

Geochemical Soil Survey for Base and Precious Metals in Dagbala-Atte District, Southwestern Nigeria

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Abstract

Geochemical survey of residual soils was undertaken in Dagbala-Atte District of Igarra Schist Belt, southwestern Nigeria to explore for base and precious metals (Cu, Pb, Zn, Au and Ag). It involved sampling residual soil from 49 sites in a grid pattern; analyzing the soil samples for the base and precious metals and other commonly associated elements including As, Cd, Fe, Hg and Sb; subjecting the geochemical data generated to statistical analysis; and preparing the geochemical distribution map of the district for each element. On the basis of the multivariate statistical analysis (correlation matrix and factor analysis) of the soil geochemical data, three mineralization types were inferred to be present in the district. The suspected types of mineralization consist of: 1) Cu-Pb-Zn-Ag mineralization with associated As, Hg, Sb and Fe; 2) Au-Cu mineralization with associated As; and 3) minor Hg-bearing Au-Pb mineralization. With the aid of the elemental geochemical distribution maps, the suspected mineralization was linked to rock types from which the sampled residual soils were apparently derived. Hence, the suspected Cu-Pb-Zn-Ag mineralization is linked to the quartz-biotite schist at the southwestern part of the district while the suspected Au-Cu mineralization is located on the silicified sheared rock at the northern and southern parts and the minor Hg-bearing Au-Pb mineralization within the granitic gneiss at the northeastern and eastern areas of the district. Based on the foregoing findings, it is recommended that lithogeochemical survey for the base and precious metals should be done in the area.

Keywords

Dagbala-Atte District, Factor Analysis, Geochemical Distribution Map, Mineralization

1. Introduction

On the basis of high potential for metallic mineralization of the Paleoproterozoic Schist Belts in northern Nigeria [1] [2] [3] [4] which are similar to the Igarra Schist Belt in southern Nigeria, the Postgraduate Mineral Exploration Research Group of the Applied Geology Department, the Federal University of Technology, Akure initiated a geochemical exploration program of the Igarra Schist Belt area. The program commenced with regional stream sediment survey which detected anomalous concentrations of base and precious metals among others around Dagbala and Atte areas [5] [6] [7]. The next step in the planned geochemical survey of the belt is to follow up these stream sediment anomalies by tracing them to their possible land sources particularly because environmental contamination around Dagbala and Atte has been ruled out based on pollution studies carried out by [8] [9] in the area. The follow-up of the stream sediment anomalies is based on the notion that the stream sediments are derived from the soil by processes of erosion and deposition in the stream beds. Hence, the present aspect of the research program is to conduct a soil geochemical survey of the Dagbala-Atte area for base and precious metals which showed up in anomalous values in the stream sediments of the area [6] [7]. Similar surveys using residual soil as a sampling medium have been carried out by other workers particularly because the results of such surveys are much easier to interpret in terms of the source of the elements [10]-[15]. The principal objective of this study therefore is to locate the possible source(s) of the metals (Cu, Pb, Zn, Au and Ag) in the residual soils of the study area.

2. Geological Setting of the Study Area

Dagbala-Atte District lies within latitude 7°10'N and 7°21'N and longitude 6°09'E and 6°17'E in the northern part of Edo State, Nigeria (Figure 1 and Figure 2) and covers an area of about 285 km² within the Igarra Schist Belt, which is surrounded by the older Basement Complex rocks believed to be of Archean to Paleoproterozoic age [16] [17]. Low-grade metasediments commonly described as the Younger metasediments [18] or Newermetasediments [19] dominate the schist belt. These metasediments consist essentially of a deformed package of pelitic to semi-peliticschists, marbles, calc-silicate gneiss, quartzites and polymic-ticmetaconglomerates [20] [21] [22] [23]. Both the Younger metasediments and the older basement rocks (*i.e.* the gneisses and migmatites) were intruded by the lgarra batholith. Minor felsic and mafic intrusives, including pegmatites, aplite, syenite, lamprophyre and dolerite cross-cut the Pan African granites and the pre-existing rocks.

In the Dagbala-Atte District the rock suites are divisible into the eastern portion dominated by granite gneiss and the western part composed mainly of the Younger metasediments with both portions separated by a narrow zone of silicified, sheared rock (**Figure 3**) [24]. The metasedimentary package consists



Figure 1. Location map of the study area. Insets: map of Nigeria showing Edo state and map of Edo state showing the location of Dagbala-Atte district.

predominantly of quartz-biotite, garnet-biotite and mica schists with minor metaconglomerate, quartzites and marble. Both the metasediments and granite gneiss are intruded by porphyritic Pan African granite. All the rock types are affected by intense weathering to varying degrees under the prevailing humid tropical climatic conditions producing ubiquitous residual soil profiles.

3. Methodology

The study entailed three stages as follows: 1) sampling of B-horizon soils from the Dagbala-Atte District of 285 km²; 2) geochemical analysis of the soil samples so as to determine the contents of the metals under consideration; and 3) data analysis, which involved qualitative and quantitative analyses.

3.1. Soil Sampling

The study area was first gridded into approximately 2×3 km portions on a topographic map of the area. This was followed by cutting traverse lines on the field to demarcate the grids (Figure 2 and Figure 3). Soil sampling was done on each grid by driving a 1-meter auger some 20 - 30 cm into the ground after



Figure 2. Topographic map of Dagbala-Atte district showing the localities of forty nine soil samples employed for this study.



Figure 3. Geological map of Dagbala-Atte district showing the locations of the forty-nine soil samples in a grid pattern (Adapted from [24]). removing the sandy top soil at every site marked by a labeled stake to ensure adequate penetration into the B-horizon which is preferred in most soil surveys owing to its high accumulation of trace elements [25]. The samples were carefully hammered out of the auger into pre-labeled baft cloth bags. Appropriate care was taken to avoid contamination during the sampling operation by cleaning the auger with water and ethanol after occupying each grid. A total of 49 samples were collected and transported to the laboratory. A global positioning system (GPS) receiver was employed for accurate sample site location on the base map. A particular attention was paid to the soil profile being sampled because the target was to sample residual soil, *i.e. in situ* soil profile, while transported soil was avoided.

3.2. Geochemical Analysis of the Soil Samples

The soil samples were first air dried at room temperature for four weeks, after which they were disaggregated prior to sieving with stainless steel screen of 200 microns size. The fraction passing through the sieve size was then pulverized to less than 75 microns using a porcelain mortar and pestle. All these were done with special care to avoid cross contamination. After the foregoing sample preparation in Nigeria, the soil samples were analyzed at Acme Laboratories, Vancouver, Canada using the following procedures.

Half of a gram (0.5 g) of each of the pulverized soil samples was weighed into graduated test tubes and 5 ml of 1:1:1 mixture of HCl-HNO₃-H₂O (modified aqua regia) was added to the samples. The mixture was heated at 95°C for one hour. After the mixture for each sample had cooled down, it was filtered and the leached solution obtained was diluted with ultra-pure water to 10 ml. The resulting solutions for all the samples were subjected to elemental analysis using an Inductively-Coupled Plasma Mass Spectrophotometer (ICP-MS) at the Acme Laboratories, Canada. The analytical results of eleven elements of interest, namely, Ag, As, Au, Cd, Cu, Fe, Hg, Mn, Pb, Sb and Zn, are presented in Table 1.

3.3. Data Analysis

The analytical results of the elements were subjected to quantitative statistical analysis and qualitative treatment. The quantitative statistical analysis comprised some univariate and multivariate statistics while the qualitative treatment involved plotting geochemical distribution map for the elements. Univariate statistics used here include histograms and box plots of both raw and log-transformed data. These plots are often useful when grouped together as they provide different ways of summarizing data and are necessary to reasonably establish the threshold values leading to the isolation of anomalous concentrations of the elements in the soils of the study area. The multivariate statistics in this paper consisted of correlation matrix and factor analysis.

3.3.1. Statistical Analysis of Geochemical Data

Quantitative statistical analysis employed in this study is widely used as a useful

HMS	Longitude	Latitude	Ag	As	Au	Cd	Cu	Fe	Hg	Mn	Pb	Sb	Zn
S01	6.27601	7.33925	12	0.6	1.2	0.01	17.31	1.9	37	80	9.41	0.03	13.4
S02	6.25559	7.34263	4	0.7	1.3	0.01	4.72	0.87	18	480	8.99	0.04	8.6
S03	6.22703	7.34360	6	0.9	1.7	0.01	11.81	2.64	23	381	14.89	0.03	18.6
S04	6.20553	7.33612	10	5.1	4.8		10.65	2.64	32	263	10.05	0.06	14.5
S05	6.20233	7.33403	3	0.7	2.3		5.15	0.83	9	120	4.67	0.02	8
S06	6.17917	7.34610	7	1.2	2.1		9.97	2.12	26	158	8.93	0.04	15.8
S07	6.15307	7.33607		0.6	8.3	0.01	3.21	0.7	15	107	6.24	0.03	11.1
S08	6.15820	7.30553	6	0.3	1.7		1.43	0.37	7	230	4.35		2.2
S09	6.17338	7.31172		0.3	0.9		1.32	0.12	10	16	1.99		2.3
S10	6.19797	7.31070	5	0.7	5.6	0.03	21.66	2.83	10	296	10.47	0.02	56.5
S11	6.21667	7.31667	4	214.5	12.5		78.9	4.43	15	21	10.74	0.11	31
S12	6.23509	7.31228	16	0.8	0.5	0.06	7.27	1.12	109	673	19.23	0.02	16.1
S13	6.25047	7.31230	9	0.6	0.5	0.05	11.8	1.46	22	279	8.21	0.04	13.8
S14	6.26623	7.30491	19	0.5	1.3		5.69	1.49	24	116	8.44		25.3
S15	6.27367	7.29119	58	1.2	9.8	0.05	11.8	3.6	26	459	24.99	0.05	68
S16	6.25689	7.28561	23	1.9	3.5		15.38	5.49	44	427	19.52	0.05	36.1
S17	6.22911	7.28183	6	0.6	2.7	0.01	32.05	2.55	10	315	2.65		41.5
S18	6.20923	7.28452	23	2.4	8.3	0.03	14.38	2.1	19	312	7.1	0.03	40.7
S19	6.19558	7.29703	27	3.5	1.8	0.02	14.86	2.21	7	356	8.12	0.06	46.3
S20	6.17358	7.28531	13	0.3	0.5		3.57	0.35		37	4.35		6.8
S21	6.15455	7.29648	18	0.5	0.7	0.03	10.91	1.17	10	386	8.23	0.05	29.8
S22	6.14990	7.25602	15	0.5		0.02	4.28	0.66	7	210	4.4		6.6
S23	6.17286	7.24761	22	1.9	1.2	0.02	24.85	3.34	15	502	13.6	0.05	75.2
S24	6.19840	7.24875	29	0.7	2.2	0.02	10.86	1.8	28	402	8.53	0.08	29.6
S25	6.21308	7.25647	5	0.1	0.7		3	0.52	13	264	5.7		4.5
S26	6.22980	7.26231	16	0.7			20.16	2.5	23	496	9.01	0.03	36.9
S27	6.25521	7.26753	14	0.8	0.6	0.06	10.87	2.16	34	893	16.5	0.05	59
S28	6.27683	7.26634	28	0.9	2.6	0.03	10.8	8.12	35	983	20.41	0.02	161.5
S29	6.27823	7.21795	14	2			12.05	3.97	14	374	8.94	0.38	164.1
S30	6.25292	7.22567	26	0.8	6.6		12.9	2.2	29	268	14.88	0.03	61
S31	6.23168	7.23347	19	1.3	3.8	0.02	14.43	2.72	32	286	16.03	0.05	33.4
S32	6.21665	7.23692	6	0.5	1.4		24.98	3.43	13	288	11.9	0.02	45.6
\$33	6.19598	7.22800	19	0.9	1.1	0.02	25.79	2.29	10	483	11.19	0.07	38.4
S34	6.16945	7.22443	155	3.5	0.8	0.04	21.76	2.79	18	658	15.92	0.23	181.4
\$35	6.15790	7.23035	41	4	0.7	0.04	31.36	3.6	32	804	16.17	0.22	47.4

Table 1. Raw concentrations of trace-elements in soils of Dagbala-Atte district. Ag, Au and Hg in ppb, Fe in % and other elements in ppm.

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S36	6.16160	7.20237	10	0.4	0.4		8.46	1.19	6	208	8.59	0.02	15.4
S37	6.17962	7.20980	10	0.5	0.8		4.32	0.86	9	306	9.2	0.04	7.3
S38	6.19903	7.21065	140	17.4	1.3	0.03	19.12	12.33	41	127	15.95	0.34	358.8
S39	6.21840	7.20993	27		0.4		20.48	2.26	17	382	11.71		102.5
S40	6.23597	7.21285	10	0.5	0.4		8.02	1.27	13	213	5.6	0.02	20.8
S41	6.25653	7.20315	11	1.4			6.85	1.51	19	296	18.83	0.05	66.1
S42	6.27455	7.21050	24	0.6	1.2	0.02	33.22	2.9	23	1159	17.25	0.05	104.4
S43	6.27132	7.18831	17	0.8	5.9		13.61	2.13	22	512	14.14	0.04	106.2
S44	6.24842	7.18557	15	0.5		0.01	6.27	1.62	26	451	20.2	0.04	167.7
S45	6.23840	7.18655	18	2.6	1.4		11.3	5.84	27	302	16.28	0.08	98.9
S46	6.21537	7.17475	43	1	0.4		7.69	2.39	36	493	26.94	0.03	145.5
S47	6.20342	7.18643	40	2.3	0.5	0.03	11.68	2.09	50	646	12.41	0.24	132.3
S48	6.17772	7.18448	7	0.7			11.72	1.72	12	254	9.09	0.07	18.9
S49	6.15867	7.17477	8	0.4	0.8	0.01	9.22	1.12	10	289	5.08	0.05	77.6

and necessary technique in interpretation of geochemical data [5] [7] [26] [27] [28] [29] [30]. Therefore, histograms, box plots, correlation matrix and factor analysis were constructed using the analytical results with the help of a software package - Minitab-16 in a DELL Inspiron 5521 Laptop computer. Both the histogram and box plots were done using the raw and log-transformed geochemical data. A summary of the basic statistics and other parameters derived from the box plots (1st quartile, 3rd quartiles and interquartile range) is tabulated for the 11 elements involved in the present study (Table 2). The table also includes threshold values calculated using formulae - Mean + 2* Standard deviation and $Q3 + 1.5^*$ interquartile range, for the logarithmically transformed data since the elements in the soils of the study area are log-normally distributed. However, the threshold values computed from formula Q3 + 1.5 interquartile range is adopted for this study because it gave relatively low values which enhance the number of anomalous sites that were recognized in this study.

3.3.2. Plotting of Geochemical Distribution Map

The distribution maps of the elements in the soils of the district based on the raw data in Table 1 and the soil sample location map of the study area (Figure 2 and Figure 3) drawn using Arc-GIS 10.4.1 program in a DELL Inspiron 5521 Laptop computer have been used to explain the distribution of the various elements in the soils of the study area.

4. Results and Discussion

4.1. Histograms

The histogram is one of the most popular graphical means of displaying distribution since it reflects the shape similar to theoretical frequency distributions.

Elements	Min. Conc.	1 st Quartile (Q ₁)	Median	Max. Conc.	3 rd Quartile (Q ₃)	No of samples where detected	Geometric Mean (X)	Standard Deviation (S)	Threshold Log ⁻¹ (X + 2S)	Threshold Log^{-1} $(Q_3 + 1.5$ $(Q_3 - Q_1)$
Ag	0.477	0.903	1.176	2.190	1.380	47	1.174	0.366	80.568	1.763
As	-1	1.020	-0.126	2.331	1.775	48	0.008	0.513	10.820	3.5
Au	-0.398	0.7	0.114	1.097	2.7	43	0.180	0.423	10.645	5.6
Cd	-2	0.01	-1.699	-1.222	0.033	26	-1.649	0.261	0.075	0.06
Cu	0.121	7.06	1.067	1.897	18.215	49	1.029	0.341	51.486	33.22
Fe	-0.921	1.18	0.328	1.091	2.81	49	0.265	0.356	9.467	4.43
Hg	0.778	10.5	1.279	2.037	28.75	48	1.272	0.265	63.264	50
Mn	1.204	221.5	2.486	3.064	481.5	49	2.452	0.374	1589.24	804
Pb	0.299	8.165	1.002	1.430	15.99	49	1.003	0.244	30.978	26.94
Sb	-1.699	0.03	-1.301	-0.420	0.065	41	-1.301	0.337	0.236	0.11
Zn	0.342	14.95	1.567	2.555	76.4	49	1.520	0.506	340.158	167.7

Table 2. Summary statistics for residual soils in Dagbala-Atte district, Nigeria.

Histograms plotted on raw data show positive skewness of the elements in varying degrees while those plotted on the log-transformed data display no skewness, which indicates that the distribution of the elements are log-normal. Examples of these two sets of histograms are presented in **Figure 4(a)** for histograms plot on raw data and **Figure 4(b)** for those plots on log-transformed data. Most of the raw data histograms show breaks in distribution, which might be interpreted as revealing mixtures of populations. Breaks in distribution occur at 75 ppb for Ag, 20 ppm for As, 37.5 ppm for Cu, 1.75% for Fe, 70 ppm for Hg, 0.12 ppm for Sb and 200 ppm for Zn. However, the log-transformed data histograms show breaks in distribution for only As and Cd.

4.2. Box Plots

The box plot is a method used to display some statistical parameters in a graphical form [31]. Provided the scale of presentation is reasonable, the box plot provides a fast visual estimate of the frequency distribution. Examples of the dispersion box plots of the trace metals based on both the raw and log-transformed data are presented in **Figure 5(a)** and **Figure 5(b)**, respectively. The raw data box plots (**Figure 5(a)**) show longer whiskers above the mean and shorter ones below it. This implies that most of the values of the elements greatly depart from the mean, which is also an indication of the extreme variability of geochemical data. This trend in any data set indicates non-normal situation. In the case of the log-transformed dispersion box plots (**Figure 5(b**)) a more refined dataset resulted with both the maximum and minimum values evenly distributed about the mean value. This provided a data set with a smaller and more stable variance



Figure 4. Histograms of (a) Raw Concentrations and (b) Logarithmically Transformed Concentrations of Ag and Cuin Soils of Dagbala-Atte District.



Figure 5. Boxplots of (a) Raw Concentrations and (b) Logarithmically Transformed Concentrations of Au and Pb in Soils of Dagbala-Atte District.

which aided greatly in data analysis. Each of the raw data and logarithmically transformed data plots shows the number of samples that contain anomalous concentrations of elements, called outlier values in the box plots (Figure 5(a) and Figure 5(b)).

4.3. Correlation Matrix

Multivariate statistical analysis, which involved correlation matrix and factor analysis, was conducted on ten elements, namely, Ag, As, Au, Cu, Fe, Hg, Mn, Pb, Sb and Zn. Cadmium (Cd), having too many censored data (i.e. values below analytical detection limit, DL), was excluded from this analysis [32]. For other elements such as Ag, Au, Hg and Sb that contain few censored data, values equal to 66% of their DLs were substituted [32] for the computation of correlation matrix and factor analysis. Table 3 shows the correlation matrix of the elements of interest in the Dagbala-Atte soil geochemical data. The matrix of simple correlation coefficients shows a fairly strong positive correlation between most metals and weak negative correlation between others. It was observed that all the correlation coefficients between metals were significant at 95% confidence level and above. The best correlation ($r \ge 0.75$) occurs between Cu-Fe and Fe-Zn, very good correlation ($0.65 \le r < 0.75$) exists between Fe-Pb and Hg-Pb, good correlation (0.55 \leq r < 0.65) occurs between each of Ag-Zn, As-Cu, As-Fe, As-Sb, Cu-Zn and Mn-Pb; and fairly good correlation ($0.35 \le r < 0.55$) exist between Ag-Fe, Ag-Hg, Ag-Mn, Ag-Pb, Ag-Sb, As-Au, As-Pb, As-Zn, Cu-Pb, Cu-Sb, Fe-Hg, Fe-Mn, Fe-Sb, Hg-Zn, Mn-Zn and Sb-Zn.

4.4. Factor Analysis

Table 4(a) is the rotated varimax factor matrix. The factor solutions provided information on eigenvalues (the amount of the total data explained by each model) and communalities (the amount of the total variability of each element explained in a given factor model) for the factor model obtained. The factor matrix

 Table 3. Pearson correlation matrix of the elements for log-data of soils in Dagbala-Atte district.

	Ag	As	Au	Cu	Fe	Hg	Mn	Pb	Sb	Zn
Ag	1									
As	0.291	1								
Au	-0.104	0.381	1							
Cu	0.250	0.609	0.223	1						
Fe	0.471	0.619	0.286	0.795	1					
Hg	0.434	0.264	0.010	0.213	0.457	1				
Mn	0.376	-0.114	-0.159	0.295	0.442	0.307	1			
Pb	0.531	0.362	0.078	0.443	0.695	0.658	0.580	1		
Sb	0.511	0.589	-0.088	0.351	0.432	0.143	0.043	0.149	1	
Zn	0.642	0.434	0.062	0.631	0.792	0.389	0.513	0.669	0.483	1

		(u)				
Variable	Factor 1	Factor 2	Factor 3	Communality		
Ag	0.761	-0.394	-0.264	0.804		
As	0.609	0.660	-0.127	0.822		
Au	0.213	0.591	0.573	0.724		
Cu	0.604	0.510	-0.068	0.630		
Fe	0.859	0.190	0.214	0.819		
Hg	0.517	-0.343	0.409	0.552		
Mn	0.202	-0.844	0.026	0.753		
Pb	0.730	-0.377	0.414	0.847		
Sb	0.682	0.124	-0.586	0.823		
Zn	0.810	-0.199	-0.132	0.713		
Variance	4.0590	2.2637	1.1642	7.4869		
% Var	0.406	0.226	0.116	0.749		
		(b)				
Fa	ctor	Elem	ents	Eigen Value (%)		
I	R-1	Fe-Zn-Ag-Pb-	Sb-As-Cu-Hg	40.6		
I	R-2	As-Au	ı-Cu	22.6		
Ι	R-3	Au-H	g-Pb	11.6		

 Table 4. (a). Varimax rotated factor loadings and communalities for log-data of soils in

 Dagbala-Atte district. (b). Element associations of the three-factor model.

(Table 4(a)) gave a three-factor model presented in Table 4(b). The total explained variance of the three-factor model is 74.9% (Table 4(a)). Table 4(a) also shows that each factor consists of significant contributions from certain variables and less important to negligible contributions from others. In summary, Factor R-1 has been deduced to be a Fe-Zn-Ag-Pb-Sb-As-Cu-Hg factor, Factor R-2, an As-Au-Cu factor and Factor R-3, a Au-Hg-Pb factor.

Factor R-1 is the most pronounced and accounts for 40.6% of the three-factor model in the area. It is interpreted to be probably a base metal (Zn, Pb and Cu) mineralization with which Fe, possibly in form of pyrite or pyrrhotite, is associated and Ag that is normally associated with Pb, probably on account of their close ionic radii. The occurrence of Sb, As and Hg, which are mobile under acidic, oxidizing soil conditions, in the base metal association suggests they can serve as pathfinder elements for these base metals [25] [33]. Factor R-2 also accounts for 22.6% of the model variance and was interpreted mainly as a gold mineralization factor with which Cu mineralization is associated and for which also the associated As can serve as a pathfinder element [33]. The high negative loading of Mn in this factor suggests that Mn is antipathetically related with this Au-Cu mineralization. Also, fairly high negative loading of Ag in this factor is a pointer to the fact that Ag mineralization does not associate with Au mineralization of this is

that the lithology that contains the base metal mineralization is different from the one that hosts the Au-Cu mineralization in the district. Factor R-3, which accounts for 11.6% of the model variance, has been interpreted as probably indicating a limited Hg mineralization with which Pb and Au are associated. The comparatively high negative loading of Sb in this factor suggests that the probable Hg mineralization might not be present in the same lithology that contains the base metal mineralization in Factor 1.

4.5. Geochemical Distribution of the Elements in Soils of Dagbala-Atte District

The geochemical distribution maps of the elements in the soils of the district have been employed to explain the distribution of the various elements in the soils of the study area. Examples of the geochemical distribution maps of the elements in the district are presented in **Figure 6**. To ensure that the anomalous sites, selected on the basis of the threshold values fixed for the elements, are meaningful these threshold values were compared with the background concentrations of selected trace elements in the world's soils (**Table 5**) [25] [33]. Where the concentration of any element in an anomalous site is lower than the average world soil's background value (**Table 5**) such an anomalous site was considered insignificant. The geological map of the study area was superimposed on the soil sample sites map (**Figure 3**) in order to link the distribution of the elements in the *in situ* soil of the district with their possible lithologic sources.

Ag has been detected in forty-seven samples, being present below the lower limit of detection of 2 ppb only in two samples. Ag concentrations range in the

Elements	Background Concentration in Soils (ppm)	Surficial Mobility	Use in Exploration
Ag*	100	High, Mn scavenged	-
As	10	Mobile, Fe scavenged	Pathfinder especially for Au
Au*	1	Low	Au deposits
Cd	0.1	High	Pathfinder for Zn deposits
Cu	15	pH > 5 low, else moderate	Most surveys
Fe			-
Hg∗	50	High	Wide use as pathfinder
Mn	300	Moderate, high at low pH	Scavenges Co, Zn, Ag
РЬ	15	Low	Wide use
Sb	1	Low	Pathfinder
Zn	35		-

Table 5. Background concentration of the analyzed trace elements in soils and utility in geochemical exploration. (Largely after Levinson, 1980 and Rose *et al.*, 1991).

*Unit of concentration in ppb.

soils of the study area between 3 and 155 ppb, with a mean of 22.51 ppb. The element therefore can be said to be fairly widely distributed in the soil of the study area. The threshold value computed for Ag in the District is 43 ppb; hence three anomalous values exist at sites DPS15 (58 ppm), DPS34 (155 ppm) and DPS38 (140 ppm). These anomalous soil sites are underlain by granitic gneiss in DPS15 and quartz-biotite schist in DPS34 and DPS38 (Figure 3). When the Ag concentrations of the anomalous sites are compared with the background concentration of Ag in crustal soils (Table 5), only two anomalous sites (DPS34 and DPS38) have values above the background and thus considered significant. The distribution map of silver tends to define an anomalous silver northwest distribution trend. On the basis of good correlation of Ag with Zn and Cu coupled with its fairly good correlation with Fe, Hg, Mn, Pb and Sb, which are all present in the Factor R-1 it is reasonable to infer that Ag is associated with the suspected Cu-Pb-Zn mineralization in the district. The negative loading of Ag in Factor-R2 (Table 4(a)), *i.e.* the Au factor, suggests that Ag and Au do not occur in the same mineralization type in the district.

As was detected in 48 samples where its values ranged from 0.1 to 214.5. The mean of As in the soils of the district is 5.95 ppm; and the computed threshold value is 10.8 ppm. With this threshold value, there exist two anomalous soil sites at DPS11 and DPS38, which are underlain by silicified sheared rock and garnet-biotite schist, respectively (Figure 3). In the north-central part of As distribution map, there is a strong anomalous region which lies over the silicified sheared rock. When the mean value of As in this area is compared with its background concentration in the earth's soils (Table 5), the As concentrations in this area are fairly low. Thus, the anomalous values recorded may not be associated with any concentration of arsenic minerals. However, they may be a pointer to the presence of sulphide deposits, with which some Au and traces of As are associated. The fairly good to very good correlation of As with Cu, Fe and Sb (Table 3) strongly suggests the presence of sulphide mineralization in the study area. The occurrence of As in Factors 1 and 2 indicates that As can serve as a pathfinder element for both the suspected Cu-Pb-Zn-Ag mineralization and Au-Cu mineralization in this district.

Au was detected in forty-two of the forty-nine samples employed for this study and it occurred below the detection limit of 0.2 ppb in other seven samples. Au geochemical distribution in the soils of the district is given in **Figure 6(a)**. Au ranges in concentrations from 0.4 to 12.5 ppb with a mean of 2.5 ppb. With the threshold value of 5.6 ppb selected for Au, six anomalous values occur at sites DPS7 (8.3 ppb), DPS11 (12.5 ppb), DPS15 (9.8 ppb), DPS18 (8.3 ppb), DPS30 (6.6 ppb) and DPS43 (5.9 ppb). These sites are underlain by quartz-biotite schist (DPS7), silicified sheared rock (DPS11), granitic gneiss (DPS15, DPS30, DPS43) and mica schist (DPS18) (**Figure 3**). Au distribution map indicates that Au forms the strongest anomaly in the north-central part of the district. This anomaly has a near ellipsoidal shape extending in the NE-SW



Figure 6. Geochemical Distribution Maps of (a) Au and (b) Zn in the soils of Dagbala-Atte District.

direction. In this north-central area, the strong anomalous sites of As and Au tend to overlap. Compared with the background concentration of Au in earth soils (**Table 5**), the concentration of Au in the soils of the area is generally very high, especially at the anomalous sites. The higher values of Au in the anomalous sites compared with the average earth's soil value suggest the occurrence of Au mineralization in parts of the study area. This coupled with the observed correlation of Au with As and the association of Au with Factors R-2 and R-3 supports the inference of Au-Cu mineralization in the silicified sheared rock, where the anomalous site recorded the highest value of 12,5 ppb of Au, and of Au-Hg-Pb mineralization in the granitic gneiss of Dagbala-Atte District.

Cu is widely distributed in the study area being detected in all the samples in concentrations ranging from 1.32 to 78.9 ppm with a mean of 14.2 ppm. With the threshold value of 33.2 ppm, only one anomalous value of copper exists at sample site DPS11 that is underlain by silicified sheared rock (Figure 3). The distribution map for Cu shows a strong anomalous area of Cu that coincides with that of As and Au. Compared with the background concentration of Cu (15 ppm) in soils (Table 5), the concentration of copper (78.9 ppm) in the soils of the study area is high. Therefore, the anomalous value is probably due to mineralization. Also, the relatively high correlation of copper with Fe, As, Zn, Pb and Sb (Table 3) and the association of copper with the trace elements that constitute Factors R-1 and R-2 (Table 4(a) & Table 4(b)) supports strongly the presence of Cu not only in the suspected Cu-Pb-Zn-Ag mineralization, which element distribution map suggests is located on the quartz-biotite schist, but also in the Au-Cu mineralization based on the silicified sheared rock of the district.

Hg has been detected in 48 of the 49 samples used for this study. Its values range from 6 to 109 ppb with mean content of 22.6 ppb. The selected threshold for mercury is 50 ppb. There are two anomalous values for mercury in the study area at site DPS12 (109 ppb) that is underlain by granitic gneiss (Figure 3). The distribution map shows a strong anomalous area with near ellipsoidal shape in the northern part of the area at DPS12 underlain by the granitic gneiss. Compared with its background concentration in soils (50 ppb) (Table 5), the concentration of mercury in the study area is moderate. The paucity of anomalous values probably suggests highly localized concentration of Hg in the area such as in site DPS 12 located on granite gneiss. The relatively high correlation of Hg with Pb, Ag, Fe and Zn, as well as the association of Hg with the elements in Factors R-1 and R-3 (Table 4) suggests the element could serve as a pathfinder element for the suspected Cu-Pb-Zn-Ag mineralization but the possibility of localized minor Hg-Pb-Au mineralization cannot be ruled out.

All the samples analyzed contain Pb in concentrations that range from 1.99 to 26.94 ppm with a mean of 11.55 ppm. With the selected threshold of 23.03 ppm, two anomalous values occur at sample sites DPS15 (24.99 ppm) and DPS46 (26.94 ppm). These sites of anomalous values are underlain by granitic gneiss (**Figure 3**) and are restricted to the eastern area of the district. Compared with

the background concentration of Pb (15 ppm) in the crustal soils (**Table 5**) the mean Pb concentration in the soils of Dagbala-Atte District is high. This indicates that there might be lead mineralization in the study area. Also the very good to fairly good correlation of Pb with Fe, Hg, Mn, Ag, As and Cu, as well as its association with the elements in Factors R-1 and R-3 supports the association of Pb with the suspected Cu-Pb-Zn-Ag mineralization and its presence in the minor Hg-bearing mineralization in the study area.

Sb has a fairly wide distribution in the study area and is present below the lower detection limit in eight of the samples. The detected values range from 0.02 to 0.38 ppm with a mean value of 0.07 ppm. The selected threshold value of 0.11 ppm indicated five anomalous values at sites DPS29 (0.38), DPS34 (0.23), DPS35 (0.22), DPS38 (0.34) and DPS47 (0.24). These anomalous sites are all distributed in the southern part of the district and lie within the quartz biotite schist, except site DPS29 that is within the granite of the district (Figure 3). The generally low concentration of Sb in the study area may be due to lack of Sb mineral or stibnite in the study area. Good correlation of Sb with As and its fairly good correlation with Ag, Cu, Fe and Zn, as well as its association with the elements in Factor R-1 probably reflect its presence as a minor trace element in the suspected Cu-Pb-Zn-Ag mineralization and its suitability as a pathfinder element for the mineralization in the district. The high negative loading of Sb in Factor R-3, unlike in the case of Factor R-1, probably indicates that the source of the Hg-bearing association in the former (Factor R-3) differs from the source of the suspected Cu-Pb-Zn-Ag mineralization associated with the latter (Factor R-1).

The distribution of Zn in the soils of the District is shown in **Figure 6(b)**. The concentrations of zinc in the soils of this district range from 2.2 to 358.8 ppm with a mean of 36.9 ppm. The selected threshold of 167.7 ppm revealed two anomalous values at sites DPS34 (181.4 ppm) and DPS38 (358.8 ppm). Quartz-biotite schist (**Figure 3**) underlies the area that yielded the anomalous zinc concentrations. The strong anomalous area for Zn falls in the south-central part of the district. However, the Zn values in the study area are low in comparison with the background concentration of Zn in soils (**Table 5**). Nevertheless, the high correlation of Zn with Pb, Ag and Cu (**Table 3**), tends to confirm the presence of Zn in the suspected Cu-Pb-Zn-Ag mineralization in the quartz-biotite schist of Dagbala-Atte District.

5. Summary and Conclusions

Geochemical soil survey for base and precious metals was undertaken in Dagbala-Atte District by sampling residual soil from 49 sites in a grid pattern. The forty nine soil samples were analyzed for eleven elements comprising base metals (Cu, Pb and Zn), precious metals (Au and Ag) and others including As, Cd, Fe, Hg, Mn and Sb commonly associated with the base and precious metals. Statistical analysis of the geochemical data generated involving correlation matrix and factor analysis enabled recognition of three major groups of elemental association or factors, viz, Fe-Zn-Ag-Pb-Sb-As-Cu-Hg factor (F1), As-Au-Cu factor (F2) and Au-Hg-Pb factor (F3). On account of high correlation coefficients between metals (**Table 3**) and high factor loadings of individual elements (**Table 4(a)** & **Table 4(b)**), three factor models were interpreted in terms of probable mineralization as follows: F1 is interpreted as a significant Cu-Pb-Zn-Ag mineralization; F2 as moderately strong Au-Cu mineralization and F3 as a minor Hg-bearing Au-Pb mineralization in the district.

By using the threshold values as cut-off points, the anomalous values for the various elements at the sample sites were determined. Superimposing the geochemical distribution maps of the various elements in the residual soils on the geological map of the district enables identification of the possible lithologic sources of these anomalous values [26] [27]. Hence, the anomalous values and the suspected mineralization types in the Dagbala-Atte District are linked with different rock types as follows:

1) The Cu-Pb-Zn-Ag mineralization is inferred to be located on quartz-biotite schist that underlies the southwestern part of the district around Sasaro and Egbigele.

2) The Au-Cu mineralization is based on the silicified sheared rock that occupies the northern and southern parts of the district around Dagbala and Atte.

3) Minor Hg-bearing Au-Pb mineralization in the granitic gneiss that underlies the northeastern and eastern part of the district around Erurhu and Atte.

The association of some trace elements such as As, Cd, Hg and Sb with the suspected mineralization suggests that they can be used as pathfinders on account of their relatively high concentration and mobility in the acidic, oxidizing soil environment [25] [33]. It would appear that As would be a suitable pathfinder for the suspected Au-Cu mineralization and As, Hg and Sb for the suspected Cu-Pb-Zn-Ag mineralization. Based on the foregoing linkage of the inferred mineralization types with some rock types in the study area, the lithogeochemical survey of the area is logically recommended.

Conflicts of Interest

The author declares no conflicts of interest regarding the publication of this paper.

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