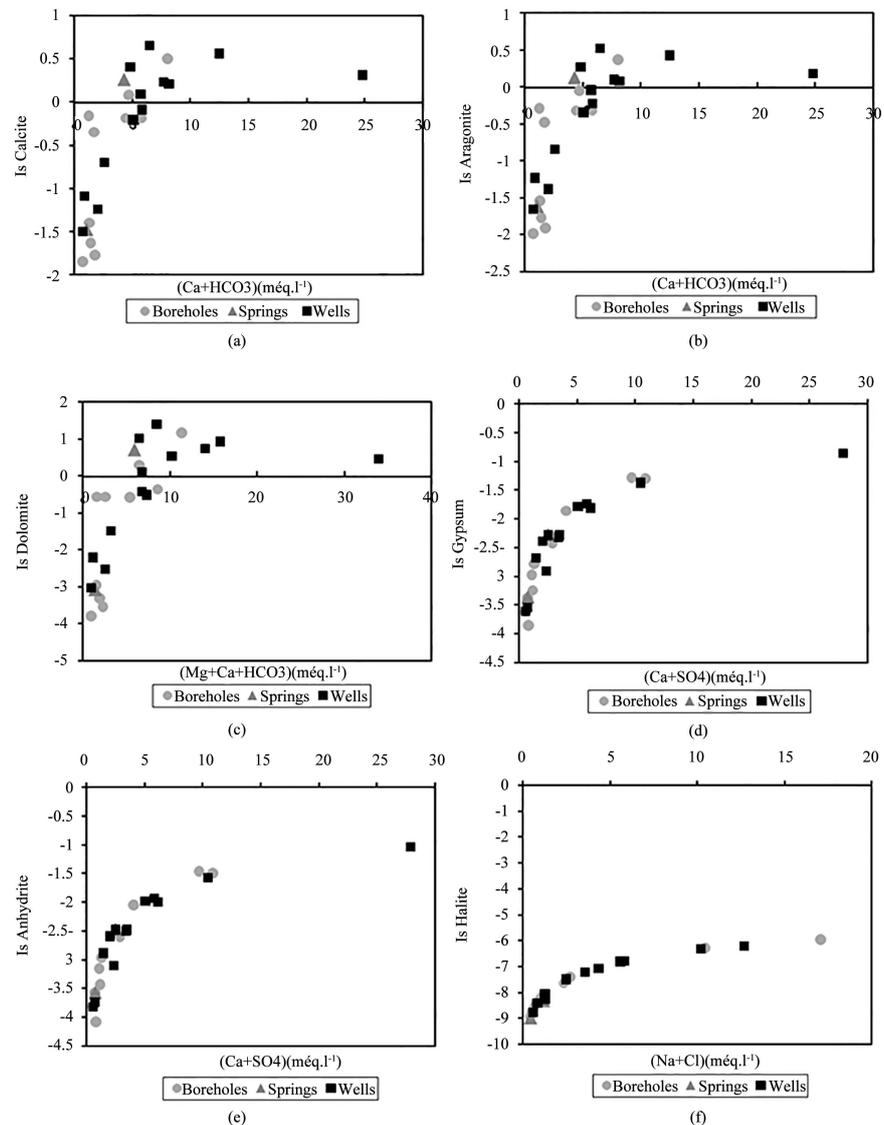


Figure 6. Correlation between the saturation indices and major elements of the Aioune sandstone water.

4.3. Isotopic Characterization

4.3.1. Surface Water and Rainfall

For rainwater, the mean value content are -5.7‰ and -36.3‰ V-SMOW, for $\delta^{18}\text{O}$ and $\delta^2\text{H}$ respectively, and are aligned on the Meteoric Water Line (MWL). For surface waters, the $\delta^{18}\text{O}$ contents vary between 23.3 and 29.7‰ V-SMOW, with an average of 26.5‰ and for the $\delta^2\text{H}$, values range between 99.6 to 130.2‰ V-SMOW, with an average of 114.9‰. Plotted in the conventional $\delta^{18}\text{O}$ vs $\delta^2\text{H}$, data scatter and deviate the meteoric line evidencing evaporation phenomenon and probably mixing.

4.3.2. Groundwaters

1) Stable Isotopes and Tritium

The $\delta^{18}\text{O}$ contents vary between -5.7 and -2.0‰ V-SMOW, with an average

of -4.6‰ and for $\delta^2\text{H}$, values range between -42.6 to -20.4‰ V-SMOW, with an average of -35.3‰ . This large range probably reflects the multiple phenomena present in the system.

The $\delta^{18}\text{O}/\delta^2\text{H}$ correlation characterizes two groups, Group A, with indirect evaporated recharge, occurring mainly at the wells and springs. Secondly, Group B, reflects mixing water between current and older recharge, located mainly at the boreholes (Figure 7). A significant variation of the stable isotopes of water, is recorded between 2011 and 2013, that characterize a water table with a fast response time, and also reflect the state of a superficial aquifer.

The $\delta^{18}\text{O}$ -Depth Correlation shows a water table affected by evaporation mainly at shallow depths between 0 - 30 m.

Correlation Cl- $\delta^{18}\text{O}$ identifies the presence of the phenomenon of dissolution, and to a less a extent evaporation phenomenon. This latter is located mainly at the shallow level.

Tritium values vary between 5.2 TU and 0.1 TU, with an average of 1.2 TU. The spatial variation of the tritium content shows the presence of local recharge, localized at the spreading zones, probably due to the topographic and favorable lithological conditions to the convergence and the water infiltration (Figure 8).

The average tritium contents of 1.5, 1 and 0.7 TU, respectively for wells, springs and boreholes, characterize the stratification of the aquifer. For some

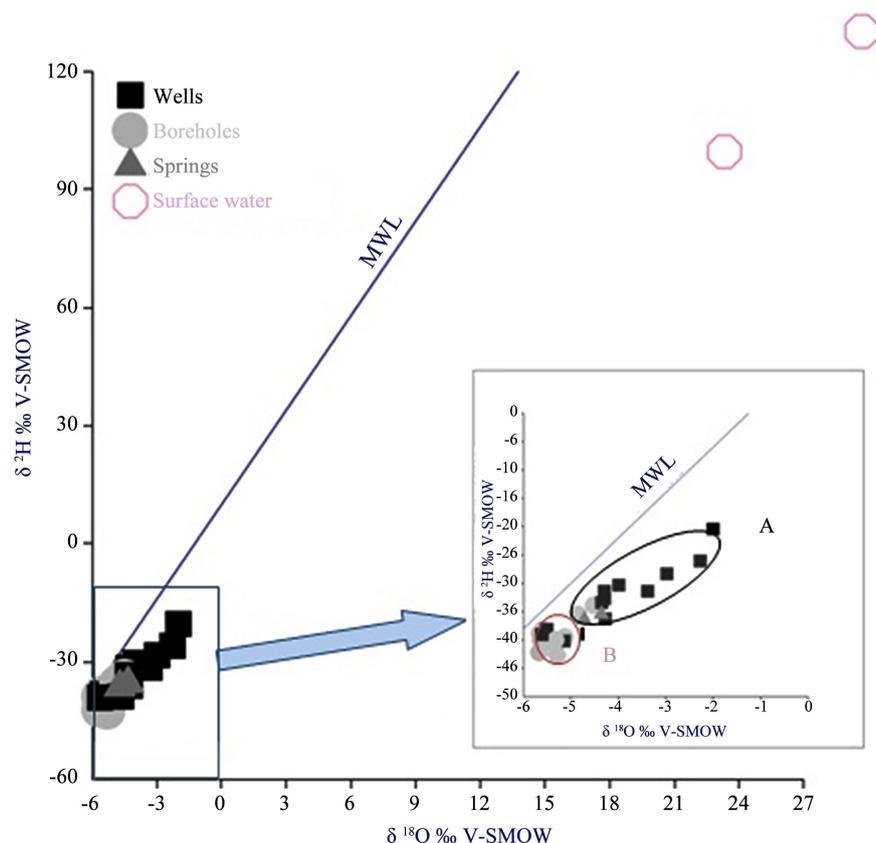


Figure 7. ^2H versus ^{18}O of Aioune Sandstone waters.

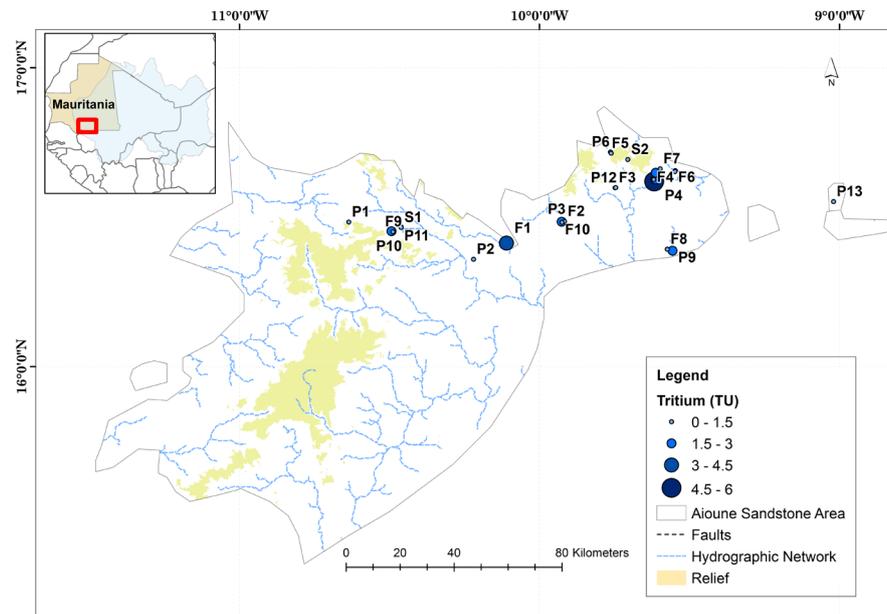


Figure 8. Tritium map of aioune aquifer (December 2011).

wells the recharge occurred after 1961 [11], while all the boreholes and springs were recharged before 1961, except the F1 borehole, which is located at Tintane City, which suffered very strong inundations in 2007, causing the evacuation of the city (Figure 9).

2) Carbon Isotopes

Correlation ^{13}C - ^{14}C exhibits large variation in ^{13}C , from -11.9 to -7.4 , indicating the presence of several carbon sources (the atmospheric reservoir and the carbonate matrix). The Aioune sandstone aquifer is characterized by a deep water enriched in $\delta^{13}\text{C}$, identifying an exchange with the carbonate matrix [2]; and surface water depleted in $\delta^{13}\text{C}$ (Figure 10), indicating the presence of isotopic exchange with soil CO_2 or organic matter [2].

The ^{13}C equilibrium, representative of soil CO_2 , defines an open to intermediate system aquifer.

For the Aioune sandstone aquifer, activities in carbon 14 vary from 60 to 104 pcm, and an average of 84 pcm, confirming the contribution of recent waters to recharge.

3) Groundwater Dating

The variation in radiocarbon contents makes it difficult to identify the initial activity A_0 . Several models were used to correct this activity taking into account the different reactions. These models evolve from the simplest (model dilution), to the more complex (exchange isotopic).

In the presence of isotopic exchange between the water and the two carbon tanks, identified previously, the F and G, IAEA and Mook models are the most appropriate since they give close values; but in the presence of an open to intermediate system, the exchange phenomena with the gaseous phase and matrix, the presence of a calcareous medium [12], the IAEA model is most indicated,

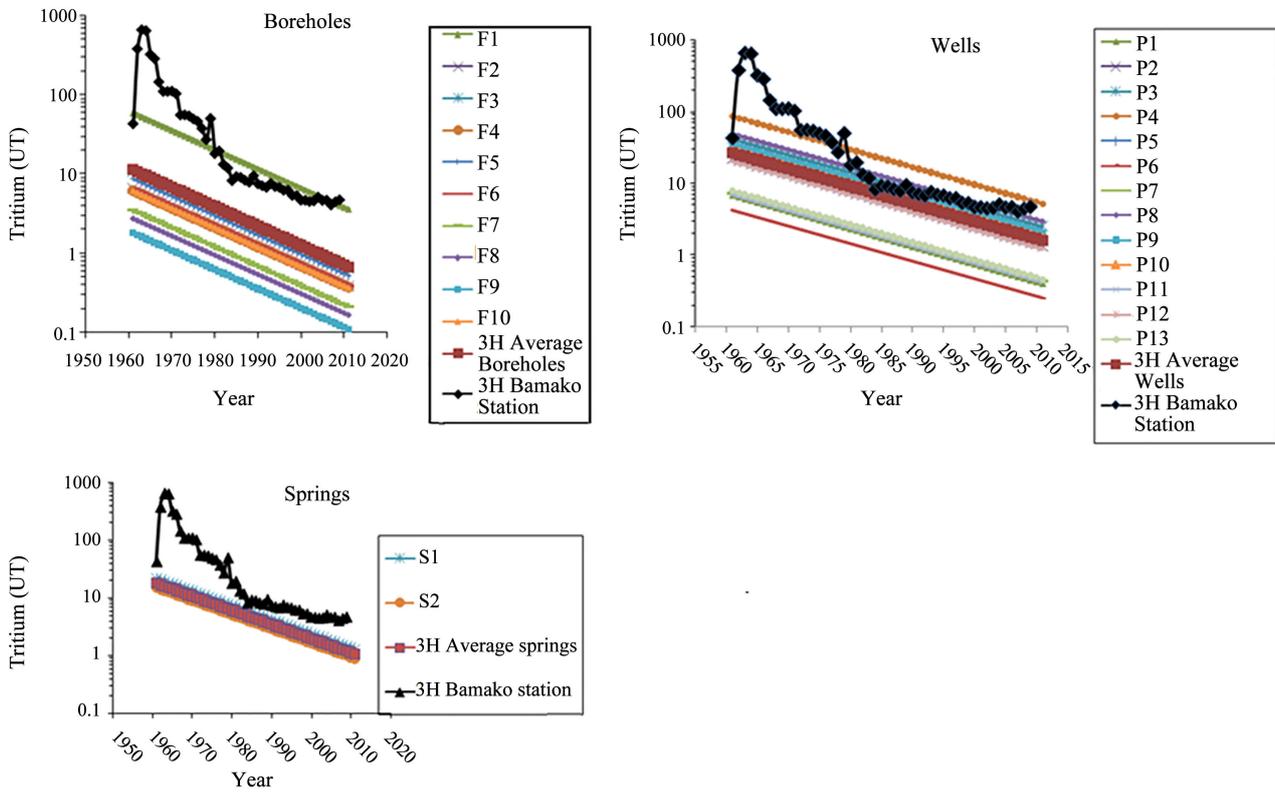


Figure 9. Datation Tritium of Aioune sandstone groundwater (2011) (pluviometric station Bamako(Mali)).

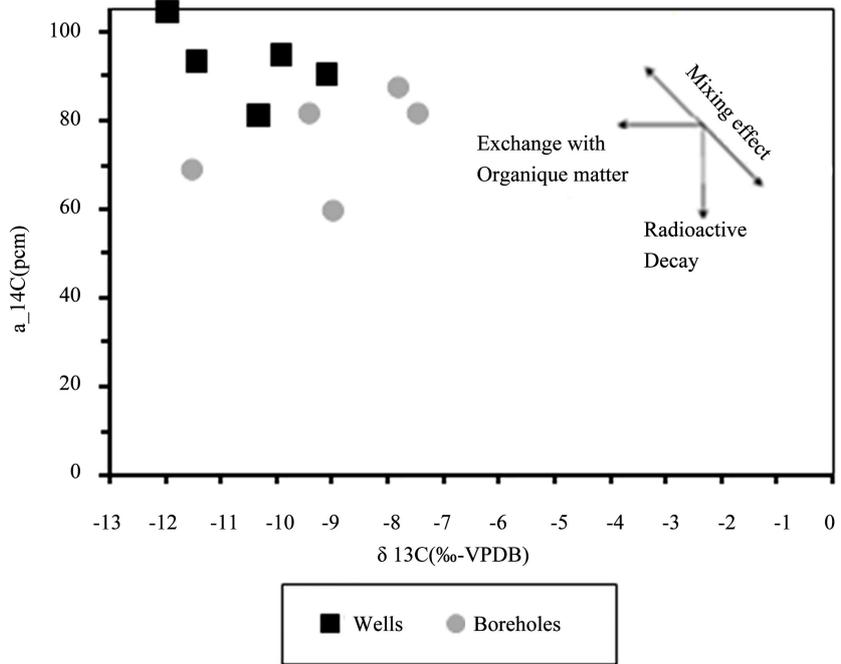


Figure 10. Relationship between activity carbone 14—carbone 13.

and gives an age closer to the piezometric age.

The age ranges from the present to 1060 years, which is in accordance with

the Tritium data, which characterize recent recharge (wells) to mixing water (boreholes) (Table 3).

4.4. Recharge

Quantification of the recharge was discussed using the chloride masse balance and the carbon 14 method.

The chloride masse balance consists of using the Allison and Hughes formula, 1978, for rainfall events greater than 5 mm sampled at the study area and groundwater samples from the recharge zones (at $TU > 1$).

$$P \times Cl_{\text{rainfall}} = R \times Cl_{\text{aquifer}}$$

where P : the average height of rain in $\text{mm}\cdot\text{year}^{-1}$, R : the recharge in $\text{mm}\cdot\text{year}^{-1}$, Cl_{rain} : the chloride content in precipitation in $\text{mg}\cdot\text{L}^{-1}$, Cl_{aq} : the content of chloride in the sample in $\text{mg}\cdot\text{L}^{-1}$.

The ^{14}C method uses the formula adapted to free groundwater [13], which takes into account the depth of structures and the thickness collected.

$$R: H\lambda\theta \cdot \ln\left(\frac{H}{H-Z} - \frac{A_0}{A}\right)$$

with: R : Recharge rate, H : Total aquifer thickness, Z : Depth of the strained area, θ : Porosity, λ : Radioactive decay constant, A : Activity measured in carbon-14, A_0 : Activity initial carbon-14.

The computed average recharges are 0.4 and 5 mm/year, respectively for the ^{14}C method and the chloride masse balance, which represents respective volumes of 0.4 to 5 million m^3/year .

The volume of water extracted from 494 water points (175 Wells and 319 boreholes), from the Aioune sandstone aquifer, is in the order of 6.5 million m^3/year by considering a rate of 1 m^3/hr for each well and 5 m^3/hr for borehole [9].

Table 3. ^{14}C Age of Aioune groundwater.

Code	brut	Tamers	Pearson	Mook	F & G	F & G _{eq}	AIEA	Evans	Eichinger	Olive
P1	863	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel
F3	1661	actuel	actuel	actuel	actuel	83	actuel	actuel	actuel	actuel
P4	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel
F5	1071	actuel	actuel	actuel	actuel	1501	actuel	actuel	actuel	actuel
F6	1683	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel
F8	3035	actuel	actuel	1362	actuel	1813	1062	actuel	actuel	actuel
P11	469	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel	actuel
F10	4258	actuel	actuel	793	actuel	2409	171	actuel	actuel	actuel
P12	599	actuel	actuel	actuel	actuel	806	actuel	actuel	actuel	actuel
P13	1747	actuel	actuel	actuel	actuel	539	actuel	actuel	actuel	actuel

The comparison between the annual recharge and the annual pumping rate, evidences an imbalance in the water budget, which results in the recurrence of water shortages during summer, and makes it urgent to take appropriate measures to make sustainable and rational this water resources.

5. Conclusions

This pioneering work in the Aioune sandstone one layer aquifer, located in the southern part of the Taoudenni Basin in Mauritania, is based on the interpretation of hydrodynamic (piezometry, porosity, etc.), hydrochemical (elements major) and isotopic ($\delta^{18}\text{O}$, $\delta^2\text{H}$, $\delta^3\text{H}$, $\delta^{13}\text{C}$ and $\delta^{14}\text{C}$) data to better understand the behavior of this aquifer.

Piezometric configuration of the water table reveals effective velocity of 94 m/year and direction flow generally from the cliff to the plains, spreading areas. The chemical data have identified a layer whose chemical facies evolving between two poles, a calcium-magnesium bicarbonated pole (low mineralized), and a second pole chloride calcium-sodium. However, the shallow water table in part of aquifer reveals high vulnerability to pollution, especially at the shallow horizons (Douerare, Aioune). The dissolution of carbonate and evaporitic minerals is likely the main process controlling mineralization, and to a lesser extent cation exchange.

The isotopic tool allowed to identify waters classified in two groups. Group A, with indirect evaporated moderne recharge is located at the shallow level of the spraying areas; and Group B consists of mixing water, between old and modern recharge water, located deep within the water table. The sandstone Aioune aquifer is characterized by an open to intermediate system, whose water age ranges from the present day to 1060 years.

The calculated average recharge is 0.4 and 5 mm/year, respectively for the ^{14}C method and the chloride masse balance.

The ^{14}C method takes into account the depth of structures and the thickness collected, and the chloride masse balance consists of using the rainfall events greater than 5 mm and groundwater samples from the recharge area (at $\text{TU} > 1$).

Recharge compared to water pumping characterizes a water table at the limit of imbalance. In order to be able to provide a sustainable solution, it is necessary to exploit these results to arrive at a model of the water table.

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