

Hyperbolic Velocity Model

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Received March 27, 2013; revised April 29, 2013; accepted May 26, 2013

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ABSTRACT

Asymptotically bounded velocity profiles describe the vertical velocity variations in compacted sediments in a more realistic way than unbounded velocity models, and allow presenting the subsurface by a smaller number of thicker layers. The first and the simplest asymptotically bounded model is the Hyperbolic velocity profile proposed by Muscat in 1937, and our paper is an extension of this early study. The Hyperbolic model has an advantage over other bounded models: The velocity increases with depth and approaches the limiting value with a more smooth and gradual rate. We derive the time-depth relationships, forward and backward transforms between the instantaneous velocity profile and the effective models (average, RMS and fourth order average velocities), study the trajectories for pre-critical and post-critical curved rays and derive the equations for traveltime, lateral propagation and arc length. We compare the ray paths obtained with the Hyperbolic model and with the other bounded velocity profiles.

Keywords: Velocity Models; Velocity Transforms; Sediments

1. Introduction

The Hyperbolic velocity model was first proposed by Muscat [1] and published in 1937. However, since then the model was not extensively studied and is unjustifiably ignored in the literature. The objective of this research is to extend the original study and to correct the inaccuracies. We show the place of the Hyperbolic model among the other asymptotically bounded models, analyze its basic relationships and attempt to develop a complete theory.

Asymptotically bounded velocity models describe the velocity profile in compacted sediments, where the velocity gradually increases with depth and eventually approaches a limiting value. These models make it possible to describe a vertical velocity profile with a smaller number of intervals as compared to the classical unbounded models, such as linear velocity vs. depth [2,3], unbounded exponent [4], linear slowness [5], “sloth” (linear variation of slowness squared), e.g., [6], parabolic model [7,8], Faust velocity model [9,10] with a reference depth and different root indices. The unbounded models are described by two parameters: the instantaneous velocity at the top interface V_a and the vertical velocity gradient k_a at the same level. The Faust model includes also the root index n , normally $n \approx 6$. Asymptotically bounded models require an additional parameter: the limiting value of velocity V_∞ at infinite depth. Two models of

this family were studied by Ravve and Koren: the Exponential asymptotically bounded model [11,12] and the Conic model [13]. The asymptotically bounded profiles can be used, in particular, as velocity trend functions for the constrained velocity inversion with the best (e.g., least-squares) fit of the input data [14]. Examples of asymptotically bounded models are presented below. For each model, we first give the original formulation of the velocity profile as it appears in the original works by the authors, and then we convert it to a canonical form in terms of the “standard” parameters V_a, k_a and V_∞ . Parameter ΔV means the instantaneous velocity range, $\Delta V = V_\infty - V_a$.

- The Hyperbolic velocity model by Muscat [1],

$$z = A \cdot \frac{V(z) - V_a}{V_\infty - V(z)}, A = \text{const.} \quad (1)$$

In our notation, the Hyperbolic profile reads

$$V(z) = V_a + \Delta V \cdot \left(1 - \frac{\Delta V}{\Delta V + k_a z} \right). \quad (2)$$

- The Exponential velocity model by Muscat [1],

$$\frac{z}{B} = \ln \frac{V_\infty + V(z)}{V_\infty - V(z)} + \ln \frac{V_\infty - V_a}{V_\infty + V_a}, B = \text{const.} \quad (3)$$

We convert it to our notation,

$$V(z) = V_\infty \tanh(A + z/z_o), \quad (4)$$

where the parameters are

$$A = \operatorname{arccosh} \frac{V_\infty}{\sqrt{V_\infty^2 - V_a^2}}, z_o = \frac{V_\infty^2 - V_a^2}{k_a V_\infty}. \quad (5)$$

- The Exponential slowness model [5,15],

$$\frac{1}{V(z)} = \frac{1}{V_\infty} + \left(\frac{1}{V_a} - \frac{1}{V_\infty} \right) \cdot \exp\left(-\frac{z}{z_o}\right). \quad (6)$$

It can be converted to canonic form,

$$V(z) = \frac{V_a V_\infty}{V_a + \Delta V \exp(-z/z_o)}, z_o = \frac{V_a \Delta V}{k_a V_\infty}. \quad (7)$$

- The Exponential asymptotically bounded (EAB) velocity model [11,12],

$$V(z) = V_a + \Delta V \cdot \left[1 - \exp\left(-\frac{k_a z}{\Delta V}\right) \right]. \quad (8)$$

- The Conic velocity model [13],

$$\frac{V(z)}{V_\infty} = \frac{Q(z+h)}{\sqrt{1+Q^2(z+h)^2}}, \quad (9)$$

where

$$Q = \frac{k_a V_\infty^2}{(V_\infty^2 - V_a^2)^{3/2}}, h = \frac{V_a (V_\infty^2 - V_a^2)}{k_a V_\infty^2}. \quad (10)$$

A detailed review on unbounded and bounded velocity models is given by Kaufman [16]. **Figure 1** shows graphs of the instantaneous velocity vs. depth for the five asymptotically bounded velocity models mentioned above.

For all models, we assume the same velocity profile parameters: $V_a = 3 \text{ km/s}$, $k_a = 1 \text{ s}^{-1}$, $V_\infty = 6 \text{ km/s}$. The vertical gradients of the velocity vs. depth are plotted in **Figure 2**. It is interesting to note that among the five models presented, the Muscat Hyperbolic model (Equation (1) and grey line on the plot) approaches the limiting value V_∞ in the slowest and the most gradual manner. The “second slow” is the Conic velocity model (red line), and the “third slow” is the EAB model (blue line). An asymptotically bounded model can be characterized by its gradient-velocity relationship, which is actually the governing differential equation of the velocity model.

This paper is structured as follows. We define the Hyperbolic model using 1) the original Muscat [1] formulation—depth vs. velocity, 2) the physical parameters: maximum gradient R , length scale Q and vertical shift h , and 3) the “technical” or geophysical parameters: top interface velocity V_a , top gradient k_a and asymptotic velocity V_∞ . We introduce the dimensionless asymptotic factor M that simplifies the transform equations. First

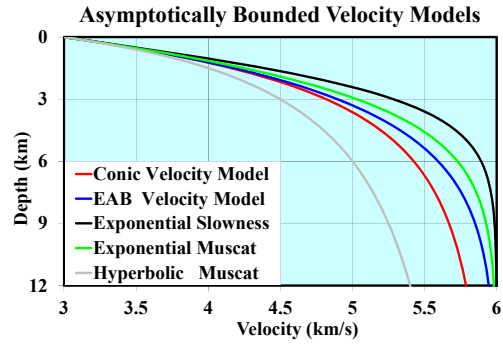


Figure 1. Asymptotically bounded velocity models: Muscat hyperbolic model, Muscat exponential model, the Exponential slowness model, the Exponential asymptotically bounded model and the Conic model: Velocity vs depth. For all models, the profile parameters are: the top interface velocity $V_a = 3 \text{ km/s}$, the top gradient $k_a = 1 \text{ s}^{-1}$ and the asymptotic velocity $V_\infty = 6 \text{ km/s}$.

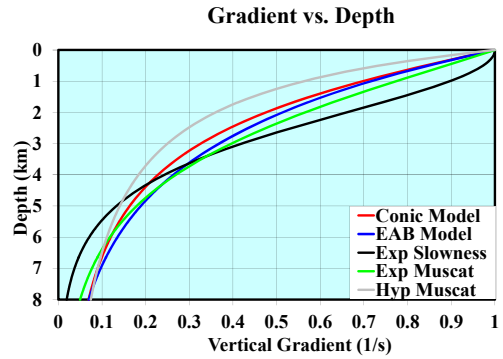


Figure 2. Vertical gradients vs. depth for asymptotically bounded velocity models.

we derive the time-depth and the depth-time relationships. Next we proceed to forward transforms from the instantaneous velocities to the effective models, such as the average, the RMS and the fourth order average velocity. Then we study the inversion problems, considering the inversion with the instantaneous velocities and gradients, and the inversion with the effective models, *i.e.*, the average or the RMS velocities given vs. time or depth. Next we comment on the two types of curved rays existing in all asymptotically bounded models, depending on the initial take-off angle, and derive the trajectories of the ray paths for the Muscat velocity profile. For both types of the curved rays we derive the lateral propagation, the traveltimes and the arc length.

2. The Hyperbolic Velocity Profile

Muscat [1] defined the Hyperbolic model by

$$\frac{z}{A} = \frac{V(z) - V_a}{V_\infty - V(z)}, \quad (11)$$

where V is the instantaneous velocity, z is depth measured

from the top interface, V_a is the top interface velocity, A is the characteristic distance (scale) that affects the top gradient k_a , and V_∞ is the asymptotic velocity. Inverting Equation (1), we obtain

$$V(z) = \frac{V_a + V_\infty z/A}{1 + z/A}. \quad (12)$$

The velocity gradient becomes

$$k(z) = \frac{dV}{dz} = \frac{\Delta V}{A} \cdot \frac{1}{(1 + z/A)^2}, \quad (13)$$

where $\Delta V = V_\infty - V_a$. At $z = 0$, the top gradient is $k = k_a$. Therefore,

$$k_a = \Delta V/A \quad \text{and} \quad A = \Delta V/k_a. \quad (14)$$

Introduce Equation (14) into Equation (13). In our notation, the Hyperbolic profile reads

$$V(z) = V_a + \Delta V \cdot \left(1 - \frac{\Delta V}{\Delta V + k_a z}\right) = \frac{V_a \Delta V + V_\infty k_a z}{\Delta V + k_a z}. \quad (15)$$

We call values V_a, k_a and V_∞ the technical parameters of the profile. At a definite height above the earth surface (above the upper interface), where $z = -h$, the instantaneous velocity vanishes. According to Equation (15),

$$h = \frac{\Delta V \cdot V_a}{k_a V_\infty}. \quad (16)$$

Introduce the absolute frame $\tilde{z} = z + h$, where the instantaneous velocity vanishes at the origin $\tilde{z} = 0$. Parameter h is the shift between the two frames of reference. In the absolute frame, the velocity profile simplifies to

$$V(\tilde{z}) = \frac{R\tilde{z}}{1 + Q\tilde{z}} = \frac{R \cdot (h + z)}{1 + Q \cdot (h + z)}, \quad (17)$$

where

$$Q = k_a V_\infty / \Delta V^2 \quad \text{and} \quad R = Q V_\infty. \quad (18)$$

We call values R, Q and h the physical parameters of the profile. Note that the linear velocity profile, where the ray trajectories are circular arcs, is a particular case of the Hyperbolic model with $Q \rightarrow 0$ and $V_\infty \rightarrow \infty$, in such way that their product $R = Q V_\infty$ remains a finite value, and parameter R becomes the constant velocity gradient of the linear model. The velocity gradient of the Hyperbolic model reads

$$k(z) = \frac{R}{(1 + Q\tilde{z})^2}. \quad (19)$$

At the absolute origin $\tilde{z} = 0$, the velocity gradient reaches its maximum value $k_{\max} = R$. Comparing Equations (17) and (19), we conclude that

$$\frac{V(z)}{V_\infty} + \sqrt{\frac{k(z)}{k_{\max}}} = 1. \quad (20)$$

Equation (20) is the governing differential equation of the Hyperbolic velocity profile. It can be used to plot the gradient-velocity diagram. Such diagrams for several asymptotically bounded velocity models are studied in Appendix A.

Introduce the normalized (dimensionless) velocity v , the normalized gradient \hat{k} and the normalized absolute depth \hat{z} ,

$$v = V/V_\infty, \quad \hat{k} = k/R, \quad \hat{z} = Q\tilde{z}. \quad (21)$$

Note that parameter Q is the reciprocal characteristic length. With these notations, the Hyperbolic velocity profile simplifies to

$$v = \frac{\hat{z}}{1 + \hat{z}}, \quad \hat{k} = \frac{1}{(1 + \hat{z})^2}. \quad (22)$$

The technical parameters of the velocity profile are related to the physical parameters,

$$V_a = \frac{Rh}{1 + Qh}, \quad k_a = \frac{R}{(1 + Qh)^2}, \quad V_\infty = \frac{R}{Q}, \quad (23)$$

The inverse relationship is

$$R = \frac{k_a V_\infty^2}{(V_\infty - V_a)^2}, \quad h = \frac{V_a}{k_a} \cdot \frac{V_\infty - V_a}{V_\infty}, \quad Q = \frac{R}{V_\infty}. \quad (24)$$

3. Asymptotic Factor

To simplify the equations for velocity transforms, it is suitable to introduce a special parameter M . This parameter can be defined at any point of the profile, and in particular, at the top and the bottom interfaces of an interval,

$$\begin{aligned} M(z) &= 1 + Q(h + z), \\ M_a &= 1 + Qh, \\ M_b &= 1 + Q(h + \Delta z), \end{aligned} \quad (25)$$

where Δz is the interval thickness (the vertical distance between the two interfaces), subscript a is related to the top interface $z = 0$, and subscript b is related to the bottom interface $z = \Delta z$. It follows from Equation (25) that

$$M_a = \frac{V_\infty}{V_\infty - V_a}, \quad M_b = \frac{V_\infty}{V_\infty - V_b}, \quad (26)$$

where V_a and V_b are the top and bottom instantaneous velocities, respectively. Next, it follows from Equation (26) that parameter M is the inverse normalized measure of the difference between the velocity at the given depth level and the asymptotic velocity V_∞ . Equation (26) can be inverted,

$$\frac{V_a}{V_\infty} = \frac{M_a - 1}{M_a}, \frac{V_b}{V_\infty} = \frac{M_b - 1}{M_b}. \quad (27)$$

The velocity gradient is also related to the asymptotic factor,

$$k_a = \frac{R}{M_a^2}, k_b = \frac{R}{M_b^2}. \quad (28)$$

It follows from Equation (25) that

$$Q = \frac{M_b - M_a}{\Delta z} = \frac{V_\infty (V_b - V_a)}{\Delta z (V_\infty - V_a) \cdot (V_\infty - V_b)}, \quad (29)$$

and

$$Qh = M_a - 1 = \frac{V_a}{V_\infty - V_a}, \quad (30)$$

$$Q(h + \Delta z) = M_b - 1 = \frac{V_b}{V_\infty - V_b}.$$

We use Equations (28) and (29) to get the interface gradients, k_a and k_b , through the increment of the asymptotic factor, $\Delta M = M_b - M_a$,

$$k_a = \frac{M_b - M_a}{M_a^2} \cdot \frac{V_\infty}{\Delta z}, k_b = \frac{M_b - M_a}{M_b^2} \cdot \frac{V_\infty}{\Delta z}. \quad (31)$$

It follows from Equation (31),

$$M_b = M_a + M_a^2 \cdot \frac{k_a \Delta z}{V_\infty}, M_a = M_b - M_b^2 \cdot \frac{k_b \Delta z}{V_\infty} \quad (32)$$

Equations (27) and (29) result in the average gradient on the interval, k_{ave} , expressed either through the interface asymptotic factors M_a and M_b , or through the interface gradients k_a and k_b ,

$$k_{ave} = \frac{\Delta V}{\Delta z} = \frac{V_b - V_a}{\Delta z} = \frac{V_\infty}{\Delta z} \cdot \frac{M_b - M_a}{M_a M_b} \quad (33)$$

$$= \frac{Q V_\infty}{M_a M_b} = \frac{R}{M_a M_b}.$$

Introduction of Equation (28) into Equation (33) leads to

$$k_{ave} = \sqrt{k_a k_b}. \quad (34)$$

The average gradient on the interval with the Hyperbolic velocity profile is the geometric average of the top and bottom interface gradients.

Given the velocity and its gradient at one interface, one can calculate these parameters at the other interface. The calculations can be done either in depth or in time. Four problems of this kind are considered in Appendix C.

4. Depth-Traveltime Relationship

Integrate the slowness to get the vertical traveltime vs. the interval thickness,

$$\Delta t = \int_0^{\Delta z} \frac{dz}{V(z)} = \int_h^{h+\Delta z} \frac{d\tilde{z}}{V(\tilde{z})} = \frac{1}{V_\infty} \int_h^{h+\Delta z} \frac{1+Q\tilde{z}}{Q\tilde{z}} d\tilde{z} \quad (35)$$

$$= \frac{\Delta z}{V_\infty} + \frac{1}{Q V_\infty} \cdot \ln \frac{h+\Delta z}{h}.$$

The traveltime equation can be written in terms of asymptotic factors at the top and bottom interfaces, M_a and M_b . With the use of Equations (25) and (29), we obtain

$$V_\infty \Delta t = \Delta z \cdot \left(1 + \frac{1}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} \right), \quad (36)$$

where the top asymptotic factor M_a is calculated with Equation (26), and the bottom asymptotic factor $M_b(\Delta z)$ with the first equation of Equation Set (32). The interval velocity (local average velocity) through the layer between the interfaces becomes

$$\frac{V_{int}}{V_\infty} = \frac{\Delta z}{V_\infty \Delta t} = \frac{M_b - M_a}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}}. \quad (37)$$

To get the vertical distance vs. traveltime we should invert Equation (26), *i.e.* find $M_b(\Delta t)$. Introduction of Equation (29) into (36) results in

$$Q V_\infty \Delta t = M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}. \quad (38)$$

Equation (38) should be solved for the unknown bottom asymptotic factor M_b ,

$$M_b - 1 + \ln(M_b - 1) \quad (39)$$

$$= M_a - 1 + \ln(M_a - 1) + R \Delta t.$$

Taking exponent from both sides of Equation (39), we get

$$(M_b - 1) \cdot \exp(M_b - 1) \quad (40)$$

$$= (M_a - 1) \cdot \exp(M_a - 1) \cdot \exp(R \Delta t).$$

Equation (40) can be solved with the Lambert function,

$$M_b - 1 = L_0 \left[(M_a - 1) \cdot \exp(M_a - 1) \cdot \exp(R \Delta t) \right], \quad (41)$$

where notation L_0 means the zero branch of the Lambert function. The Lambert function $y = L(x)$ delivers the solution of the transcendent equation $x = y \cdot \exp y$, see Appendix B for details. In terms of the interface velocities, Equation (41) reduces to

$$\frac{V_b}{V_\infty - V_b} = L_0 \left[\frac{V_a}{V_\infty - V_a} \cdot \exp \frac{V_a}{V_\infty - V_a} \cdot \exp(R \Delta t) \right]. \quad (42)$$

After the bottom asymptotic factor M_b or the bottom interface velocity V_b is found, the interval thickness can be established with Equation (29),

$$\Delta z = \frac{M_b - M_a}{Q} = \frac{V_\infty (V_b - V_a)}{Q(V_\infty - V_a)(V_\infty - V_b)}. \quad (43)$$

5. Hyperbolic and Non-Hyperbolic Moveout

In the absence of the intrinsic anellipticity, the hyperbolic parameter W and the non-hyperbolic parameter H on the interval are defined as

$$\begin{aligned} W &= \int_{t_a}^{t_b} V^2 dt = \int_{z_a}^{z_b} V dz, \\ H &= \int_{t_a}^{t_b} V^4 dt = \int_{z_a}^{z_b} V^3 dz. \end{aligned} \quad (44)$$

Introduce the velocity profile from Equation (17). The hyperbolic parameter W becomes

$$\begin{aligned} W &= \int_{z_a}^{z_b} V dz = \int_{\tilde{z}=h}^{\tilde{z}=h+\Delta z} \frac{R\tilde{z} dz}{1+Q\tilde{z}} \\ &= V_\infty \Delta z - \frac{V_\infty}{Q} \cdot \ln \frac{1+Q(h+\Delta z)}{1+Qh}. \end{aligned} \quad (45)$$

The non-hyperbolic parameter H becomes

$$\begin{aligned} H &= \int_{z_a}^{z_b} V^3 dz = \int_{\tilde{z}=h}^{\tilde{z}=h+\Delta z} \frac{R^3 \tilde{z}^3 d\tilde{z}}{(1+Q\tilde{z})^3} \\ &= V_\infty^3 \Delta z + \frac{3V_\infty^3}{Q} \cdot \frac{1}{1+Qh} - \frac{3V_\infty^3}{Q} \cdot \frac{1}{1+Q(h+\Delta z)} \\ &\quad - \frac{V_\infty^3}{2Q} \cdot \frac{1}{(1+Qh)^2} + \frac{V_\infty^3}{2Q} \cdot \frac{1}{[1+Q(h+\Delta z)]^2} \\ &\quad - \frac{3V_\infty^3}{Q} \cdot \ln \frac{1+Q(h+\Delta z)}{1+Qh}. \end{aligned} \quad (46)$$

With the use of the top and bottom asymptotic factors, the hyperbolic parameter becomes

$$\begin{aligned} W &= \frac{V_\infty}{Q} \cdot \left(M_b - M_a - \ln \frac{M_b}{M_a} \right) \\ &= V_\infty \Delta z \cdot \left(1 - \frac{1}{M_b - M_a} \cdot \ln \frac{M_b}{M_a} \right). \end{aligned} \quad (47)$$

Introducing Equation (37) for the traveltime into Equation (47), we obtain the local RMS velocity U over the interval. By definition, $U = \sqrt{W/\Delta t}$, so

$$\frac{U^2}{V_\infty^2} = \frac{M_b - M_a - \ln \frac{M_b}{M_a}}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}}. \quad (48)$$

The non-hyperbolic parameter becomes

$$\begin{aligned} H &= \frac{V_\infty^3}{Q} \cdot \left(M_b - M_a + \frac{3}{M_a} - \frac{3}{M_b} \right) \\ &\quad + \frac{V_\infty^3}{Q} \cdot \left(-\frac{1}{2M_a^2} + \frac{1}{2M_b^2} - 3 \cdot \ln \frac{M_b}{M_a} \right). \end{aligned} \quad (49)$$

With the use of Equation (29), the non-hyperbolic parameter simplifies to

$$\begin{aligned} H &= V_\infty^3 \Delta z \\ &\quad \times \left(1 + \frac{3}{M_a M_b} - \frac{M_a + M_b}{2M_a^2 M_b^2} - \frac{3}{M_b - M_a} \cdot \ln \frac{M_b}{M_a} \right). \end{aligned} \quad (50)$$

When the parameters of the velocity profile are specified, the top asymptotic factor M_a is a known value. The bottom asymptotic factor M_b can be presented either vs. depth (interval thickness) or vs. traveltime. Thus, the hyperbolic and non-hyperbolic parameters become functions of depth or traveltime, accordingly. The anellipticity induced by the vertically varying velocity is defined as the fractional difference between the fourth-order average velocity V_4 and the RMS velocity V_2 ,

$$\eta = \frac{V_4^4 - V_2^4}{8V_2^4}. \quad (51)$$

Parameter η can be also considered as a function of depth or vertical time. For a particular case of a single infinite layer (half-space) with any vertical velocity profile,

$$V_2^2 = \frac{W}{t}, V_4^4 = \frac{H}{t} \rightarrow \eta = \frac{H \cdot t - W^2}{8 \cdot W^2}. \quad (52)$$

The graph for the induced anellipticity η is plotted vs. depth in **Figure 3** for three asymptotically bounded velocity models: Exponential, Conic and Hyperbolic. For all the three models, the parameters of the velocity profile are: $V_a = 3$ km/s, $k_a = 1$ s⁻¹ and $V_\infty = 6$ km/s. At the surface, the anellipticity is zero as there are yet no accumulated variations of the instantaneous velocity. The induced anellipticity is always positive. It reaches a maximum value a definite depth and then vanishes at the

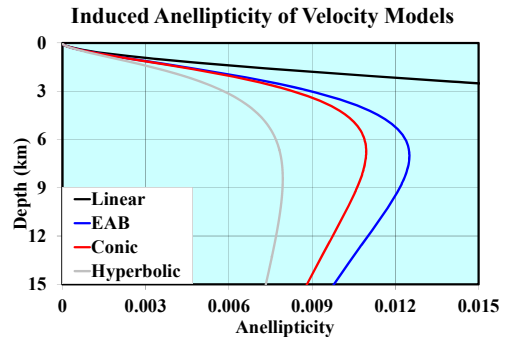


Figure 3. Induced anellipticity vs. depth for asymptotically bounded velocity models.

infinity, where the medium velocity is asymptotically constant.

6. Forward Dix Transform

Consider a package of n layers (vertical intervals), where the nodes (interfaces) are enumerated from zero, and layers are enumerated from 1. Interval n connects nodes $n-1$ (top interface) and n (bottom interface). The nodal average velocity $V_{1,n}$, RMS velocity $V_{2,n}$ and fourth-order average velocity $V_{4,n}$ are

$$\begin{aligned} V_{1,n} &= \frac{V_{1,n-1} \cdot t_{n-1} + \Delta z_n}{t_{n-1} + \Delta t_n}, \\ V_{2,n}^2 &= \frac{V_{2,n-1}^2 \cdot t_{n-1} + W_n}{t_{n-1} + \Delta t_n}, \\ V_{4,n}^4 &= \frac{V_{4,n-1}^4 \cdot t_{n-1} + H_n}{t_{n-1} + \Delta t_n}. \end{aligned} \quad (53)$$

where Δt_n is the one-way interval traveltime, Δz_n is the layer thickness, W_n and H_n are the interval hyperbolic and non-hyperbolic parameters, respectively. For $n=1$ we set $t_0=0$ in Equation (53). The effective velocities (average, RMS and fourth-order average) can be also defined for any internal point of the interval.

7. Inverse Dix Transform

Recall that the Hyperbolic velocity profile on the interval is defined by the three parameters: the top interface instantaneous velocity V_a , the top interface gradient k_a , and the asymptotic velocity V_∞ . We consider that the asymptotic velocity is always given a priori. When the two other parameters, V_a and k_a , are also known, then velocity transforms are considered forward. When one or both parameters are unknown (with another data specified instead), we deal with the velocity inversion. There are three groups of inverse transforms studied in Appendices D, E and F.

Appendix D considers the inversion that does not involve the RMS velocity. These formulations deal with the instantaneous velocity and its gradient only. We solve a problem where the two velocities are given at the interfaces, V_a and V_b , or—alternatively—the two gradients, k_a and k_b . Another kind of problem is when the velocity and its gradient are given at the different interfaces of the interval, *i.e.* the velocity is given at the top interface and the gradient—at the bottom interface, V_a and k_b , or vice versa, k_a and V_b . We solve also a problem where the instantaneous velocity is given at the bottom interface and at the intermediate point of the interval, V_b and V_c . These problems are studied both vs. depth and vs. time.

Appendix E considers the inversion with the RMS ve-

locity specified at the interfaces vs. depth or time with a single parameter unknown, either V_a or k_a . We consider also a problem with the traveltime specified vs. the interval thickness, also with a single parameter unknown. Finally, we consider the RMS velocity specified vs. *both* depth and time, with the two parameters unknown, V_a and k_a .

In Appendix F we study the two-interval inversion. The RMS velocity is given vs. depth or time at the two interfaces and at an internal point of the interval. Alternatively, depth can be specified vs. traveltime at the three points. Both parameters of the velocity profile are unknown. This is a so-called three-point or two-interval inversion.

8. Ray Trajectories

In this section we establish the trajectories of non-vertical rays. Due to Snell's law, in 1D medium the horizontal slowness p is constant, and the ray angle α (measured from the vertical axis) becomes

$$\sin \alpha = pV(z). \quad (54)$$

Introduce the ray parameter m ,

$$m = \frac{Q}{pR} = \frac{1}{pV_\infty} \equiv \frac{1}{P}, \quad (55)$$

where Q and R are the physical parameters of the Hyperbolic velocity profile, $P = pV_\infty$ is the normalized ray slowness, and m is its inverse value. We call parameter m "eccentricity of the ray trajectory" as it is very similar to the eccentricity of the hyperbolic and elliptic rays of the Conic velocity model [13]. With Equation (17), the sine of the ray angle becomes

$$\sin \alpha = \frac{pR\tilde{z}}{1+Q\tilde{z}} = \frac{1}{m} \cdot \frac{Q\tilde{z}}{1+Q\tilde{z}}, \quad (56)$$

so that the tangent of this angle is

$$\begin{aligned} \tan \alpha &= \pm \frac{\sin \alpha}{\sqrt{1 - \sin^2 \alpha}} \\ &= \pm \frac{Q\tilde{z}}{\sqrt{m^2 + 2m^2Q\tilde{z} - m^2Q^2\tilde{z}^2}}, \end{aligned} \quad (57)$$

where $m'^2 = 1 - m^2$. Parameter m'^2 (the conjugate eccentricity squared) may be positive or negative. Introduce the dimensionless coordinates,

$$\hat{x} = Qx, \hat{z} = Q\tilde{z}. \quad (58)$$

The tangent of the ray angle becomes

$$\tan \alpha = \pm \frac{1}{m} \cdot \frac{\hat{z}}{\sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2}} = \frac{dx}{d\tilde{z}} = \frac{d\hat{x}}{d\hat{z}}. \quad (59)$$

Integrating Equation (59), we obtain

$$\pm m(\hat{x} - \hat{x}_c) = \int \frac{\hat{z} d\hat{z}}{\sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2}}, \quad (60)$$

where x_c is the constant of integration. This integral can be reduced to

$$\begin{aligned} & \int \frac{\hat{z} d\hat{z}}{\sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2}} \\ &= -\frac{m^2}{m'^2} \cdot \sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2} \\ &+ \frac{m^2}{m'^2} \int \frac{d\hat{z}}{\sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2}}. \end{aligned} \quad (61)$$

To obtain the integral on the right side of Equation (61), we consider two cases, or two ranges of the eccentricity: $m > 1$ (pre-critical rays) and $m < 1$ (post-critical rays),

$$\begin{aligned} & \int \frac{d\hat{z}}{\sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2}} \\ &= \frac{m}{\sqrt{-m'^2}} \cdot \operatorname{arccosh} \frac{m^2 - m'^2 \hat{z}}{m} \text{ for } m > 1, \\ & \int \frac{d\hat{z}}{\sqrt{1 + 2\hat{z} - m'^2/m^2 \cdot \hat{z}^2}} \\ &= \frac{m}{\sqrt{+m'^2}} \cdot \operatorname{arccos} \frac{m^2 - m'^2 \hat{z}}{m} \text{ for } m < 1. \end{aligned} \quad (62)$$

For a limiting case $m = 1$ (critical rays),

$$\pm(\hat{x} - \hat{x}_c) = \int \frac{\hat{z} d\hat{z}}{\sqrt{1 + 2\hat{z}}} = \frac{(\hat{z} - 1) \cdot \sqrt{1 + 2\hat{z}}}{3}, \quad (63)$$

We emphasize that two kinds of rays exist for any monotonously increasing and asymptotically bounded velocity model, and in particular, for the Hyperbolic model. The pre-critical rays that may start on the earth surface, propagate to the infinite depth, and their curvature asymptotically vanishes. The post-critical rays have a limited propagation depth. Their arc-like trajectories have a finite minimum curvature at the turning point, and these rays return to the earth surface. Note that at any point of the trajectory, the ray path curvature Λ depends on the velocity gradient k only,

$$\Lambda(z) = pk(z). \quad (64)$$

In particular, the linear velocity model with a constant velocity gradient leads to ray trajectories of constant curvatures, *i.e.* to the circular arcs.

The critical rays with the unit eccentricity $m = 1$ are the limit case between the two types of rays. Their take-off angle (the ray angle at the upper interface) is called the critical angle α_c . It follows from Equations (54) and (55) that the critical take-off angle is

$$\alpha_c = \arcsin \frac{V_a}{V_\infty}. \quad (65)$$

It follows from Equations (61) and (62) that the trajectories of the pre-critical and post-critical rays are

$$\begin{aligned} & \text{for } m > 1, \pm(\hat{x} - \hat{x}_c) \\ &= -\frac{m}{m'^2} \times \left(\sqrt{1 + 2\hat{z} - \frac{m'^2 \hat{z}^2}{m^2}} - \frac{m}{\sqrt{-m'^2}} \cdot \operatorname{arccosh} \frac{m^2 - m'^2 \hat{z}}{m} \right), \\ & \text{for } m < 1, \pm(\hat{x} - \hat{x}_c) \\ &= -\frac{m}{m'^2} \times \left(\sqrt{1 + 2\hat{z} - \frac{m'^2 \hat{z}^2}{m^2}} - \frac{m}{\sqrt{+m'^2}} \cdot \operatorname{arccos} \frac{m^2 - m'^2 \hat{z}}{m} \right). \end{aligned} \quad (66)$$

At infinite depth, pre-critical rays become asymptotically straight. Equation (57) leads to

$$\tan \alpha_\infty = \frac{1}{\sqrt{m^2 - 1}}, \sin \alpha_\infty = \frac{1}{m} = pV_\infty. \quad (67)$$

However, although the slope of these rays converges to a constant value and their curvature becomes infinitesimal, the pre-critical rays of the Hyperbolic model have no asymptotic straight line, unlike the pre-critical rays of the EAB and the Conic models. The pre-critical, critical and post-critical rays are plotted in **Figure 4** for the Hyperbolic, the Conic and the Exponential (EAB) models. Parameters of the velocity profile are the same

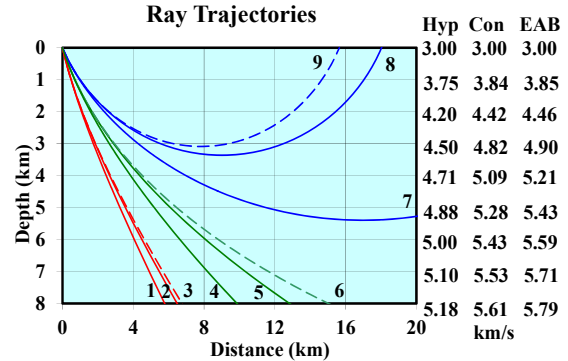


Figure 4. Pre-critical (red lines), critical (green lines) and post-critical (blue lines) ray trajectories for Hyperbolic (lines 1, 4, 7), Conic (lines 2, 5, 8) and EAB (lines 3, 6, 9) velocity models. Pre-critical rays propagate to an infinite depth and become asymptotically straight. Critical ray propagate to an infinite depth, and the ray angle approaches $\pi/2$ at large depth, but this ray has no asymptote. Post-critical rays pass the turning point and return to the earth surface, their trajectories are symmetric arcs. The Conic rays pass above the Hyperbolic model rays because their curvature is larger. For the same reason, the EAB rays pass above the Conic rays. For all models, parameters of the medium are the same as in Figure 1. The take-off angle of the pre-critical rays is $3\pi/24$, that of the critical rays $\pi/6$, and that of the post-critical rays $5\pi/24$.

as above. The three columns of numbers to the right of the plot area are velocities for the three models at the specified depth levels.

9. Maximum Penetration Depth

Pre-critical rays penetrate to infinite depth. The maximum penetration depth of post-critical rays follows from Equation (59). At the turning point, the ray angle $\alpha = \pi/2$, and thus, its tangent becomes infinite. This leads to a quadratic equation with a single positive root,

$$\hat{z}_{\max} = \frac{m}{1-m}. \tag{68}$$

Recall that \hat{z} is the dimensionless depth measured from the absolute origin (above the upper interface). The maximum penetration depth in units of length, measured from the upper interface reads

$$\begin{aligned} z_{\max} &= \frac{V_{\infty} - V_a}{k_a} \cdot \frac{1 - pV_a}{pV_{\infty} - 1} \\ &= \frac{V_a}{k_a} \cdot \frac{(1 - \sin \alpha_a) \cdot (1 - \sin \alpha_c)}{\sin \alpha_a - \sin \alpha_c}. \end{aligned} \tag{69}$$

10. Lateral Propagation, Traveltime and Arc Length

In a 1D medium, it is convenient to express the lateral propagation distance x , traveltime t_s and arc length s through the ray angle α and angle-dependent gradient $k(\alpha)$. These relationships are [12,16] (Kaufman, 1953; Ravve and Koren, 2006)

$$\begin{aligned} x &= \frac{1}{p} \int_{\alpha_a}^{\alpha_b} \frac{\sin \alpha d\alpha}{k(\alpha)}, \\ t_s &= \int_{\alpha_a}^{\alpha_b} \frac{d\alpha}{\sin \alpha k(\alpha)}, \\ s &= \frac{1}{p} \int_{\alpha_a}^{\alpha_b} \frac{d\alpha}{k(\alpha)}. \end{aligned} \tag{70}$$

where α_a and α_b are ray angles at the departure and the destination points, respectively. Equations (17) and (19) make it possible to eliminate depth and to express the gradient through the velocity,

$$k = R \cdot \frac{(V_{\infty} - V)^2}{V^2}. \tag{71}$$

Next we apply Snell's law and obtain the vertical gradient vs. the ray angle,

$$k(\alpha) = R \cdot (1 - m \sin \alpha)^2. \tag{72}$$

Note that

$$\begin{aligned} pR \cdot \frac{dx}{d\alpha} &= \frac{Q}{m} \cdot \frac{dx}{d\alpha} = \frac{\sin \alpha}{(1 - m \sin \alpha)^2} \\ &= \frac{1}{m(1 - m \sin \alpha)^2} - \frac{1}{m(1 - m \sin \alpha)}, \end{aligned} \tag{73}$$

$$\begin{aligned} R \cdot \frac{dt_s}{d\alpha} &= \frac{1}{\sin \alpha \cdot (1 - m \sin \alpha)^2} \\ &= \frac{1}{\sin \alpha} + \frac{m}{1 - m \sin \alpha} + \frac{m}{(1 - m \sin \alpha)^2}, \end{aligned}$$

$$\int \frac{d\alpha}{\sin \alpha} = \ln \tan \frac{\alpha}{2}, \tag{74}$$

$$\begin{aligned} &\int \frac{d\alpha}{(1 - m \sin \alpha)^2} \\ &= \frac{1}{1 - m^2} \times \left(\int \frac{d\alpha}{1 - m \sin \alpha} - \frac{m \cos \alpha}{1 - m \sin \alpha} \right), m \neq 1. \end{aligned} \tag{75}$$

The indefinite integral on the right side of this equation essentially depends on the range of the eccentricity m , resulting in

$$\begin{aligned} I^{\text{pre}} &\equiv \int \frac{d\alpha}{1 - m \sin \alpha} \\ &= \frac{2}{\sqrt{-m^2}} \times \operatorname{arccoth} \frac{m - \tan(\alpha/2)}{\sqrt{-m^2}} \text{ for } m > 1, \text{ pre-critical,} \\ I^{\text{pst}} &\equiv \int \frac{d\alpha}{1 - m \sin \alpha} \\ &= \frac{2}{\sqrt{+m^2}} \times \operatorname{arccot} \frac{m - \tan(\alpha/2)}{\sqrt{+m^2}} \text{ for } m < 1, \text{ post-critical.} \end{aligned} \tag{76}$$

For the critical ray, $m = 1$

$$\begin{aligned} \int \frac{d\alpha}{1 - \sin \alpha} &= \frac{2 \sin(\alpha/2)}{\sqrt{1 - \sin \alpha}}, \\ \int \frac{d\alpha}{(1 - \sin \alpha)^2} &= \frac{\cos(3\alpha/2) + 3 \sin(\alpha/2)}{3(1 - \sin \alpha)^{3/2}}. \end{aligned} \tag{77}$$

Let α_a and α_b be the ray angles at the start point and the destination point of the ray path, respectively. Equation Set (76) can be re-arranged as follows

- For the pre-critical rays, $m > 1$

$$\begin{aligned} I^{\text{pre}} &\equiv \frac{2}{\sqrt{-m^2}} \operatorname{arccoth} \frac{m - \tan(\alpha/2)}{\sqrt{-m^2}} \Big|_{\alpha_a}^{\alpha_b} \\ &= \frac{2}{\sqrt{-m^2}} \operatorname{arccoth} \frac{\cos \frac{\alpha_b - \alpha_a}{2} - m \sin \frac{\alpha_a + \alpha_b}{2}}{\sqrt{-m^2} \cdot \sin \frac{\alpha_b - \alpha_a}{2}}, \end{aligned} \tag{78}$$

- For the post-critical rays, $m < 1$

$$I^{pst} \equiv \frac{2}{\sqrt{+m'^2}} \operatorname{arccot} \frac{m - \tan(\alpha/2) \Big|_{\alpha_b}}{\sqrt{+m'^2} \Big|_{\alpha_a}} \quad (79)$$

$$= \frac{2}{\sqrt{+m'^2}} \operatorname{arccot} \frac{\cos \frac{\alpha_b - \alpha_a}{2} - m \sin \frac{\alpha_a + \alpha_b}{2}}{\sqrt{+m'^2} \cdot \sin \frac{\alpha_b - \alpha_a}{2}}.$$

$I^{pre}(\alpha_a, \alpha_b)$ and $I^{pst}(\alpha_a, \alpha_b)$ are functions of the ray angles at the endpoints of the path. The following identities were used,

$$\operatorname{arccot} B - \operatorname{arccot} A = \operatorname{arccot} \frac{A \cdot B + 1}{A - B}, \quad (80)$$

$$\operatorname{arccoth} B - \operatorname{arccoth} A = \operatorname{arccoth} \frac{A \cdot B - 1}{A - B}.$$

To simplify the notations, we introduce one more function of the ray angles at the endpoints,

$$J(\alpha_a, \alpha_b) = \frac{\cos \alpha \Big|_{\alpha_b}}{1 - m \sin \alpha \Big|_{\alpha_a}} \quad (81)$$

$$= 2 \sin \frac{\alpha_b - \alpha_a}{2} \cdot \frac{m \cos \frac{\alpha_b - \alpha_a}{2} - \sin \frac{\alpha_a + \alpha_b}{2}}{(1 - m \sin \alpha_a) \cdot (1 - m \sin \alpha_b)}.$$

The normalized lateral propagation becomes

$$\frac{Qx}{m} = \frac{m \cdot I(\alpha_a, \alpha_b) - J(\alpha_a, \alpha_b)}{m'^2}, m \neq 1, \quad (82)$$

The normalized traveltime is

$$Rt_s = \ln \frac{\tan(\alpha_b/2)}{\tan(\alpha_a/2)} + \frac{m(1+m'^2) \cdot I(\alpha_a, \alpha_b) - m^2 \cdot J(\alpha_a, \alpha_b)}{m'^2}, \quad (83)$$

$m \neq 1$.

The normalized arc length is

$$\frac{Qs}{m} = \frac{I(\alpha_a, \alpha_b) - m \cdot J(\alpha_a, \alpha_b)}{m'^2}, m \neq 1. \quad (84)$$

For the critical rays, $m = 1$, and the departure angle is critical, $\alpha_a = \alpha_c$. The ray path parameters are,

$$Q \cdot x = \left[\frac{1}{1 - \sin \alpha} - \frac{2}{3} \cdot \frac{\cos(3\alpha/2)}{(1 - \sin \alpha)^{3/2}} \right] \Big|_{\alpha_c}^{\alpha_b},$$

$$R \cdot t_s = \left[\ln \tan(\alpha/2) - \frac{1}{1 - \sin \alpha} \right] \Big|_{\alpha_c}^{\alpha_b} + \frac{2}{3} \cdot \frac{3 \sin(\alpha/2) + 2 \cos(3\alpha/2)}{(1 - \sin \alpha)^{3/2}} \Big|_{\alpha_c}^{\alpha_b}, \quad (85)$$

$$Q \cdot s = \frac{\cos(3\alpha/2) + 3 \sin(\alpha/2)}{3(1 - \sin \alpha)^{3/2}} \Big|_{\alpha_c}^{\alpha_b}.$$

The current depth can be also expressed through the ray angle. It follows from Equation (17) that

$$\tilde{z}(\alpha) = \frac{1}{R} \cdot \frac{V}{1 - V/V_\infty} \rightarrow Q \cdot (z_b - z_a) \quad (86)$$

$$= \frac{V_b/V_\infty}{1 - V_b/V_\infty} - \frac{V_a/V_\infty}{1 - V_a/V_\infty}.$$

Recall that

$$\frac{V}{V_\infty} = \frac{pV}{pV_\infty} = m \sin \alpha, \quad (87)$$

and therefore

$$Q \cdot (z_b - z_a) = \frac{m \sin \alpha_b}{1 - m \sin \alpha_b} - \frac{m \sin \alpha_a}{1 - m \sin \alpha_a}$$

$$= \frac{1}{1 - m \sin \alpha_b} - \frac{1}{1 - m \sin \alpha_a} \quad (88)$$

$$= \frac{2m}{(1 - m \sin \alpha_a) \cdot (1 - m \sin \alpha_b)}$$

$$\times \sin \frac{\alpha_b - \alpha_a}{2} \cdot \cos \frac{\alpha_a + \alpha_b}{2}.$$

In **Figure 5** we plot the graphs for the traveltime vs. ray path arc length for the pre-critical, the critical and the post-critical rays of the Hyperbolic velocity profile.

The “trigonometric” solution for the lateral propagation and traveltime of the post-critical rays was obtained (in a different form) by Muscat (1937). However, it was not pointed out in this early study that the solution was related to the post-critical rays only, and that the other, “hyperbolic” solution exists for the pre-critical rays (and a “transient” solution for the critical rays, which are the limit case between the two basic types of rays).

Note that for the vanishing or infinitesimal parameter $Q \rightarrow 0$, the shape of the trajectory, the lateral propagation, the traveltime and the arc length of the Hyperbolic model ray path converge to the corresponding character-

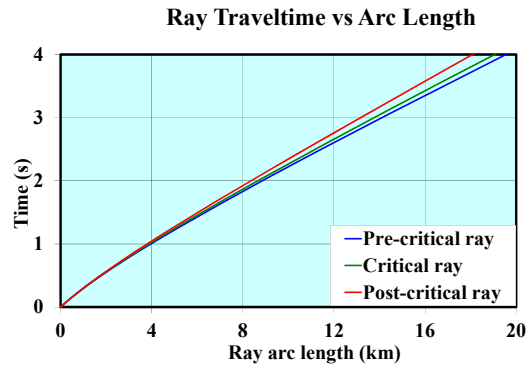


Figure 5. Traveltime vs. arc length of ray path for the three kinds of rays of the Hyperbolic velocity model: the pre-critical ray $\alpha_a = 22.5^\circ$, the critical ray $\alpha_a = \alpha_c = 30^\circ$ and the post-critical ray $\alpha_a = 37.5^\circ$.

istics of the linear velocity profile. In this case the asymptotic velocity V_∞ becomes unbounded, so that the product $R = QV_\infty$ remains a finite value and converges to a constant gradient of the linear velocity model. The eccentricity m becomes infinitesimal, and

$$\lim_{m \rightarrow 0} \frac{Q}{m} = \lim_{V_\infty \rightarrow \infty} \frac{pRV_\infty}{V_\infty} = pR. \quad (89)$$

Functions I and J from Equations (79) and (81) simplify to

$$I = \alpha_b - \alpha_a, J = \cos \alpha_b - \cos \alpha_a. \quad (90)$$

Equation (66) comes to

$$(x - x_c)^2 p^2 R^2 + \tilde{z}^2 p^2 R^2 = 1. \quad (91)$$

This is an equation for the circular arc of radius $\frac{1}{pR}$, whose center is located at $x = x_c, \tilde{z} = 0$, where p is the ray slowness. The lateral propagation, Equation (82), becomes

$$pRx = \cos \alpha_a - \cos \alpha_b. \quad (92)$$

Equation (83) yields the traveltime for this limiting case,

$$Rt_s = \ln \frac{\tan(\alpha_b/2)}{\tan(\alpha_a/2)}, \quad (93)$$

and finally, Equation (84) for the arc length converges to

$$pRs = \alpha_b - \alpha_a. \quad (94)$$

11. Full Arc of Post-Critical Ray

Consider two points on the earth surface, the transmitter and the receiver, located Δx distance apart. The goal is to trace the full arc of the post critical turning ray that connects the two points. Note that due to the symmetry of the arc, the ray angle at the destination point α_b is related to the take-off angle α_a ,

$$\alpha_b = \pi - \alpha_a. \quad (95)$$

Applying Equations (79), (81) and (82), we obtain

$$\frac{m^2}{m'^3} \operatorname{arccot} \frac{\sin \alpha_a - m}{m' \cos \alpha_a} + \frac{m \cos \alpha_a}{m'^2 (1 - m \sin \alpha_a)} = \frac{Q\Delta x}{2}, \quad (96)$$

where $m' = \sqrt{1 - m^2}$ is the conjugate eccentricity of the post-critical ray path. Recall that

$$m \sin \alpha_a = \sin \alpha_c = V_a / V_\infty. \quad (97)$$

Equation (96) simplifies to

$$\frac{m^2}{m'^3} \operatorname{arccot} \frac{\sin \alpha_c - m^2}{m' \sqrt{m^2 - \sin^2 \alpha_c}} + \frac{\sqrt{m^2 - \sin^2 \alpha_c}}{m'^2 (1 - \sin \alpha_c)} = \frac{Q\Delta x}{2}. \quad (98)$$

Equation (98) should be solved numerically for the unknown eccentricity m . To obtain the initial guess, we assume that the distance Δx is small. Then the take-off angle α_a approaches $\pi/2$, and according to Equation (97), the eccentricity exceeds the sine of the critical angle only slightly. We assume

$$m = \sin \alpha_c + \Delta m, \quad (99)$$

where Δm is a small positive value. Next we expand Equation (98) into the Taylor series and neglect the high order terms,

$$\frac{\sqrt{2 \sin \alpha_c}}{(1 - \sin \alpha_c)^2} \cdot \Delta m^{1/2} + \frac{3 + 13 \sin \alpha_c}{6 \sqrt{2 \sin \alpha_c} \cdot (1 - \sin \alpha_c)^3} \times \Delta m^{3/2} + O(\Delta m^{5/2}) = \frac{Q\Delta x}{2}. \quad (100)$$

The cubic Equation (100) has a single positive root. For example, for the velocity profile $V_a = 3 \text{ km/s}$, $k_a = 1 \text{ s}^{-1}$ and $V_\infty = 6 \text{ km/s}$, and the offset $\Delta x = 10 \text{ km}$, the critical angle becomes $\alpha_c = \pi/6$. Equation (100) leads to $\Delta m \approx 0.23$, and Equation (99) yields the initial guess $m \approx 0.73$. Solving Equation (98) with the Newton method, we obtain the eccentricity $m = 0.67638$. The take-off angle becomes $\alpha_a = 0.83193 = 47.67^\circ$. The arcs are plotted in **Figure 6** for the three asymptotically bounded velocity models. In the shallow region, the Hyperbolic model has a smaller vertical gradient (and thus, a smaller curvature) than the Conic and the EAB models, and thus, the Hyperbolic model yields a smaller take-off angle. The ray path arc of the Hyperbolic model passes above the Conic and the EAB arcs.

12. Boundary Value Ray Tracing

Given data are the departure point (x_a, z_a) and the arrival point (x_b, z_b) , and the goal is to trace the ray path. The ray path is an explicit function of the eccentricity m , and this parameter is so far unknown. Without any loss of generality, we assume here that $x_b > x_a$, i.e. that the lateral distance $\Delta x = x_b - x_a$ and the horizontal ray slowness

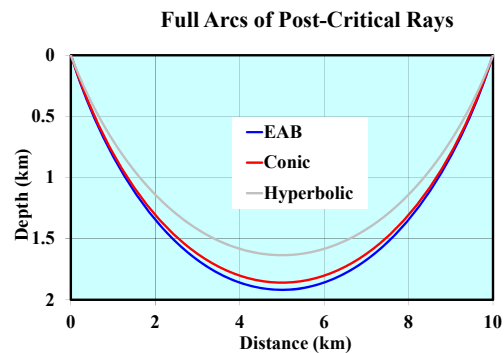


Figure 6. Full arcs of post-critical rays for the EAB, the Conic and the Hyperbolic velocity profiles.

p are positive. Assume also $z_b > z_a$, which is also not a limitation (one can reverse the endpoints otherwise). Since the ray tracing equations depend on the type of ray, we need to determine, whether the ray is pre-critical or post-critical. For this, one can plot a critical path that starts at the departure point at the critical take-off angle $\alpha_a = \alpha_c = \arcsin(V_a/V_\infty)$. If the destination point lays to the left from the critical trajectory, then the ray path is pre-critical. The ray path is post-critical if the destination point lays to the right. The critical lateral propagation Δx_c is delivered by Equation (63), which can be rearranged as

$$3Q\Delta x_c = [Q(z_b + h) - 1] \cdot \sqrt{1 + 2Q(z_b + h)} - [Q(z_a + h) - 1] \cdot \sqrt{1 + 2Q(z_a + h)}. \quad (101)$$

Given the vertical coordinates of the source and the receiver, z_a and z_b , we calculate the critical lateral propagation, and then apply the criterion

$$\begin{aligned} \Delta x < \Delta x_c & - \text{pre-critical ray,} \\ \Delta x = \Delta x_c & - \text{critical ray,} \\ \Delta x > \Delta x_c & - \text{post-critical ray.} \end{aligned} \quad (102)$$

The velocities at the end points of the trajectory $V_a(z_a)$ and $V_b(z_b)$ are known values. It follows from Snell's law that the ray angles at the end points of the trajectory are the functions of the eccentricity alone,

$$\sin \alpha_a = \frac{V_a}{mV_\infty}, \sin \alpha_b = \frac{V_b}{mV_\infty}. \quad (103)$$

Note that for the pre-critical rays and for the post critical rays before the turning point, the ray angle is acute, while for the turning rays after the turning point the ray angle is obtuse,

$$\begin{aligned} \alpha(z) &= \arcsin \frac{V(z)}{mV_\infty}, \text{ before turning point,} \\ \alpha(z) &= \pi - \arcsin \frac{V(z)}{mV_\infty}, \text{ after turning point.} \end{aligned} \quad (104)$$

Equation (82) relates the lateral propagation Δx to the ray angles at the endpoints, which, in turn, depend on the eccentricity according to Equations (103) and (104),

$$\frac{Q\Delta x}{m} = \frac{m \cdot I(m) - J(m)}{1 - m^2}, \quad (105)$$

where

$$\begin{aligned} J(m) &= \frac{2V_\infty^2}{(V_\infty - V_a) \cdot (V_\infty - V_b)} \cdot \sin \frac{\alpha_b - \alpha_a}{2} \\ &\times \left(m \cos \frac{\alpha_b - \alpha_a}{2} - \sin \frac{\alpha_a + \alpha_b}{2} \right). \end{aligned} \quad (106)$$

Function

$$I(m) = I[\alpha_a(m), \alpha_b(m)]. \quad (107)$$

is delivered by Equations (78) and (79). It was initially defined as a function of the endpoints' ray angles, but due to Equations (103) and (104) it can be considered as a function of the eccentricity alone. Next we solve nonlinear Equation (105) numerically for the unknown eccentricity m . Then the ray angles at the endpoints can be established, and the ray path can be plotted with Equation (66). Numerical examples for the boundary value ray tracing with the Hyperbolic velocity profile are presented in Appendix G.

13. Conclusion

The Hyperbolic asymptotically bounded exponential velocity model has been studied and compared to other asymptotically bounded models, in particular, the Exponential and the Conic. The forward and the inverse velocity transforms are derived. The Hyperbolic model allows a better representation of the vertical velocity variations in compacted sediments, especially in the case of thick layers. An advantage of the Hyperbolic model is that the instantaneous velocity reaches the asymptotic value in a more slow and gradual fashion, as compared to other asymptotically bounded models. Ray tracing equations have been derived. The ray trajectories, traveltimes and arc lengths have been studied analytically, and the boundary value ray tracing problem have been solved. We have tried to present a complete theory for both vertical and non-vertical rays propagating through the Hyperbolic model. Application of the Hyperbolic velocity distribution enables us to present realistic geological models using fewer parameters, as compared to the classical linear velocity function. We showed that the linear velocity function is a limiting particular case of the Hyperbolic model.

14. Acknowledgements

We are grateful to Paradigm Geophysical for the financial and technical support of this study and for the permission to publish its results.

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Appendix A. Gradient-Velocity Diagrams

In this appendix we derive the gradient-velocity diagrams for the five asymptotically bounded velocity models. In all cases we pass to a shifted frame \tilde{z} , in which the governing equations are essentially simplified. The value of the vertical shift h is different for all models. For all models the “absolute” origin corresponds to a point of maximum gradient. For all models except the Exponential slowness, this is also the point of a vanishing instantaneous velocity. As we show below, for the Exponential slowness model, the absolute origin corresponds to the half-limiting velocity $V(\tilde{z}) = V_\infty/2$.

A.1. The Hyperbolic Muscat Model

The velocity profile is given by

$$V(z) = V_a + \Delta V \cdot \left(1 - \frac{\Delta V}{\Delta V + k_a z} \right), \Delta V = V_\infty - V_a. \quad (\text{A-1})$$

Establish the depth level where the velocity vanishes. This point is located above the earth surface,

$$z = -h, h = \frac{V_a \Delta V}{k_a V_\infty}. \quad (\text{A-2})$$

Introduce the shifted frame, $\tilde{z} = z + h$. The velocity profile becomes

$$\frac{V(\tilde{z})}{V_\infty} = \frac{k_a V_\infty \tilde{z}}{\Delta V^2 + k_a V_\infty \tilde{z}}. \quad (\text{A-3})$$

The vertical gradient is

$$\frac{k(\tilde{z})}{V_\infty} = \frac{k_a \Delta V^2 V_\infty}{(\Delta V^2 + k_a V_\infty \tilde{z})^2}. \quad (\text{A-4})$$

At the absolute origin $\tilde{z} = 0$, the gradient is maximal,

$$k_{\max} = k_a V_\infty^2 / \Delta V^2. \quad (\text{A-5})$$

Introduction of Equation (A-5) into (A-3) and (A-4) leads to

$$\frac{V(\tilde{z})}{V_\infty} = \frac{k_{\max} \tilde{z}}{V_\infty + k_{\max} \tilde{z}}, \frac{k(\tilde{z})}{k_{\max}} = \frac{V_\infty^2}{(V_\infty + k_{\max} \tilde{z})^2}. \quad (\text{A-6})$$

Finally, elimination of depth \tilde{z} from the two equations of Equation Set (A-6) results in

$$\sqrt{\frac{k}{k_{\max}}} + \frac{V}{V_\infty} = 1. \quad (\text{A-7})$$

A.2. The Exponential Muscat Model

The velocity model is described by

$$V(z)/V_\infty = \tanh(A + z/z_o). \quad (\text{A-8})$$

The vertical shift is $h = Az_o$, and in the shifted frame,

$\tilde{z} = z + h$, Equation (A-8) simplifies to

$$V(\tilde{z})/V_\infty = \tanh(\tilde{z}/z_o). \quad (\text{A-9})$$

The velocity gradient is

$$\frac{k(\tilde{z})}{V_\infty} = \frac{1}{z_o \cosh^2(\tilde{z}/z_o)}, \quad (\text{A-10})$$

with $k_{\max} = V_\infty/z_o$, so that

$$\frac{k(\tilde{z})}{k_{\max}} = \frac{1}{\cosh^2(\tilde{z}/z_o)}. \quad (\text{A-11})$$

Eliminate the absolute depth from Equations (A-9) and (A-11) and obtain the gradient-velocity relationship,

$$\frac{k}{k_{\max}} + \left(\frac{V}{V_\infty} \right)^2 = 1. \quad (\text{A-12})$$

A.3. The Exponential Slowness Model

The profile equation reads

$$V(z) = \frac{V_a V_\infty}{V_a + \Delta V \exp(-z/z_o)}, z_o = \frac{V_a \Delta V}{k_a V_\infty}. \quad (\text{A-13})$$

Unlike the other asymptotically bounded models mentioned in the introduction, the velocity in the Exponential slowness model does not vanish at a finite negative depth. The velocity vanishes at $z \rightarrow -\infty$ and approaches to the asymptotic value V_∞ at $z \rightarrow +\infty$. The gradient of the velocity is

$$k(z) = \frac{V_a \Delta V V_\infty}{z_o} \cdot \frac{\exp(-z/z_o)}{[V_a + \Delta V \exp(-z/z_o)]^2}. \quad (\text{A-14})$$

The gradient $k(z)$ vanishes at both remote ends, $z \rightarrow \pm\infty$, and $k(z)$ has a single critical point: the maximum of the gradient occurs at $z = z^*$,

$$z^* = z_o \ln \frac{\Delta V}{V_a}. \quad (\text{A-15})$$

We emphasize that the logarithm in Equation (A-15) may prove to be both positive and negative. At the point $z = z^*$, the maximum gradient and the velocity are

$$k_{\max} = \frac{V_\infty}{4z_o}, V^* = \frac{V_\infty}{2}. \quad (\text{A-16})$$

Note that in case when the depth of the maximum gradient is positive, $z^* \geq 0$, this point really exists underground, and the gradient first increases, then accepts the maximum value

$$k_{\max} = \frac{k_a}{4} \cdot \frac{V_\infty^2}{V_a \Delta V}. \quad (\text{A-17})$$

Below this point, the gradient begins to decay, and eventually vanishes at the infinite depth. In case when

the depth z^* of the maximum gradient is negative, this point is above the earth surface, and throughout the whole depth range $0 \leq z < \infty$ the velocity gradient $k(z)$ is actually a monotonously decreasing function. Note that

$$z^* \geq 0 \rightarrow \Delta V = V_\infty - V_a \geq V_a \text{ or } V_a \leq V_\infty/2. \quad (\text{A-18})$$

Next we assume the shift $h = -z^*$, and pass to the shifted frame, $\tilde{z} = z + h$. The gradient accepts now a maximum value at the origin. Rearrange Equation (A-13),

$$\begin{aligned} \frac{V(\tilde{z})}{V_\infty} &= \frac{1}{1 + \exp(-\tilde{z}/z_o)}, \\ k(\tilde{z}) &= \frac{4 \exp(-\tilde{z}/z_o)}{[1 - \exp(-\tilde{z}/z_o)]^2}. \end{aligned} \quad (\text{A-19})$$

Note that the velocity profile in Equation (A-19) can be set in an alternative way,

$$\begin{aligned} \frac{V(\tilde{z})}{V_\infty} &= \frac{\exp(a\tilde{z})}{\exp(a\tilde{z}) + \exp(-a\tilde{z})} \\ &= \frac{\cosh(a\tilde{z}) + \sinh(a\tilde{z})}{2 \cosh(a\tilde{z})} = \frac{1 + \tanh(a\tilde{z})}{2} \end{aligned} \quad (\text{A-20})$$

where

$$a = \frac{1}{2z_o} = \frac{2k_{\max}}{V_\infty}, \quad (\text{A-21})$$

so that

$$\begin{aligned} \frac{V(\tilde{z})}{V_\infty} &= \frac{1}{2} \cdot \left(1 + \tanh \frac{2k_{\max} \tilde{z}}{V_\infty} \right), \\ \frac{k(\tilde{z})}{k_{\max}} &= \cosh^{-2} \left(\frac{2k_{\max} \tilde{z}}{V_\infty} \right). \end{aligned} \quad (\text{A-22})$$

Finally, we eliminate depth from Equation (A-22) and obtain

$$\frac{k}{4k_{\max}} = \frac{V}{V_\infty} \cdot \left(1 - \frac{V}{V_\infty} \right). \quad (\text{A-23})$$

A.4. The Exponential Asymptotically Bounded Velocity Model (EAB)

In this case the velocity profile is

$$V(z) = V_a + \Delta V \cdot \left[1 - \exp\left(-\frac{k_a z}{\Delta V}\right) \right]. \quad (\text{A-23})$$

The velocity vanishes above the earth surface at $z = -h$,

$$h = \frac{\Delta V}{k_a} \cdot \ln \frac{V_\infty}{\Delta V}. \quad (\text{A-24})$$

In the shifted frame $\tilde{z} = z + h$, the velocity profile

simplifies to

$$\begin{aligned} \frac{V(\tilde{z})}{V_\infty} &= 1 - \exp\left(-\frac{k_a \tilde{z}}{\Delta V}\right), \\ k(\tilde{z}) &= \frac{k_a \cdot V_\infty}{\Delta V} \cdot \exp\left(-\frac{k_a \tilde{z}}{\Delta V}\right). \end{aligned} \quad (\text{A-25})$$

At the shifted origin the gradient is maximal,

$$k_{\max} = \frac{k_a \cdot V_\infty}{\Delta V}. \quad (\text{A-26})$$

Introduce Equation (A-26) into (A-25),

$$\begin{aligned} \frac{V(\tilde{z})}{V_\infty} &= 1 - \exp\left(-\frac{k_{\max} \tilde{z}}{V_\infty}\right), \\ \frac{k(\tilde{z})}{k_{\max}} &= \exp\left(-\frac{k_{\max} \tilde{z}}{V_\infty}\right). \end{aligned} \quad (\text{A-27})$$

Finally, we eliminate depth from Equation (A-27) and obtain the governing equation,

$$\frac{k}{k_{\max}} + \frac{V}{V_\infty} = 1. \quad (\text{A-28})$$

Note that only for the EAB velocity model the diagram Equation (A-28) is linear.

A.5. Conic Velocity Model

For the Conic profile, the velocity and its gradient in the absolute frame are given by

$$\begin{aligned} V(\tilde{z}) &= \frac{R\tilde{z}}{\sqrt{1 + Q^2 \tilde{z}^2}}, \\ k(\tilde{z}) &= \frac{R}{(1 + Q^2 \tilde{z}^2)^{3/2}}. \end{aligned} \quad (\text{A-29})$$

This equation can be rearranged as

$$\begin{aligned} \frac{V(\tilde{z})}{V_\infty} &= \frac{k_{\max} \tilde{z}}{\sqrt{V_\infty^2 + k_{\max}^2 \tilde{z}^2}}, \\ \frac{k(\tilde{z})}{k_{\max}} &= \frac{V_\infty^3}{(V_\infty^2 + k_{\max}^2 \tilde{z}^2)^{3/2}}. \end{aligned} \quad (\text{A-30})$$

Next we eliminate the absolute depth \tilde{z} from Equation (A-29) and get the governing differential equation of the Conic velocity model,

$$\left(\frac{k}{k_{\max}} \right)^{2/3} + \left(\frac{V}{V_\infty} \right)^2 = 1. \quad (\text{A-31})$$

The Conic velocity profile, Equation (A-30), can be also set in an equivalent form, through a hyperbolic and an inverse hyperbolic function,

$$\frac{V(\tilde{z})}{V_\infty} = \tanh \operatorname{arcsinh} \frac{k_{\max} \tilde{z}}{V_\infty}. \quad (\text{A-32})$$

A.6. Comments on Diagrams

Summarize the gradient-velocity diagrams for the five asymptotically bounded velocity models. The governing differential equations are

$$\begin{aligned}
 \sqrt{\frac{k}{k_{\max}} + \frac{V}{V_{\infty}}} &= 1 && \text{Muscat Hyperbolic Model,} \\
 \frac{k}{k_{\max}} + \left(\frac{V}{V_{\infty}}\right)^2 &= 1 && \text{Muscat Exponential Model,} \\
 \frac{k}{4k_{\max}} = \frac{V}{V_{\infty}} \cdot \left(1 - \frac{V}{V_{\infty}}\right) &&& \text{Exponential Slowness Model,} \\
 \frac{k}{k_{\max}} + \frac{V}{V_{\infty}} &= 1 && \text{EAB Velocity Model,} \\
 \left(\frac{k}{k_{\max}}\right)^{2/3} + \left(\frac{V}{V_{\infty}}\right)^2 &= 1 && \text{Conic Velocity Model.}
 \end{aligned}
 \tag{A-33}$$

The gradient-velocity diagrams for the five asymptotically bounded models are plotted in **Figure 7**. Note that in the original frame of reference the asymptotically bounded models are described by the three parameters. As we mentioned, in the shifted frame where the velocity vanishes at the origin (or the vertical gradient accepts a maximum value at the origin), only two parameters are needed. These two parameters may be the maximum gradient k_{\max} and the asymptotic velocity V_{∞} as in Equation Set (A-33). The constant value that appears upon the integration of each equation is not a new parameter as it should be adjusted to match the maximum values k_{\max} and V_{∞} .

We emphasize that only for the EAB model the gradient-velocity relationship is linear: Derivative of an exponent is proportional to the same exponent.

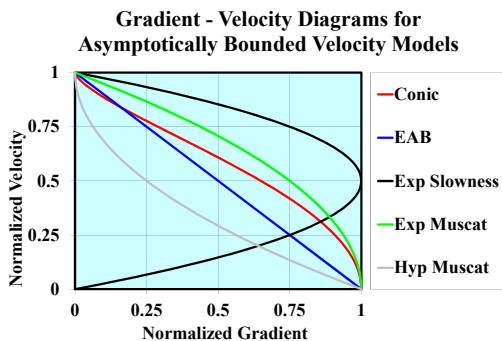


Figure 7. Diagrams “Gradient-Velocity” for asymptotically bounded velocity models. Only for the EAB model the diagram is linear. For all models, except the Exponential slowness model, the gradient decreases with the increase of velocity. For the Exponential slowness model the gradient reaches maximum when the velocity becomes one half of the asymptotic value. Note the central symmetry between the two Muscat models, Hyperbolic and Exponential.

For all models except the Exponential slowness [8], the gradient decreases with the increase of velocity (and depth). In case of the Exponential slowness, the vertical gradient increases along with the velocity, until the velocity reaches one half of the asymptotic value $V_{\infty}/2$. At this point the gradient reaches its maximum value k_{\max} and then begins to decrease with depth. The point of maximum gradient in the Exponential slowness model may really exist in the subsurface, or it may be an imaginary point located *above* the earth surface (or above the upper interface of a layer). This point is real in case when the initial velocity V_a does not exceed the half-limiting value, $V_a \leq V_{\infty}/2$. Furthermore, we comment on the special central symmetry between the two Muscat (1937) models: Hyperbolic (*HM*) and Exponential (*EM*), see Equation (A-33) and the two corresponding plots in **Figure 7**,

$$F_{HM} \left(\frac{k}{k_{\max}}, \frac{V}{V_{\infty}} \right) = 1 \rightarrow F_{EM} \left(1 - \frac{k}{k_{\max}}, 1 - \frac{V}{V_{\infty}} \right) = 1. \tag{A-34}$$

where F_{HM} and F_{EM} are the corresponding gradient-velocity functions in Equation (A-33) for these two models. Although these two models are described by essentially different vertical velocity profiles, there is an apparent similarity in the gradient-velocity diagrams.

Appendix B. Lambert Function

The Lambert function [17] $y = L(x)$ delivers the solution of the transcendent equation

$$x = y \cdot \exp y. \tag{B-1}$$

Its graph is plotted in **Figure 8** and consists of two branches: branch zero $L_0(x)$ and branch minus one $L_{-1}(x)$. The argument range is $-\exp(-1) \leq x < \infty$ for branch zero and $-\exp(-1) \leq x < 0$ for branch minus one. The value range is $-1 \leq y < \infty$ for branch zero and $-\infty \leq y \leq -1$ for branch minus one. In particular, this

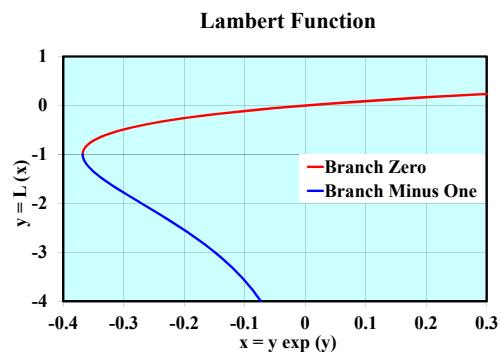


Figure 8. Lambert function $y = L(x)$ is the solution of the transcendent equation $y \cdot \exp(y) = x$ for a given value x and an unknown value y . The function has two branches. Branch zero is plotted in red, and branch minus one is plotted in blue.

means that for a positive argument x only branch zero exists, while for a negative argument both branches do exist. Therefore, in the latter case, the branch index should be specified to avoid ambiguity. The derivative of the Lambert function is

$$\frac{dL}{dx} = \frac{1}{x} \cdot \frac{L(x)}{1+L(x)} \text{ for } x \neq 0, \quad (\text{B-2})$$

and for the infinitesimal argument

$$\lim_{x \rightarrow 0} \frac{dL(x)}{dx} = 1. \quad (\text{B-3})$$

Comment. A general comment is related to Appendices C to F. The transform equations are formulated in the dimensionless form, with the unknown top and bottom asymptotic factors, M_a and M_b . After the transform equation or equation set is resolved, we apply Equation (27) to find the top and bottom instantaneous velocities, V_a and V_b . If the transform is formulated in depth (*i.e.*, the interval thickness Δz is specified), we apply Equation (31) to find the top and bottom gradients of velocity, k_a and k_b . If the transform is formulated in time (*i.e.*, the interval travelttime Δt is specified), then we first apply Equation (37) to establish the interval thickness Δz , and then Equation (31) to find the top and bottom gradients.

Appendix C. Swapping Interfaces

In this appendix, we find the instantaneous velocity and its gradient at the bottom interface, given these parameters at the top interface, and vice versa, and consider these problems both vs. depth and vs. time.

Problem C1. Given the velocity V_a and its gradient k_a at the top interface and the layer thickness Δz , one can establish the corresponding parameters at the bottom interface. For this, we calculate the top asymptotic factor M_a with the first equation of Equation Set (26). The bottom asymptotic factor M_b can be obtained with the first equation of Equation Set (32).

Problem C2. The instantaneous velocity and its gradient are specified at the bottom interface and the interval thickness Δz is given. Velocity and gradient should be found at the top interface. Thus, parameters V_b, k_b are given, and parameters V_a, k_a are to be found. In this case calculate the bottom asymptotic factor M_b with the second equation of Equation Set (26) and apply the second equation of Equation Set (32) to get the top get M_a .

Problem C3. The velocity and its gradient at the top interface, V_a and k_a are given, the interval travelttime Δt is known, and the bottom interface parameters V_b and k_b should be found. Combining the first equation of Equation Set (32) and Equation (36), we eliminate the interval thickness Δz and obtain

$$M_b - M_a + \ln \frac{M_b - 1}{M_a - 1} = M_a^2 k_a \Delta t, \quad (\text{C-1})$$

where the top asymptotic factor M_a is known from Equation (26). Equation (C-1) should be solved for the unknown bottom asymptotic factor M_b . We use a cubic approximation for Equation (C-1) to get the initial guess, assuming the increment of the asymptotic factor $\Delta M = M_b - M_a$ is small,

$$\frac{\Delta M^3}{3(M_a - 1)^2} - \frac{\Delta M^2}{2(M_a - 1)} + \frac{M_a \Delta M}{M_a - 1} = M_a^2 k_a \Delta t. \quad (\text{C-2})$$

Problem C4. The velocity and its gradient are specified at the bottom interface, along with the interval travelttime, and the profile parameters should be found at the top interface. Thus, parameters V_b, k_b and Δt are given, while parameters V_a and k_a are to be established. For this, we combine the second equation of Equation Set (32) and Equation (36),

$$M_b - M_a + \ln \frac{M_b - 1}{M_a - 1} = M_b^2 k_b \Delta t. \quad (\text{C-3})$$

The bottom asymptotic factor M_b is known from Equation (26), and Equation (C-3) should be solved for the unknown top asymptotic factor M_a . The initial guess for $\Delta M = M_b - M_a$ can be found from

$$\frac{\Delta M^3}{3(M_b - 1)^2} + \frac{\Delta M^2}{2(M_b - 1)} + \frac{M_b \Delta M}{M_b - 1} = M_b^2 k_b \Delta t. \quad (\text{C-4})$$

Appendix D. Inversion with Instantaneous Velocity

In this appendix we consider inversion problems that do not involve the effective models (average and RMS velocity). In this inversion group, one or both parameters at the top interface, V_a and k_a are unknown, with the other data given instead. The group includes four problems vs. depth (interval thickness) and four similar problems vs. interval travelttime.

Problem D1. Instantaneous velocities vs. depth. Given data are the top and bottom interface instantaneous velocities, V_a and V_b , the asymptotic velocity V_∞ and the interval thickness Δz . Find the top interface gradient k_a . *Solution.* Apply Equation (26) to calculate the top and bottom asymptotic factors, M_a and M_b .

Problem D2. Instantaneous velocities vs. time. Given data are the top and bottom interface instantaneous velocities, V_a and V_b , the asymptotic velocity V_∞ and the interval travelttime Δt . Find the top interface gradient k_a . *Solution.* Apply Equation (26) to calculate the top and bottom asymptotic factors, M_a and M_b .

Problem D3. Gradients vs. depth. Given data are the top and bottom interface vertical gradients, k_a and k_b ,

the asymptotic velocity V_∞ and the interval thickness Δz . Find the top interface instantaneous velocity V_a . *Solution.* Solve Equation Set (32) for the unknown asymptotic factors, M_a and M_b ,

$$M_a = A\sqrt{k_b}, M_b = A\sqrt{k_a}, \quad (D-1)$$

where $A \equiv \frac{V_\infty}{\Delta z} \cdot \frac{\sqrt{k_a} - \sqrt{k_b}}{k_a k_b}$.

Problem D4. Gradients vs. time. Given data are the top and bottom interface vertical gradients, k_a and k_b , the asymptotic velocity V_∞ and the interval traveltime Δt . Find the top interface instantaneous velocity V_a . *Solution.* Introduce solution (D-1) into the traveltime Equation (36). This leads to a nonlinear equation vs. the unknown interval thickness Δz

$$V_\infty \Delta z + \frac{k_a k_b \Delta z^2}{B^2} \cdot \ln \frac{BV_\infty \sqrt{k_a} - k_a k_b \Delta z}{BV_\infty \sqrt{k_b} - k_a k_b \Delta z} = V_\infty^2 \Delta t, \quad (D-2)$$

$$\text{where } B = \sqrt{k_a} - \sqrt{k_b}.$$

Equation (D-2) should be solved numerically. It is suitable to normalize the gradients and the interval thickness,

$$\tilde{k}_a = k_a \Delta t, \tilde{k}_b = k_b \Delta t, \Delta \tilde{z} = \frac{\Delta z}{V_\infty \Delta t} = \frac{V_{\text{int}}}{V_\infty} < 1. \quad (D-3)$$

The normalized equation becomes

$$\Delta \tilde{z} + \frac{\tilde{k}_a \tilde{k}_b \Delta \tilde{z}^2}{\tilde{B}^2} \cdot \ln \frac{\tilde{B} V_\infty \sqrt{\tilde{k}_a} - \tilde{k}_a \tilde{k}_b \Delta \tilde{z}}{\tilde{B} V_\infty \sqrt{\tilde{k}_b} - \tilde{k}_a \tilde{k}_b \Delta \tilde{z}} = 1, \quad (D-4)$$

$$\text{where } \tilde{B} = \sqrt{\tilde{k}_a} - \sqrt{\tilde{k}_b}.$$

To get an initial guess, we expand Equation (D-4) into a power series and obtain a cubic approximation,

$$\frac{\tilde{k}_a^{3/2} \tilde{k}_b^{3/2}}{\tilde{B}^2} \cdot \Delta \tilde{z}^3 + \frac{\tilde{k}_a \tilde{k}_b}{2\tilde{B}^2} \cdot \ln \frac{\tilde{k}_a}{\tilde{k}_b} \cdot \Delta \tilde{z}^2 + \Delta \tilde{z} = 1. \quad (D-5)$$

After the interval thickness is found, apply solution (D-1).

Problem D5. Velocity and gradient vs. depth at different interfaces. Given data are the top interface velocity V_a , the bottom interface vertical gradient k_b , the asymptotic velocity V_∞ and the interval thickness Δz . Find the top interface gradient k_a . *Solution.* Apply Equation (26) to calculate the top asymptotic factor M_a . Introduce the normalized bottom gradient,

$$\hat{k}_b = k_b \Delta z / V_\infty. \quad (D-6)$$

With this notation, the second equation of Equation Set (32) becomes

$$\hat{k}_b M_b^2 - M_b + M_a = 0. \quad (D-7)$$

Taking into account that for an interval of vanishing

thickness ($\Delta z \rightarrow 0$) the top and bottom asymptotic factors coincide ($M_b \rightarrow M_a$), one can establish the single physical root of the quadratic equation,

$$M_b = \frac{1 - \sqrt{1 - 4M_a \hat{k}_b}}{2\hat{k}_b}. \quad (D-8)$$

Problem D6. Velocity and gradient vs. depth at different interfaces. Given data are the top interface gradient k_a , the bottom interface instantaneous velocity V_b , the asymptotic velocity V_∞ and the interval thickness Δz . Find the top interface velocity V_a . *Solution.* Calculate the bottom asymptotic factor M_b with equation 26. Normalize the top gradient,

$$\hat{k}_a = k_a \Delta z / V_\infty. \quad (D-9)$$

Get the top asymptotic factor

$$M_a = \frac{\sqrt{1 + 4M_b \hat{k}_a} - 1}{2\hat{k}_a}. \quad (D-10)$$

Problem D7. Velocity and gradient vs. time at different interfaces. Given data are the top interface velocity V_a , the bottom interface vertical gradient k_b , the asymptotic velocity V_∞ and the interval traveltime Δt . Find the top interface gradient k_a . *Solution.* Apply Equation (26) to calculate the top asymptotic factor M_a . Then solve the nonlinear Equation (C-3) for the unknown bottom asymptotic factor M_b . To get the initial guess, we assume $M_b = M_a + \Delta M$. Expand Equation (C-3) for a small increment of the asymptotic factor ΔM to get a cubic approximation,

$$\frac{\Delta M^3}{3(M_a - 1)^3} - \frac{1 + 2k_b \Delta t (M_a - 1)^2}{2(M_a - 1)^2} \cdot \Delta M^2 + \frac{[1 - 2k_b \Delta t (M_a - 1)] M_a}{M_a - 1} \cdot \Delta M = k_b \Delta t M_a^2. \quad (D-11)$$

Problem D8. Given data are the top interface gradient k_a , the bottom interface instantaneous velocity V_b , the asymptotic velocity V_∞ and the interval traveltime Δt . Find the top interface velocity V_a . *Solution.* Apply Equation (26) to calculate the bottom asymptotic factor M_b . To calculate the unknown top asymptotic factor, we solve the nonlinear Equation (C-1). To get the initial guess, assume $M_a = M_b - \Delta M$. This leads to

$$\frac{\Delta M^3}{3(M_b - 1)^3} + \frac{1 - 2k_a \Delta t (M_b - 1)^2}{2(M_b - 1)^2} \cdot \Delta M^2 + \frac{[1 + 2k_a \Delta t (M_b - 1)] M_b}{M_b - 1} \cdot \Delta M = k_a \Delta t M_b^2. \quad (D-12)$$

Problem D9. Given data are the instantaneous velocity V_b at the bottom interface, and at an intermediate

level inside the interval, V_c , and the asymptotic velocity V_∞ . Two vertical distances are specified: the full interval thickness Δz_b (the distance between the top and bottom interfaces) and the partial interval thickness Δz_c (the distance between the top interface and the intermediate level). Find the top interface velocity and gradient, V_a and k_a . *Solution.* Use Equation (26) to calculate the bottom asymptotic factor M_b and the intermediate asymptotic factor M_c . Apply Equation (29) for the full interval and for the partial interval,

$$Q = \frac{M_b - M_a}{\Delta z_b} = \frac{M_c - M_a}{\Delta z_c}. \quad (\text{D-13})$$

Solve Equation (D-13) for the top asymptotic factor M_a ,

$$M_a = \frac{M_c \Delta z_b - M_b \Delta z_c}{\Delta z_b - \Delta z_c}. \quad (\text{D-14})$$

Problem D10. Given data are the instantaneous velocity V_b at the bottom interface, and at an intermediate level V_c inside the interval, and the asymptotic velocity V_∞ . Two vertical traveltimes are specified: the full interval traveltime Δt_b (the traveltime between the top and bottom interfaces) and the partial traveltime Δt_c (the traveltime between the top interface and the intermediate level). Find the top interface velocity and gradient, V_a and k_a . *Solution.* Use Equation (26) to calculate the bottom asymptotic factor M_b and the intermediate asymptotic factor M_c . Apply Equation (28) for the full interval and for the partial interval, parameter R becomes,

$$\frac{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}}{\Delta t_b} = \frac{M_c - M_a + \ln \frac{M_c - 1}{M_a - 1}}{\Delta t_c}. \quad (\text{D-15})$$

Introduce parameter $m_i = M_i - 1$, Equation (D-15) becomes

$$\frac{m_b - m_a + \ln(m_b/m_a)}{\Delta t_b} = \frac{m_c - m_a + \ln(m_c/m_a)}{\Delta t_c}. \quad (\text{D-16})$$

Equation (D-16) can be solved with the Lambert function, branch zero,

$$m_a = L_0 \left[\exp(m_c f_b - m_b f_c) \cdot m_b^{-f_c} \cdot m_c^{f_b} \right], \quad (\text{D-17})$$

$$f_b \equiv \frac{\Delta t_b}{\Delta t_b - \Delta t_c} \text{ and } f_c \equiv \frac{\Delta t_c}{\Delta t_b - \Delta t_c}.$$

Appendix E. Two-Point Effective Model Inversion

In this appendix we consider the two-point (single-interval) inversion where the RMS velocity is specified at the interfaces of an interval vs. depth or traveltime, or

alternatively, depth is specified vs. traveltime instead of the RMS velocity. One of the two parameters at the top interface—either the top velocity V_a , or the top gradient k_a —is a known value, while the other one is unknown and should be established. In all cases, the asymptotic velocity V_∞ and the top interface absolute traveltime t_a are assumed known values. We consider also a special case when the RMS velocity is specified vs. both depth and time, and both parameters, V_a and k_a , are unknown.

Problem E1. RMS vs. depth with unknown gradient. Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, the interval thickness Δz and the top interface velocity V_a . Find the top gradient k_a . *Solution.* It follows from the definition of the hyperbolic parameter, Equation (44),

$$W = V_{2,b}^2 \cdot t_b - V_{2,a}^2 \cdot t_a, t_b = t_a + \Delta t. \quad (\text{E-1})$$

Recall that t_a and t_b are one-way absolute top and bottom interface traveltimes (measured from the earth surface), Δt is the one-way interval traveltime, and W is the hyperbolic parameter through the interval. Equation (E-1) can be arranged as

$$W - V_{2,b}^2 \cdot \Delta t = (V_{2,b}^2 - V_{2,a}^2) \cdot t_a. \quad (\text{E-2})$$

Introduction of Equation (37) for the traveltime Δt and Equation (47) for the hyperbolic parameter W into Equation (E-2) results in

$$\frac{1}{M_b - M_a} \cdot \ln \frac{M_b}{M_a} + \frac{B}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} = A, \quad (\text{E-3})$$

where A and B are known dimensionless parameters,

$$A = 1 - B - \frac{V_{2,b}^2 - V_{2,a}^2}{V_\infty \Delta z} \cdot t_a, B = V_{2,b}^2 / V_\infty^2. \quad (\text{E-4})$$

The top asymptotic factor M_a is delivered by Equation (26), and the nonlinear Equation (E-4) should be solved for the unknown bottom asymptotic factor M_b . To get the initial guess, we assume a small increment of the asymptotic factor on the interval, $\Delta M = M_b - M_a$, and expand Equation (E-3) into a power series. The cubic approximation reads

$$C_3 \cdot \Delta M^3 + C_2 \cdot \Delta M^2 + C_1 \cdot \Delta M + C_0 = A, \quad (\text{E-5})$$

where

$$C_i = \frac{(-1)^i}{i+1} \cdot \left[\frac{B}{(M_a - 1)^{i+1}} + \frac{1}{M_a^{i+1}} \right], i = 0, 1, \dots \quad (\text{E-6})$$

Equation (E-3) should be solved for M_b .

Problem E2. RMS vs. depth with unknown velocity. Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, the interval thickness Δz and the top interface gradient k_a . Find the top interface

velocity V_a . *Solution.* Use Equation (E-3) and the first equation of Equation Set (32),

$$\frac{1}{M_b - M_a} \cdot \ln \frac{M_b}{M_a} + \frac{B}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} = A, \quad (\text{E-7})$$

$$M_b = M_a + M_a^2 \cdot \frac{k_a \Delta z}{V_\infty}.$$

We solve Equation Set (E-7) for the unknown asymptotic factors M_a and M_b . To obtain the initial guess for M_a , assume that the increment of the asymptotic factor ΔM is small, and linearize Equation Set (E-7). This leads to

$$\begin{aligned} (2A + B + \hat{k}_a \cdot B) \cdot M_a^3 - 2 \cdot (2A + B + \hat{k}_a + 1) \cdot M_a^2 \\ + (2A + 2B + \hat{k}_a + 4) \cdot M_a = 2. \end{aligned} \quad (\text{E-8})$$

Problem E3. RMS vs. time with unknown gradient. Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, the interval traveltime Δt and the top interface velocity V_a . Find the top gradient k_a . *Solution.* First we apply Equation (E-1) and calculate the hyperbolic parameter W . At this time, W is a known value. Next we apply Equation (48),

$$\frac{M_b - M_a - \ln \frac{M_b}{M_a}}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}} = C, \quad (\text{E-9})$$

where C is a known dimensionless parameter, the normalized local RMS velocity,

$$C = \frac{U^2}{V_\infty^2} = \frac{W}{V_\infty^2 \Delta t}, \quad (\text{E-10})$$

and U is the non-normalized local RMS velocity on the interval. The top asymptotic factor M_a is delivered by Equation (26), and the bottom asymptotic factor M_b is established from Equation (E-9). To solve this nonlinear equation, an initial guess is needed. Assume that the increment of the asymptotic factor ΔM on the interval is small, and expand Equation (E-9) into a power series. The cubic approximation reads

$$\begin{aligned} \frac{3M_a - 4}{6M_a^5} \cdot \Delta M^3 - \frac{4M_a - 5}{6M_a^4} \cdot \Delta M^2 \\ + \frac{M_a - 1}{M_a^3} \cdot \Delta M + \frac{(M_a - 1)^2}{M_a^2} = C. \end{aligned} \quad (\text{E-11})$$

Problem E4. RMS vs. time with unknown velocity. Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, the interval traveltime Δt and the top interface gradient k_a . Find the top interface velocity V_a . *Solution.* We solve a set consisting of two equations: the first is Equation (E-9) combined with

Equation (C-1), and the second is (C-1) itself,

$$\begin{aligned} \frac{M_b - M_a}{M_a^2} - \frac{1}{M_a^2} \cdot \ln \frac{M_b}{M_a} = C k_a \Delta t, \\ \frac{M_b - M_a}{M_a^2} + \frac{1}{M_a^2} \cdot \ln \frac{M_b - 1}{M_a - 1} = k_a \Delta t. \end{aligned} \quad (\text{E-12})$$

To get the initial guess, we linearize Equation Set (E-12) for a small increment of the asymptotic factor ΔM and obtain

$$\begin{aligned} M_a = \frac{V_\infty}{V_\infty - U}, \\ \Delta M = \frac{k \Delta t U V_\infty}{(V_\infty - U)^2}. \end{aligned} \quad (\text{E-13})$$

Problem E5. Depth vs. time with unknown gradient. Given data are the interval thickness Δz , the traveltime Δt and the top interface velocity V_a . Find the top gradient k_a . *Solution.* Use Equation (26) to find the top asymptotic factor M_a . Next we apply Equation (37) to establish the bottom asymptotic factor M_b

$$\frac{M_b - M_a}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}} = D, \quad (\text{E-14})$$

where D is a known dimensionless parameter—the normalized interval velocity,

$$D = \frac{V_{\text{int}}}{V_\infty} = \frac{\Delta z}{V_\infty \Delta t}. \quad (\text{E-15})$$

To get the initial guess for Equation (E-14), we expand it into a power series for a small increment of the asymptotic factor ΔM ,

$$\begin{aligned} \frac{6M_a^2 - 8M_a + 3}{24M_a^4 (M_a - 1)^2} \cdot \Delta M^3 - \frac{4M_a - 3}{12M_a^3 (M_a - 1)} \cdot \Delta M^2 \\ + \frac{\Delta M}{2M_a^2} + \frac{M_a - 1}{M_a} = D. \end{aligned} \quad (\text{E-16})$$

Equation (E-14) should be solved for the bottom asymptotic factor M_b .

Problem E6. Depth vs. time with unknown velocity. Given data are the interval depth Δz , the interval traveltime Δt and the top interface gradient k_a . Find the top interface velocity V_a . *Solution.* Apply Equation (E-14) and the first equation of Equation Set (32),

$$\begin{aligned} \frac{1}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} = \frac{1 - D}{D}, \\ M_b - M_a = M_a^2 \cdot \frac{k_a \Delta z}{V_\infty}. \end{aligned} \quad (\text{E-17})$$

Next we solve Equation Set (E-17) for the unknown asymptotic factors M_a and M_b . To obtain the initial

guess, we linearize Equation Set (E-17) for a small increment of the asymptotic factor ΔM , and consider the first equation of the set,

$$\lim_{M_b \rightarrow M_a} \left(\frac{1}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} \right) = \frac{1}{M_a - 1}. \quad (\text{E-18})$$

Hence we obtain the initial guess for the top asymptotic factor M_a ,

$$M_a = \frac{1}{1 - D} = \frac{V_\infty \Delta t}{V_\infty \Delta t - \Delta z}. \quad (\text{E-19})$$

Problem E7. RMS vs. depth and time with unknown velocity and gradient. Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, the interval traveltime Δt and the interval thickness Δz . Find the top interface velocity V_a and gradient k_a . *Solution.* The resolving set follows from Equations (E-9) and (E-14),

$$\begin{aligned} \frac{1}{M_b - M_a} \cdot \ln \frac{M_b}{M_a} &= A, \\ \frac{1}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} &= B. \end{aligned} \quad (\text{E-20})$$

where A and B are known dimensionless parameters,

$$A = \frac{V_{\text{int}} \cdot V_\infty - U^2}{V_{\text{int}} \cdot V_\infty}, B = \frac{V_\infty - V_{\text{int}}}{V_{\text{int}}}. \quad (\text{E-21})$$

To obtain the initial guess, we assume that the top and bottom values of the asymptotic factor, M_a and M_b , are close, and solve the two equations of Equation Set (E-20) apart. Each equation yields a root; the smaller root is the top asymptotic factor M_a , and the larger root is the bottom asymptotic factor M_b . The initial guess becomes

$$\begin{aligned} M_a &= \frac{1}{B} + 1 = \frac{V_\infty}{V_\infty - V_{\text{int}}}, \\ M_b &= \frac{1}{A} = \frac{V_{\text{int}} V_\infty}{V_{\text{int}} V_\infty - U^2} \rightarrow M_b \geq M_a. \end{aligned} \quad (\text{E-22})$$

Appendix F. Three-Point Inversion

In this appendix we consider three problems where the RMS velocity is given at the interfaces and at an intermediate (inner) point of the interval vs. depth or time, or depth is given vs. time at the interfaces and at an intermediate point, and both parameters of the velocity profile, V_a and k_a , are unknown.

Problem F1. RMS vs. depth with the unknown top interface velocity V_a and top gradient k_a . Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, and the RMS velocity at the inner point of the interval, $V_{2,c}$, the full interval thickness Δz_b (between the top and bottom interfaces) and the partial thickness

Δz_c (between the top interface and the intermediate level). Find the top velocity V_a and the gradient k_a . *Solution.* The resolving equation set follows from Equation (E-3),

$$\begin{aligned} \frac{1}{M_c - M_a} \cdot \ln \frac{M_c}{M_a} + \frac{B_c}{M_c - M_a} \cdot \ln \frac{M_c - 1}{M_a - 1} &= A_c, \\ \frac{1}{M_b - M_a} \cdot \ln \frac{M_b}{M_a} + \frac{B_b}{M_b - M_a} \cdot \ln \frac{M_b - 1}{M_a - 1} &= A_b. \end{aligned} \quad (\text{F-1})$$

Coefficients A_b, A_c, B_b and B_c are known values; they follow from Equation (E-4),

$$\begin{aligned} A_c &= 1 - B_c - \frac{V_{2,c}^2 - V_{2,a}^2}{V_\infty \Delta z_c} \cdot t_a, B_c = V_{2,c}^2 / V_\infty^2, \\ A_b &= 1 - B_b - \frac{V_{2,b}^2 - V_{2,a}^2}{V_\infty \Delta z_b} \cdot t_a, B_b = V_{2,b}^2 / V_\infty^2, \end{aligned} \quad (\text{F-2})$$

where the asymptotic factors M_a, M_b and M_c are to be found,

$$M_a = \frac{V_\infty}{V_\infty - V_a}, M_b = \frac{V_\infty}{V_\infty - V_b}, M_c = \frac{V_\infty}{V_\infty - V_c}. \quad (\text{F-3})$$

It follows from Equation (32),

$$\begin{aligned} M_b - M_a &= M_a^2 \cdot \frac{k_a \Delta z_b}{V_\infty}, \\ M_c - M_a &= M_a^2 \cdot \frac{k_a \Delta z_c}{V_\infty}, \end{aligned} \quad (\text{F-4})$$

Divide the second equation of Equation Set (F-4) over its first equation,

$$\frac{M_c - M_a}{M_b - M_a} = \frac{\Delta z_c}{\Delta z_b}. \quad (\text{F-5})$$

Thus, we solve Equation Set (F-1) together with Equation (F-5), to obtain the three unknown asymptotic factors. To obtain the initial guess, we assume that the layer is thin, *i.e.* that the differences in the asymptotic factors $M_b - M_a$ and $M_c - M_a$ are small. This assumption allows linearization of Equation Set (F-1), which, in turn, results in

$$\begin{aligned} \frac{M_c - M_a}{2M_a \cdot (M_a - 1)} &\approx \frac{B_c M_a - (M_a - 1) \cdot (A_c M_a - 1)}{(M_a - 1)^2 + B_c M_a^2}, \\ \frac{M_b - M_a}{2M_a \cdot (M_a - 1)} &\approx \frac{B_b M_a - (M_a - 1) \cdot (A_b M_a - 1)}{(M_a - 1)^2 + B_b M_a^2}. \end{aligned} \quad (\text{F-6})$$

Introduction of solution (F-6) into Equation (F-5) leads to a fourth order polynomial equation for the top interface asymptotic factor $M_a > 1$,

$$\begin{aligned} \frac{B_c M_a - (M_a - 1) \cdot (A_c M_a - 1)}{B_b M_a - (M_a - 1) \cdot (A_b M_a - 1)} &= \frac{\Delta z_c}{\Delta z_b} \cdot \frac{(M_a - 1)^2 + B_c M_a^2}{(M_a - 1)^2 + B_b M_a^2}. \end{aligned} \quad (\text{F-7})$$

We solve Equation (F-7), choose a proper root, and this yields the initial guess for M_a . Then we use Equation (F-6) to obtain the initial guess for M_b and M_c . Next we solve Equation Set (F-1), (F-5) numerically.

Problem F2. RMS vs. time with the unknown top interface velocity V_a and top gradient k_a . Given data are the RMS velocities at the top and bottom interfaces, $V_{2,a}$ and $V_{2,b}$, and at the inner point of the interval, $V_{2,c}$, the full interval traveltime Δt_b (between the top and bottom interfaces) and the partial traveltime Δt_c (between the top interface and the intermediate level). Find the top velocity V_a and the gradient k_a . *Solution.* The resolving equation set follows from Equation (E-8),

$$\frac{M_c - M_a - \ln \frac{M_c}{M_a}}{M_c - M_a + \ln \frac{M_c - 1}{M_a - 1}} = C_c, \quad (F-8)$$

$$\frac{M_b - M_a - \ln \frac{M_b}{M_a}}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}} = C_b,$$

where C_c and C_b are the normalized local RMS velocities, for the partial interval and the full interval, respectively. These are the known dimensionless values,

$$C_c = \frac{U_c^2}{V_\infty^2} = \frac{W_c}{V_\infty^2 \Delta t_c} = \frac{V_{2,c}^2 \cdot t_c - V_{2,a}^2 \cdot t_a}{V_\infty^2 \Delta t_c}, \quad (F-9)$$

$$C_b = \frac{U_b^2}{V_\infty^2} = \frac{W_b}{V_\infty^2 \Delta t_b} = \frac{V_{2,b}^2 \cdot t_b - V_{2,a}^2 \cdot t_a}{V_\infty^2 \Delta t_b}.$$

We make use of Equation (D-14),

$$\frac{M_c - M_a + \ln \frac{M_c - 1}{M_a - 1}}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}} = \frac{\Delta t_c}{\Delta t_b}, \quad (F-10)$$

and solve Equation Set (F-8) numerically along with Equation (F-10). To obtain the initial guess, we assume that the layer is thin, and thus, we linearize Equation Set (F-8) for the small variations of the asymptotic factor, $M_b - M_a$ and $M_c - M_a$,

$$\frac{M_c - M_a}{M_a} \approx \frac{C_c M_a^2 - (M_a - 1)^2}{M_a - 1}, \quad (F-11)$$

$$\frac{M_b - M_a}{M_a} \approx \frac{C_b M_a^2 - (M_a - 1)^2}{M_a - 1}.$$

Note that for the small increments of the asymptotic factor

$$M_c - M_a + \ln \frac{M_c - 1}{M_a - 1} \approx \frac{M_a \cdot (M_c - M_a)}{M_a - 1}, \quad (F-12)$$

$$M_b - M_a + \ln \frac{M_b - 1}{M_a - 1} \approx \frac{M_a \cdot (M_b - M_a)}{M_a - 1}.$$

Linearization of Equation (F-10) leads to

$$\frac{M_c - M_a}{M_b - M_a} \approx \frac{\Delta t_c}{\Delta t_b}. \quad (F-13)$$

Equation (F-13) is similar to (F-5); however, (F-5) is exact and valid for any layer thickness, while (F-13) is an approximation for a thin layer only. Introduction of (F-13) into (F-11) results in a quadratic equation,

$$\frac{C_c M_a^2 - (M_a - 1)^2}{C_b M_a^2 - (M_a - 1)^2} = \frac{\Delta t_c}{\Delta t_b}. \quad (F-14)$$

The equation has two positive roots, but only one of them exceeds 1. The solution is

$$M_a = \frac{\sqrt{\Delta t_b - \Delta t_c}}{\sqrt{\Delta t_b - \Delta t_c} - \sqrt{C_c \cdot \Delta t_b - C_b \cdot \Delta t_c}}. \quad (F-15)$$

Next we apply Equation (F-11) to obtain the initial guess for M_b and M_c . Then we solve Equation Set (F-8), (F-10) numerically and deliver the three asymptotic factors.

Problem F3. Depth vs. time with the unknown top interface velocity V_a and top gradient k_a . Given data are the traveltime and layer thickness for the full and the partial intervals, respectively: $\Delta t_b, \Delta z_b$ and $\Delta t_c, \Delta z_c$. Find the top velocity V_a and the gradient k_a . *Solution.* The resolving equation set follows from Equation (E-14),

$$\frac{M_c - M_a}{M_c - M_a + \ln \frac{M_c - 1}{M_a - 1}} = D_c, \quad (F-16)$$

$$\frac{M_b - M_a}{M_b - M_a + \ln \frac{M_b - 1}{M_a - 1}} = D_b,$$

where D_b and D_c are the normalized interval velocities for the full and the partial intervals,

$$D_c = \frac{V_{\text{Int},c}}{V_\infty} = \frac{\Delta z_c}{V_\infty \Delta t_c}, \quad D_b = \frac{V_{\text{Int},b}}{V_\infty} = \frac{\Delta z_b}{V_\infty \Delta t_b}. \quad (F-17)$$

Equation Set (F-16) should be solved numerically along with Equation (F-5). To obtain the initial guess, we linearize Equation Set (F-16) for the small increments of the asymptotic factors,

$$\frac{M_c - M_a}{2M_a} \approx D_c M_a - (M_a - 1), \quad (F-18)$$

$$\frac{M_b - M_a}{2M_a} \approx D_b M_a - (M_a - 1).$$

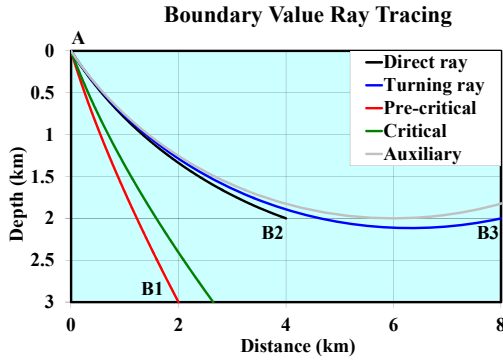


Figure 9. Boundary value ray tracing with the Hyperbolic velocity model: AB_1 —pre-critical ray, AB_2 —post-critical ray without the turning point, AB_3 —post-critical turning ray, green line—critical ray, grey line—auxiliary post-critical ray with the turning point at the given depth.

Introduction of Equation Set (F-18) into Equation (F-5) results in

$$\frac{D_c M_a - (M_a - 1)}{D_b M_a - (M_a - 1)} = \frac{\Delta z_c}{\Delta z_b}. \quad (F-19)$$

Equation (F-19) yields the initial guess for the top asymptotic factor M_a ,

$$M_a = \frac{\Delta z_b - \Delta z_c}{\Delta z_b - \Delta z_c - (D_b \Delta z_c - D_c \Delta z_b)}. \quad (F-20)$$

Equation (F-18) yields the initial guess for the bottom and the inner asymptotic factors, M_b and M_c . Then we solve Equation Set (F-16) numerically, along with Equation (F-5).

Appendix G. Numerical Examples of BVRT

In this appendix, we present three examples of the boundary value ray tracing, **Figure 9**. The parameters of the velocity profile are the same as above. For each case, the source point A is located at the origin of the frame.

The destination point is different for each case: B_1 ($x = 2$ km, $z = 3$ km), B_2 ($x = 4$ km, $z = 2$ km), and B_3 ($x = 8$ km, $z = 2$ km). As we will show, the first case corresponds to the pre-critical ray path, the second corresponds to the post-critical path that does not include the turning point, and the last case leads to the post-critical path with the turning point. First we calculate the critical

lateral propagation with Equation (101): $\Delta x_c = 1.587$ km for the destination point depth $z_b = 2$ km, and $\Delta x_c = 2.646$ km for $z_b = 3$ km. Next we compare the horizontal offset $\Delta x = x_B - x_A$ to the critical lateral propagation and apply the criterion in Equation (102): $\Delta x_1 < \Delta x_c$ for the destination point B_1 and $\Delta x_{2,3} > \Delta x_c$ for the destination points B_2 and B_3 . We conclude that ray AB_1 is pre-critical, while rays AB_2 and AB_3 are post-critical.

We solve Equation (105) for pre-critical ray AB_1 and establish its eccentricity: $m = 1.18647$. The ray angles at the departure and the destination points are: $\alpha_a = 0.43501 \approx 24.92^\circ$ and $\alpha_b = 0.68430 \approx 39.21^\circ$.

Next we figure out whether the post-critical rays AB_2 and AB_3 include the turning point or not. If the post-critical ray includes a turning point, then the ray angle at the destination is obtuse, and this should be taken into account in Equation (104). For this we establish an auxiliary post-critical ray, whose maximum penetration depth is equal to the destination depth of points B_1 and B_2 : $z_b = 2$ km. We solve Equation (68) for the eccentricity,

$$m = \frac{\hat{z}}{1 + \hat{z}}, \text{ where } \hat{z} = Q(z + h). \quad (G-1)$$

Compare Equation (G-1) with Equation (22). It is interesting to note that the eccentricity of a post-critical ray with the turning point at the given depth is equal to the normalized velocity at this depth. We obtain the eccentricity of the auxiliary ray $m = 0.7$, and the take-off angle $\alpha_a = 0.79560 \approx 45.58^\circ$. Calculate the half-chord of this ray with Equation (98), $\Delta x_{aux}/2 = 5.994$ km. Thus, we see that for the destination point B_2 , the offset is less than the half-chord of the auxiliary ray, $\Delta x_2 < \Delta x_{aux}/2$, while for the destination point B_3 , $\Delta x_3 > \Delta x_{aux}/2$. Hence we conclude that ray AB_2 does not include the turning point, while ray AB_3 includes the turning point. We solve Equation (105) for ray AB_2 and obtain the eccentricity, $m = 0.71798$, and the ray angles at the endpoints, $\alpha_a = 0.77036 \approx 44.14^\circ$, $\alpha_b = 1.34651 \approx 77.15^\circ$. For ray AB_3 the results are: $m = 0.70681$, $\alpha_a = 0.78582 \approx 45.02^\circ$, and $\alpha_b = 1.70973 \approx 97.96^\circ$. Note that the last angle is obtuse. The three traced rays, along with the auxiliary ray (grey line) and the critical ray (green line) are plotted in **Figure 9**.