

Study on Quantitative Model of Karst Drainage Basin Water-Holding Based on Principal Component Analysis: A Case Study of Guizhou, China^{*}

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Received June 19, 2013; revised July 19, 2013; accepted August 8, 2013

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ABSTRACT

In Karst drainage basins, there are the ground water and underground water exchanging frequently, and the shortage of water resources due to having the special double aquifer mediums and unique surface and subsurface river systematic structure. This paper is to select 20 research sampling areas coming from Guizhou Province, and according to the spectral characteristics of the catchment water-holding mediums and vegetations, and using the remote sensing technique, extract the watershed vegetation index. According to the principle of principal component analysis, using the software of Spss and Matlab is to analyze the impacts of watershed vegetation type on the catchment water-holding ability, and establish the principal component analysis function. Studies have shown that: 1) the watershed vegetation coverage rate plays an important role in Karst basin water-holding ability; 2) the catchment water-holding ability is the comprehensive reflection and manifestation of the Catchment Water-storing Capacity (*CWC*); 3) it is much better effects and higher accuracy to monitor/forecast the catchment water-holding volume by using the vegetation indices.

Keywords: Karst Drainage Basin; Watershed Vegetation Index; Catchment Water-Holding Ability; Catchment Water-Storing Capacity; Quantitative Model

1. Introduction

Karst is a kind of fragile eco-environment composed mainly of Water Resources [1]. In Karst areas, there are the surfaces broken, the slopes steep and mountains high, and the valleys deeply cut that will cause the surface water infiltrated seriously, the subsurface water deeply hidden, and the shortage of the catchment water resources. The broken Karst surfaces, shorter soil-formed times, thinner soil-layer and lower soil-fertility result in the difficulty of vegetation growth and the water/soil out-flow severely. The development of the fractures and conduits within the rocky layer below surfaces causes the surface water flowing rapidly and the difficulty of

catchment water-holding. Hence these will affect on the Catchment Water-storing Capacity (*CWC*). The catchment water-holding is the comprehensive reflection of the strong/weak of the *CWC*, and the manifestation of the spatial distribution of water resources. The Vegetation Index (*VI*) is the important information of the catchment water-holding, and through calculating and analyzing the watershed vegetation index is to reveal the catchment water-holding rules, and reflect the spatial distribution characteristic of water resources. After the developments of nearly 20 years, the Vegetation Index (*VI*) has dozens of kinds. It is the commonly used forms for these Vegetation Indices (*VI*s) like the Vegetation Index (*RVI*), Difference Vegetation Index (*DVI*), Normalized Difference Vegetation Index (*NDVI*), Transformational Vegetation Index (*TVI*), Return Difference Vegetation Index (*RDVI*) and Enhanced Vegetation Index (*EVI*). The *VI*s have widely applied to the studies to the global and regional land cover, vegetation classification and environ-

^{*}This study was jointed funded by Natural Science Funds from Department of Water Resources of Guizhou Province, China ((KT 201105), (KT 201010), (KT200802)), and Guizhou Science and Technology Department (QKHJ (2010)2026), (QKHJ (2013)2208), and Department of Education of Guizhou Province (QJK 2009 (0039) and QJK-2006307).

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mental variation [2-14]. At present, as for the studies about the Karst catchment water-holding, He zhonghua *et al.*, [15-17] had some relevant discussions, but, the less *VI*s-based studying Karst catchment water-holding. This paper is to select the 20 sampling sites having the continuously 5-year observation hydrological and remote sensing data of Guizhou Province. Using the remote sensing technique extracts the watershed vegetation indices of the *RVI*, *DVI*, *NDVI*, *TVI*, *RDVI* and *EVI* from TM images. Utilizing mathematical methods explores the relationship between catchment water-holding and *VI*s, and establishes the quantitative remote sensing model of Karst drainage basin water-holding. At last, utilizing 5 sampling areas is to be made the significant test for the model, which has a good predictive effect. This paper is useful for the no-data regions to calculate the water resources, and provides a stronger theoretical basis for us to more fully and reasonably estimate and utilize karst water resources.

2. Sample Area Selection & Data Acquisition

2.1. Selections of the Sample Area

According to the study purpose, this paper is to select 20 sampling sites located in the central part, southern, southwest area, western and northern of Guizhou Province, respectively. They are all distributed in the typical Karst areas, with the condition of Geology and geomorphology of them similar as possible, which reflects the spatial variation characteristics and rules of catchment water-holding. Meanwhile, the 20 sample sites been in the same climatic zones is to guarantee the precipitation or watershed prophase-water content the same as possible.

2.2. Hydrologic Data

Hydrologic data comes from the *Guizhou Statistics on Mean Monthly Flows per Calendar Year* compiled by Guizhou Hydrologic Station and *Water Resources Bulletin of Guizhou Province* compiled by the Department of Guizhou Hydrology & Water Resources. This paper is to select 20 hydrologic section data that are in the same climatic zones during the period from Sep. 2005 to Sep. 2010. Drainage area is generally dominated by small watershed, which is to ensure that the geological conditions of basin underlying surfaces can be as far as possible the same or similar. To calculate the mean monthly runoff depth of 20 research sampling sites and make its standardized processing using the Formula (7) is shown on **Table 4**.

2.3. Remote Sensing Data

2.3.1. Preprocesses of Remote Sensing Images

(1) Selection of remote sensing data

In order to guarantee less than 30% of the cloud-cover

amount per research time, this paper is to select the TM images of a total of 6 periods in 6 years during the period from Sep. 2005 to Sep. 2010.

(2) Atmospheric correction

Currently, there are many methods of atmospheric correction, and the atmospheric radiative transfer model is more accurate method. It is utilizing the radioactive transfer principles of electromagnetic wave in the atmosphere to be established the atmospheric correction model of remote sensing image. This paper is to adopt the FLAASH model, namely improved MORTRAN model, which can be made atmospheric correction not only for hyperspectral data, but also for multispectral data like LANDSAT, SPOT, AVHRR, MERIS, IRS and ASTER, etc.

2.3.2. Apparent Reflectance Calculation

(1) Calculation of spectral radiance

$$L = G_{ain} \cdot DN + B_{ias} \quad (1)$$

If there are no calibration parameter data of G_{ain} and B_{ias} , a band L can be calculated by the Formula (2)

$$L = \frac{L_{max} - L_{min}}{QCAL_{max} - QCAL_{min}} \cdot (QCAL - QCAL_{min}) + L_{min} \quad (2)$$

where $QCAL$ is the DN value of a pixel, namely $QCAL=DN$; $QCAL_{max}$ is the maximum value 255, and $QCAL_{min}$ is the minimum value. The Formula (2) can be changed to the Formula (3) for LandSat-7 (namely $QCAL_{min} = 1$) H, and to the Formula (4) for LandSat-5 (namely $QCAL_{min} = 0$).

$$L = \frac{L_{max} - L_{min}}{254} \cdot (DN - 1) + L_{min} \quad (3)$$

$$L = \frac{L_{max} - L_{min}}{255} \cdot DN + L_{min} \quad (4)$$

(2) Calculation of apparent reflectivity [18-20].

$$\rho = \frac{\pi \cdot L \cdot D^2}{ESUN \cdot \cos \theta} \quad (5)$$

where ρ is the apparent reflectivity of the Top of the Atmosphere (TOA) (Dimensionless); π is a constant (steradian sr); D is the distance of Sun-Earth, calculated the Sun-Earth distance of any day of the year according to the **Table 1**; $ESUN$ is the mean Solar spectral irradiance of the TOA , can be lookup from the **Table 2** [21,22]; θ is the Solar Zenith Angle (Viz., $\theta = 90^\circ - \beta$), and $\cos \theta$ can directly calculated by the Formula (6) [23].

$$\cos \theta = \sin \Phi \sin \delta + \cos \Phi \cos h \quad (6)$$

where Φ is the Geographic Latitude, δ is the Solar Declination, and h is the Sun Angle

2.3.3. Selection & Calculation of Vegetation Index

From the theoretical analysis, the *DN* value of the original remote sensing image is without any correction, including the radiation calibration, is only a digital conversion form of radiant energy got into the sensor, and can't essentially reflect the radiation characteristics of target objects. The *L* and ρ have been made radiometric calibration correction. The ρ is the reflectivity of the surface features after atmospheric correction, and can essentially reflect the radiation characteristics of target objects. Therefore, the *VI* established by the ρ can reflect the vegetation coverage rate and its changing of watershed underlying surfaces. The common vegetation indices are listed in the following **Table 3** [24-30].

The Vegetation Index (*VI*) is a quantitative indicator of catchment water-holding, indicating that the larger (or smaller) the *VI*, the more (or less) the biomass and the more developed (or less developed) the plant roots, which results in the more (or less) the interception amount of vegetation to precipitation, the higher (or lower) the amount of rainfall infiltration, and shows the stronger

Table 1. Sun-Earth distance at different lime (Astronomical units).

Day	Distance								
1	0.9832	74	0.9945	152	1.0140	227	1.0128	305	0.9925
15	0.9836	91	0.9993	166	1.0158	242	1.0092	319	0.9892
32	0.9853	106	1.0033	182	1.0167	258	1.0057	335	0.9860
46	0.9878	121	1.0076	196	1.0165	274	1.0011	349	0.9843
60	0.9909	135	1.0109	213	1.0149	288	0.9972	365	0.983

Table 2. Mean solar spectral irradiance at the atmospheric top for Landsat-7 ($w \cdot m^{-2} \cdot \mu m^{-1}$).

Band	1	2	3	4	5	7
Landsat-7 ESUN	1969.00	1840.00	1551.00	1044.00	225.70	82.07

(or weaker) the Catchment Water-storing Capacity (*CWC*) and the more significant (or less significant) the catchment water-holding. Therefore, the different vegetation indices are to reflect the catchment water-holding status from a different point of view. This paper is to select the remote sensing data *LandSat-7*, firstly, using the Formula (3) is to compute the spectral radiance of surface objects; secondly, using Formulas (5) and (6) and based on **Tables 1** and **2** calculates the apparent reflectivity of them; thirdly, using the formulas of the **Table 3** are to calculate the *VI*s of the apparent reflectivity like *RVI*, *DVI*, *NDVI*, *TVI*, *RDVI* and *EVI* and make standardized processing for them by use of the Formula (4), shown on **Table 4**.

3. Quantitative Analysis of Catchment Water-Holding

3.1. Principles of Quantitative Analysis

In Karst drainage basins, there are many factors affecting on catchment water-holding ability, and which are complex and changeable relationships. In this study, taking vegetation factors as an example is to explore the impacts of vegetation factors under different level (or factor) on catchment water-holding ability. Utilizing the principle of principal component analysis is to compute the contribution rate of each factor of vegetations to catchment water-holding [31,32].

1) The original data matrix *X*

$$X = \begin{bmatrix} x_{11} & x_{12} & \dots & x_{1p} \\ x_{21} & x_{22} & \dots & x_{2p} \\ \vdots & \vdots & & \vdots \\ x_{n1} & x_{n2} & \dots & x_{np} \end{bmatrix}$$

And make standardized processing, Viz.,

$$x_{aj}^* = \frac{x_{aj} - \bar{x}_j}{\sigma_j} \tag{7}$$

where *p* is the number of variables; *n* is the number of

Table 3. The *VI* formulas in this paper.

Vegetation Index	Formula	References
<i>RVI</i>	$RVI = \rho_{nir} / \rho_{red}$	R. L. Pearson & D. L. Miller (1972)
<i>DVI</i>	$DVI = \rho_{nir} - \rho_{red}$	Jordan (1969)
<i>NDVI</i>	$NDVI = (\rho_{nir} - \rho_{red}) / (\rho_{nir} + \rho_{red})$	Rouse et al. (1974)
<i>TVI</i>	$TVI = \sqrt{NDVI + 0.5}$	Huete (1988)
<i>RDVI</i>	$RDVI = \sqrt{DVI * NDVI}$	Reujean & Breon (1995)
<i>EVI</i>	$EVI = 2.5 \times [(\rho_{nir} - \rho_{red}) / (\rho_{nir} + 6.0 \times \rho_{red} - 7.5 \times \rho_{blue} + 1)]$	

Note: the ρ_{blue} , ρ_{red} and ρ_{nir} is the blue band, red band and near-infrared band of TM, respectively.

Table 4. The hydrological data and VIs of Karst basin sample areas.

Serial Number	County Name	Hydrological station name	Runoff Depth	Vegetation Indices (VIs)						
				RVI	DVI	NDVI	TVI	RDVI	EVI	Z ₁
1	Guiding	Xiawan	-0.66686	-0.24342	-0.01505	0.04061	-0.02034	-0.05701	0.11732	-0.01697
2	Jianhe	Nanshao	-0.32800	0.12790	0.24738	0.19943	0.11850	0.04866	0.19168	0.160535
3	Tongren	Tongren	1.11470	0.60338	0.56929	0.53327	0.52682	0.73610	0.59310	0.599093
4	Xiuwen	Xiuwendianchang	0.25244	0.20499	0.29112	0.28091	0.27275	0.17761	0.36667	0.274688
5	Yuqing	Yuqing	-1.13658	-0.89842	-0.39819	-0.44166	-0.31584	-0.72364	-0.74032	-0.57372
6	Xishui	Shisun	-0.74738	-0.26973	-0.06928	-0.00637	-0.08198	-0.39852	-0.02133	-0.13113
7	Guiyang	Guiyang	-1.05605	-0.36946	-0.27047	-0.30654	-0.23273	-0.64937	-0.29417	-0.35263
8	Dushan	Xiasi	-0.50582	0.09221	0.21239	0.09133	0.02292	0.00297	0.18497	0.103495
9	Jiangkou	Jiangkou	-0.70377	-0.26423	-0.06578	0.01192	-0.03485	-0.32196	0.08321	-0.08601
10	Pu'an	Caopingtou	0.11488	0.20003	0.27362	0.22229	0.18480	0.11164	0.21963	0.206102
11	Bijie	Xuhuatun	-1.32111	-2.75310	-4.09841	-4.24200	-4.38618	-3.44043	-4.16139	-3.97863
12	Anlong	Pojiao	1.73875	1.54560	0.76524	0.83261	0.80207	1.23974	0.97832	1.008591
13	Daozhen	wujiayuanzi	-0.27431	0.12995	0.26663	0.21315	0.13253	0.05536	0.19168	0.170397
14	zhenning	Gaoche	0.46381	0.34936	0.45907	0.43806	0.45242	0.52318	0.40916	0.44879
15	Shibing	Shidong	1.21535	0.67195	0.68651	0.58108	0.55647	0.89698	0.61267	0.672225
16	Puding	sanchahe	1.55757	1.33536	0.73725	0.77523	0.74036	1.19828	0.66579	0.892018
17	Pingba	Xujiadu	-0.94198	-0.33798	-0.21624	-0.12652	-0.12729	-0.63490	-0.24106	-0.27441
18	Zheng'an	Zheng'an	0.27592	0.26537	0.39959	0.29297	0.29332	0.34343	0.37170	0.335193
19	Majiang	Xiasier	0.43697	0.32650	0.41359	0.37487	0.34172	0.39935	0.39742	0.383134
20	Jinsha	Mukong	1.29252	0.97093	0.73550	0.75736	0.72222	1.15012	0.65404	0.830153

**Note: WATER is the runoff depth of the surface water; Z₁ is the first principal component of the VIs.

samples; \bar{x}_j , σ^2 is the sample mean and total variance, respectively. Its expression is as follows:

$$\bar{x}_j = \frac{1}{N} \sum_a x_{aj}$$

$$\sigma^2 = \frac{1}{N} \sum_a (x_{aj} - \bar{x}_j)^2$$

2) Calculate the correlation coefficient matrix R

$$r_{ij} = \frac{\frac{1}{N} \sum_a (x_{ai} - \bar{x}_i)(x_{aj} - \bar{x}_j)}{\sigma_i \sigma_j} = \frac{1}{N} \sum_a x_{ai}^* x_{aj}^* \quad (8)$$

where r is the correlation coefficient; N is the total number of samples.

3) Calculation of eigenvalues and eigenvectors

According to the characteristic equation $|R - \lambda I| = 0$ is to calculate the eigenvalue, viz., solving the characteristic polynomial:

$$r_n \lambda^p + r_{n-1} \lambda^{p-1} + \dots + r_1 \lambda + r_0 = 0 \quad (9)$$

To compute the value of $\lambda_1, \lambda_2, \dots, \lambda_p$, and sorted λ_i by size as the following:

$$\lambda_1 \geq \lambda_2 \geq \dots \geq \lambda_p \geq 0$$

The eigenvector (l_k) of the eigenvalue (λ_k) is listed by

$$l_k = [l_{k1}, l_{k2}, \dots, l_{kp}]^T$$

$$Rl_k = \lambda l_k \quad (10)$$

4) To calculate the contribution ratio $\lambda_k / \sum_{i=1}^p \lambda_i$ and the cumulative contribution ratio

$$\sum_{j=1}^k \left(\lambda_j / \sum_{i=1}^p \lambda_i \right) \quad (11)$$

To take generally the principal component of the eigenvalue $\lambda_1, \lambda_2, \dots, \lambda_m$ ($m \leq k \leq p$) of greater than 80% of cumulative contribution ratio.

expresses the score of six factors like *RVI*, *DVI*, ..., *EVI*. The Z_1 is represented the *FPC*, then, the **Table 7** may be written as the score function

$$Z_1 = 0.123RVI + 0.174DVI + \dots + 0.182EVI \quad (14)$$

4. Quantitative Model Establishment of Catchment Water-Holding

4.1. Principle of Quantitative Model

Suppose that the relationship between hydrologic sectional observation value y and vegetation indices x in Karst drainage basins can be expressed by the following model [31,32].

$$y = b_0 + b_1x_1 + b_2x_2^2 + \dots + b_nx_n^n + \varepsilon \quad (15)$$

where b_0 is the constant, and b_1, b_2, b_n is the coefficient of independent variables x_1, x_2, x_n , respectively; ε is the random variable obeyed normal distribution with the mean 0 and variance σ^2 . In order to evaluate the precision of regression equation that needs to be made significant test with the statistic F .

$$R^2 = \frac{\sum(\hat{y}_k - \bar{y})^2}{\sum(y_k - \bar{y})^2} \quad (16)$$

$$F = \frac{\frac{S_{\text{Regression value}}}{m}}{\frac{S_{\text{Residual value}}}{n-m-1}} \sim F(m, n-m-1) \quad (17)$$

where $S_{\text{Regression value}}$ is the regression value, and is equal to $\sum_{k=1}^n (\hat{y}_k - \bar{y})^2$; $S_{\text{Residual value}}$ is the residual value, and is equal to $\sum_{k=1}^n (y_k - \hat{y}_k)^2$.

The statistic F is obeyed the F -distribution with the first freedom degree m and second freedom degree $n-m-1$. The critical value $F_{\alpha}(m, n-m-1)$ can be lookup from F -distribution table with a given α (such as $\alpha = 0.05$). If the statistic value F is greater than critical value $F_{\alpha}(m, n-m-1)$, indicating that the regression model established with these data is significant and can be used for the analysis of regional water resources, In contrast, can not be used.

4.2. Establishment of Quantitative Model

Firstly, on basis of **Table 4**, using the Formula (14) is to calculate the *FPC* (z_1) of the *VI*s of six factors, and shown on the **Table 4**. Secondly, on basis of the **Table 4** again, using the Formula (15) and the Spss and Matlab Software is to compute the quantitative model coefficient between catchment water-holding and *FPC* (Z_1) of the *VI*s, and shown on the **Table 8**. Thirdly, the model-fitted

degree between catchment water-holding and *FPC* (Z_1) is computed utilizing the Formula (16) and made significant test with the F -distribution by the Formula (17) (**Table 8**).

(1) We know from the **Table 8** that the catchment water-holding (*WATER*) of Karst drainage basins is fitted by the *FPC* (Z_1) of the *VI*s, with very good its fitting effect and very high the model-fitted coefficient ($R = 0.974$). The dynamic change model of catchment water-holding is established by using the Formula (15), with very high the multiple correlation coefficient ($R^2 = 0.974$) and very small the root mean square error ($RMSE = 0.1513$). The quantitative model established by using the Formula (17) is to be made the F -test. The statistic value ($F = 313.732$) is greater than the critical value ($F_{0.01} = 6.11$), indicating that the built quantitative model is significant, and illustrating that it is very good to monitor dynamically the catchment water-holding with the first principal component of vegetation index.

(2) On basis of the **Table 8** and using the Formula (15), the dynamic change monitoring, forecasting quantitative model of Karst drainage basin water-holding can be expressed as the following:

$$WATER = -0.451 + 1.886Z_1 + 0.419Z_1^2 \quad (18)$$

To sum up, in Karst areas, there are the hydrology dynamic changes violently, surface water infiltrated severely and the catchment water-holding ability badly due mainly to the rugged surface, caves and conduits crisscrossed subsurface; the thinner soil-layer and lower soil-fertility result in the difficulty of vegetation growth, and the water/soil outflow severely; which is to be formed a special, fragile karst environment, and restricts severely the Catchment Water-storing Capacity (*CWC*). The catchment water-holding is the manifestation of the *CWC*, while the *VI* is the important information of the catchment water-holding and an important indicator of the *CWC*. A large/small of the *VI*s is to affect on the velocity and residence time of rainfall in the catchment surface and the rainfall infiltration amount, which further influences the large/small of the *CWC*. Therefore, the vegetation index is the comprehensive reflection and manifestation of the catchment water-storing and water-holding.

4.3. Evaluation of the Model Accuracy

In order to assess the accuracy of the monitoring, forecasting quantitative model, the 5 research sampling areas selected randomly in Karst drainage basin, made some processing using the above methods, are to be extracted the *VI*s, respectively, e.g., *RVI*, *DVI*, *NDVI*, *TVI*, *RDVI* and *EVI*. Utilizing the Formula (14) is to calculate the *FPC* (z_1) value (**Table 9**). The **Table 9** is computed by using the Formula (18) and Eviews software, and the

Table 8. The coefficient table of models.

Variables Coefficient	Constant	Z_1	Z_1^2	<i>RMSE</i>	<i>R</i>	<i>R</i> ²	<i>F</i>	<i>F</i> _{α}	<i>Sig.</i>
	<i>b</i> ₀	<i>b</i> ₁	<i>b</i> ₂						
1	-0.451	1.886	0.419	0.1513	0.987	0.974	313.732	6.11	**

Note: * is a significant as $\alpha = 0.05$; ** is a particularly significant as $\alpha = 0.01$.

Table 9. The model-test table.

Serial Number	County Name	Hydrological station name	Vegetation Indices (<i>VIs</i>)							Runoff depth		Absolute error	Relative error absolute %
			<i>RVI</i>	<i>DVI</i>	<i>NDVI</i>	<i>TVI</i>	<i>RDVI</i>	<i>EVI</i>	Z_1	Measured value	Predicted value		
1	Yanhe	Tangba	-2.18091	-1.63160	-1.22409	-0.89234	-1.32676	-1.35420	-1.4089	-1.9768	-2.2753	0.2985	15.10
2	Zhenfeng	Datianhe	0.46246	0.46782	0.44970	0.46815	0.58299	0.50644	0.4968	0.7658	0.5898	0.1760	22.98
3	Cengong	Chebian	-0.33798	-0.16026	-0.11779	-0.08295	-0.47244	-0.12141	-0.2064	-0.9420	-0.8218	-0.1202	12.76
4	Luodian	Shimenkan	-1.16721	-0.34046	-0.44166	-0.26261	-0.67654	-0.29417	-0.4978	-1.0561	-1.2855	0.2294	21.72
5	Qingzhen	Yachihe	1.53644	0.74074	0.81182	0.80207	1.23515	0.68424	0.9451	1.6280	1.7063	-0.0783	4.81

results are compared with the runoff depth measured value, which concluded the 22.98% of the maximum relative error value, 4.81% of the minimum relative error value. It indicates that the *FPC* of the *VIs* can be used to dynamically monitor, forecast the catchment water-holding, with very good the effect and very high the accuracy.

5. Conclusions and Analysis

(1) The correlation coefficient between the *VIs* and the runoff depth is greater than 0.5, indicating that the vegetation coverage rate in Karst drainage basins plays an important role in the catchment water-holding. The descending order of the impacts of the different vegetation indices on the catchment water-holding is the *RVI* (0.847) > *RDVI* (0.82) > *EVI* (0.629) > *DVI* (0.621) > *NDVI* (0.613) > *TVI* (0.601).

(2) The first principal component function model based on the *VIs* can be expressed as the following:

$$Z_1 = 0.123RVI + 0.174DVI + \dots + 0.182EVI$$

(3) The dynamic monitoring, forecasting quantitative model of catchment water-holding based on the *FPC* can also be expressed as the following:

$$\text{WATER} = -0.451 + 1.886Z_1 + 0.419Z_1^2$$

This model through the variance analyzing and the sample areas testing shows very good monitoring, prediction effect.

In short, in Karst areas, there are the mountains high

and slopes steep, carbonate rocks distributed widely, and the surface and subsurface connected, interlinked everywhere, which causes the surface and subsurface water exchanging frequently, the catchment water-holding difficultly. The watershed vegetation index is the important information of the catchment water-holding situation and its spatial distribution. The catchment water-holding model based on the *FPC* of the *VIs*, on the one hand, reflects the catchment water-holding situation from the different angles of the *VIs*; on the other hand, extracting the principal component will reach to reduce the number of variables and the relevance between the variables, make that the model can comprehensively reflect the catchment water-holding situation and its spatial distribution characteristics, and the model prediction accuracy is higher. Establishing the Karst drainage basin water-holding model based on the *FPC* of the *VIs*, is to be applied to solving the calculations of the water resources volume in no-data areas, and provides a strong theoretical basis for us to more fully estimate, more rationally utilize the Karst water resources.

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