

# GOSAT CH<sub>4</sub> and CO<sub>2</sub>, MODIS Evapotranspiration on the Northern Hemisphere June and July 2009, 2010 and 2011

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## ABSTRACT

The Greenhouse gases Observing Satellite (GOSAT) affords an ability to assess and monitor CH<sub>4</sub> and CO<sub>2</sub> near-surface atmospheric concentrations globally on monthly scales pertaining to biogeochemical cycles and anthropogenic emissions. In addition to GOSAT our investigation incorporates global-monthly estimates of evapotranspiration (ET) from the Moderate Resolution Spectroradiometer (MODIS) and fire/wildfire locations for correspondence and comparison. We restrict the investigation to the months of June and July in years 2009, 2010 and 2011. After processing and assessment on the northern hemisphere we focus on two regions in Eurasia for interrogation: 40° to 80°E by 50° to 58°N and 100° to 140°E by 50° to 58°N. The regions allow for contrasting regional settings, an agricultural-industrial-urban west-region to a boreal-steppe discontinuous permafrost zone tundra and thaw lake east-region. Joint probability density functions allow us to identify significant modes, the highest probable values of background levels of CH<sub>4</sub> and CO<sub>2</sub> to ET and develop regressions for correlated relationships. We found that background levels of CH<sub>4</sub>, CO<sub>2</sub> and ET were not affected by the wildfires of 2010. Regressions indicate significant inverse relationships of CH<sub>4</sub> and CO<sub>2</sub> to ET in the west-region and no significant relationships in the east-region. The east-region shows significantly higher background levels of CH<sub>4</sub>, CO<sub>2</sub> and ET owing to the heterogeneity of ecosystems, hydrology, physical processes and terrain in the discontinuous permafrost zone of the central Siberian Plateau.

**Keywords:** GOSAT; CH<sub>4</sub>; CO<sub>2</sub>; MODIS; Evapotranspiration; Wildfire

## 1. Introduction

Addressing the uncertainties in near-surface atmospheric CH<sub>4</sub> and CO<sub>2</sub> sources and sinks is a task that space-based remote sensing has engaged in recent years to complement and extend ground-based sensing methods. GOSAT a joint project by the Japan National Institute of Environmental Studies, the Japan Ministry of Environment and the Japan Aerospace Exploration Agency (JAXA) employs dual-channel Fourier transform spectrometry to measure dry-air concentrations of CH<sub>4</sub> and CO<sub>2</sub> globally since launch in early 2009 [1].

CH<sub>4</sub> and CO<sub>2</sub> are constituents of the Earth's atmosphere. Their concentrations derive from physical processes of long-lived biogeochemical cycles on lands and in oceans. Additionally variations in their concentration owe to anthropogenic activities since wide spread industrialization in the 19th century. Their affects on the Earth's near-surface energy budget, *i.e.* greenhouse gases warming, have been speculated, postulated and debated extensively [2-5].

Our purpose is to investigate correspondences of background levels of CH<sub>4</sub>, CO<sub>2</sub> and ET through joint probability density functions and develop regressions for correlated relationships. We compose and co-geolocate the datasets on the northern hemisphere, and then focus on two regions of interest in Eurasia during June and July 2009, 2010 and 2011. We test a hypothesis that the wildfires in the summer of 2010 affected background levels of CH<sub>4</sub>, CO<sub>2</sub> and ET on a regional and hemispheric basis.

## 2. Data and Methods

### 2.1. GOSAT

Launched and operating since January 2009 the Greenhouse gases Observing Satellite (GOSAT) has been providing global estimates of the column average dry air mole fractions of CH<sub>4</sub> and CO<sub>2</sub> [1]. The GOSAT sensing instrument is the Thermal and Near Infrared Sensor for Carbon Observations (TANSO) Fourier Transform Spectrometer (FTS). TANSO-FTS exploits absorption band characteristics of backscatter sunlight from surface and

atmosphere in two channels: a spectral interferometer. The spectral resolution of TANSO-FTS is 0.3 cm. The nadir footprint radius is about 5 km. A second instrument the TANSO-Cloud and Aerosol Imager (CAI) observes surface and atmosphere reflectance in the shortwave and near infrared to provide cloud screening flags at several hundred points within a FTS footprint. Vertical profiles of pressure, temperature and humidity together with surface pressure and wind speed are taken from European Center for Medium-Range Weather Forecasts. Surface elevation data (within the FTS footprint) are taken from the GTOPO30 database.

The retrieval algorithm uses a radiative transfer inverse method based on Phillips-Tikhonov regularization and the L-curve method to estimate retrieval parameters to simultaneously derive near surface and atmosphere column average concentrations of CH<sub>4</sub> and CO<sub>2</sub> [1]. The algorithm uses radiances in 4 spectral windows cover the O<sub>2</sub>A-band, weakly absorbing CO<sub>2</sub> band, CH<sub>4</sub> band and strongly absorbing CO<sub>2</sub> band. Retrieval parameters are in 12-layer vertical profiles of CH<sub>4</sub> and CO<sub>2</sub> column number and total number density of H<sub>2</sub>O (interfering absorber) and scattering-spectral shift parameters.

During April 2009 through July 2010 validation campaigns were performed using six Total Carbon Column Observing Network (TCCON) sites that use ground-based FTS: two in Australia, Europe and North America. Bias in CH<sub>4</sub> is  $-0.3\% \pm 0.26\%$  and in CO<sub>2</sub> is  $-0.05\% \pm 0.37\%$ , on average.

GOSAT TANSO-FTS by design has a critical dependence on solar angle and cloud (contamination by meteorological cloud water vapor) that reduces the number of reliable CH<sub>4</sub> and CO<sub>2</sub> retrievals in the high latitudes. Therefore we limit our time period of interest to June and July 2009, 2010 and 2011 to obtain the maximum number of retrievals. From the files we extract geolocation information (latitude-longitude relative to the WGS-84 reference ellipsoid), concentration of CH<sub>4</sub> and CO<sub>2</sub> in ppb and ppm, respectively, and the standard error values.

June and July retrievals across the northern hemisphere 40°N latitude and higher are irregular grids which become sparse at 70°N. We apply a least squares Green's Function interpolation such that the original concentration values per geolocation are constraints to produce regular equal-area projection grids at 5 km intervals [6]. The resultant CH<sub>4</sub> and CO<sub>2</sub> mixing ratios (concentrations) and standard error grids are shown in **Figure 1**. Gray-areas indicate no data. The oceans are masked; generally too few data for reliable processing. Values on Greenland are mostly artifacts from too few data at times.

## 2.2. MODIS

Proto-Flight Model and Flight Model 1 of the Moderate

Resolution Imaging Spectroradiometer (MODIS) operate onboard the NASA-Terra launched in December 1999 and NASA-Aqua launched in May 2002, respectively [7]. The orbits of Terra and Aqua are near polar and sun synchronous. They acquire reflectance and radiance (atmosphere and surface) in the optical through thermal portion of the electromagnetic spectrum in 36 bands. Spatial (nadir) resolutions are from 250 m, 500 m and 1 km. Terra acquires its data with equator crossing times of 10:30 (daytime) with repeat pass 22:30 (nighttime). Aqua acquires its data with equator crossing times of 13:30 (daytime) with repeat pass 01:30 (nighttime). Both instruments have a viewable swath width of 2330 km. Temporal resolutions of the MODIS products are 1-day, 8-day, 16-day, 1-month, 4-month and 1-year. Detailed information can be found online at the Land Processes Distributed Active Archive Center USGS website<sup>1</sup>.

## 2.3. MODIS Fire Information for Resource Management System (FIRMS)

MODIS FIRMS produces daily near-real time and science quality fire location hotspots data from standard MODIS MOD14 (Terra) and MYD14 (Aqua) Fire and Thermal Anomalies products, Level 3 version 5 [8]. The fire detection algorithm of MOD14/MYD14 (version 2.4 as of Oct. 2006) uses radiances in thermal 4 μm (two channels) and 11 μm (one channel) for fire detection relative to background brightness temperature, 0.65, 0.68 and 2.1 μm water-body (glint false alarm rejection) and 11 and 12 μm for fire/cloud masks [9]. Daily FIRMS products of hotspot locations give the detection thermal brightness temperature in Kelvin, probability of detection quality flag and geolocations given by latitude and longitude relative to the WGS-84 reference ellipsoid.

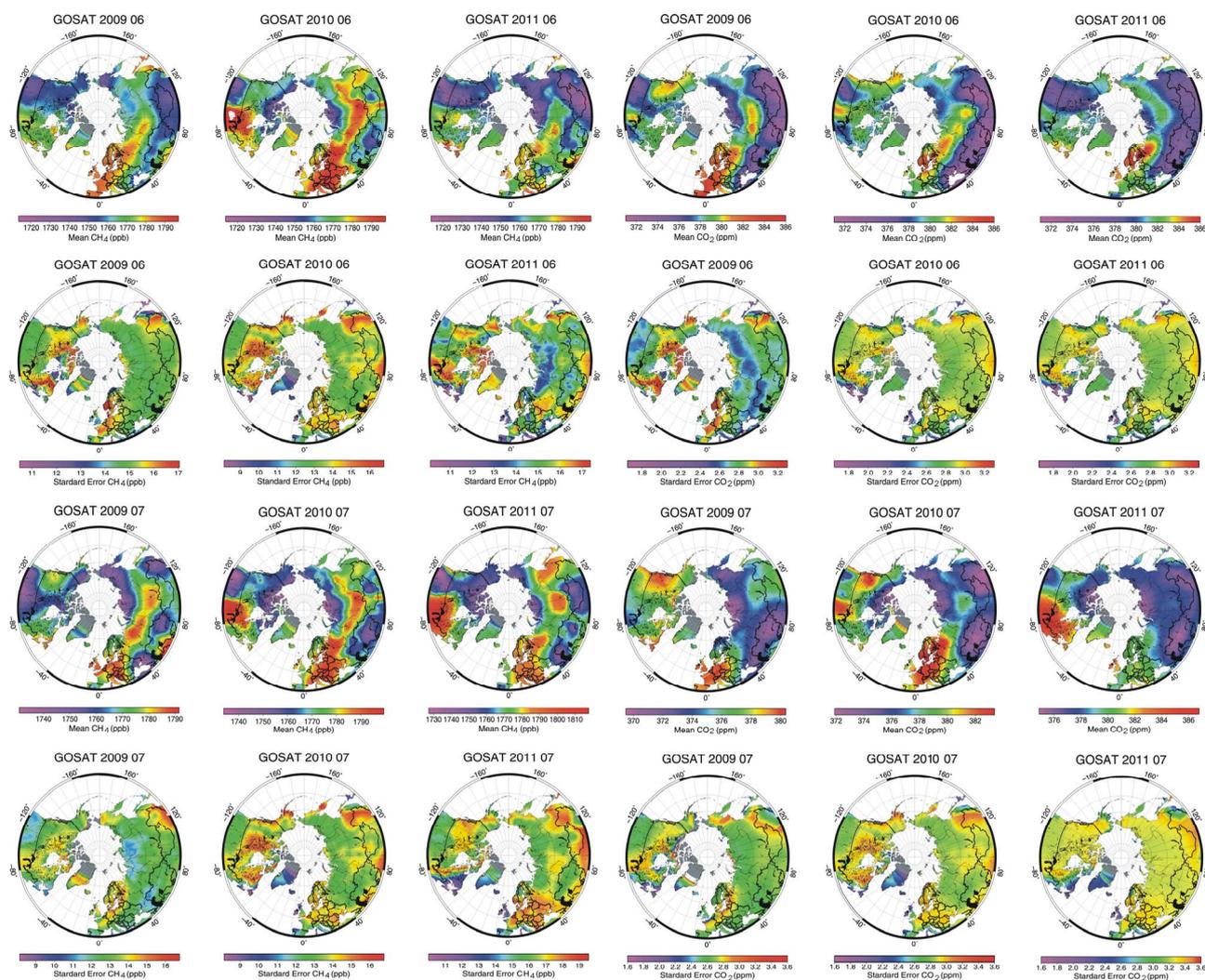
Validation campaigns have been carried out globally since 2000 using co-located thermal images from ASTER, Landsat 5/7 and Geostationary Operational Environment Satellites (GOES) for example [10]. On a global basis estimated commission errors vary with scan angle from 1.5% up to 8%. Commission errors in dense vegetation equatorial regions can approach 35% [11].

Our processing of the MODIS FIRMS data takes the daily detection files to produce June and July month fire-spot sums at 1 km resolution on the northern hemisphere, **Figure 2**. This gives us a monthly landscape and regional fire-spot geolocations for comparison to the GOSAT and other MODIS monthly datasets.

## 2.4. MODIS Evapotranspiration

The Global Terrestrial Evapotranspiration Data Set project of the Numerical Terradynamic Simulation Group at the University of Montana College of Forestry produces

<sup>1</sup>[https://lpdaac.usgs.gov/products/modis\\_overview](https://lpdaac.usgs.gov/products/modis_overview).



**Figure 1. GOSAT mean CH<sub>4</sub> and CO<sub>2</sub> mixing ratio (concentrations) and standard error (ppm) fields during June and July 2009-2011.**

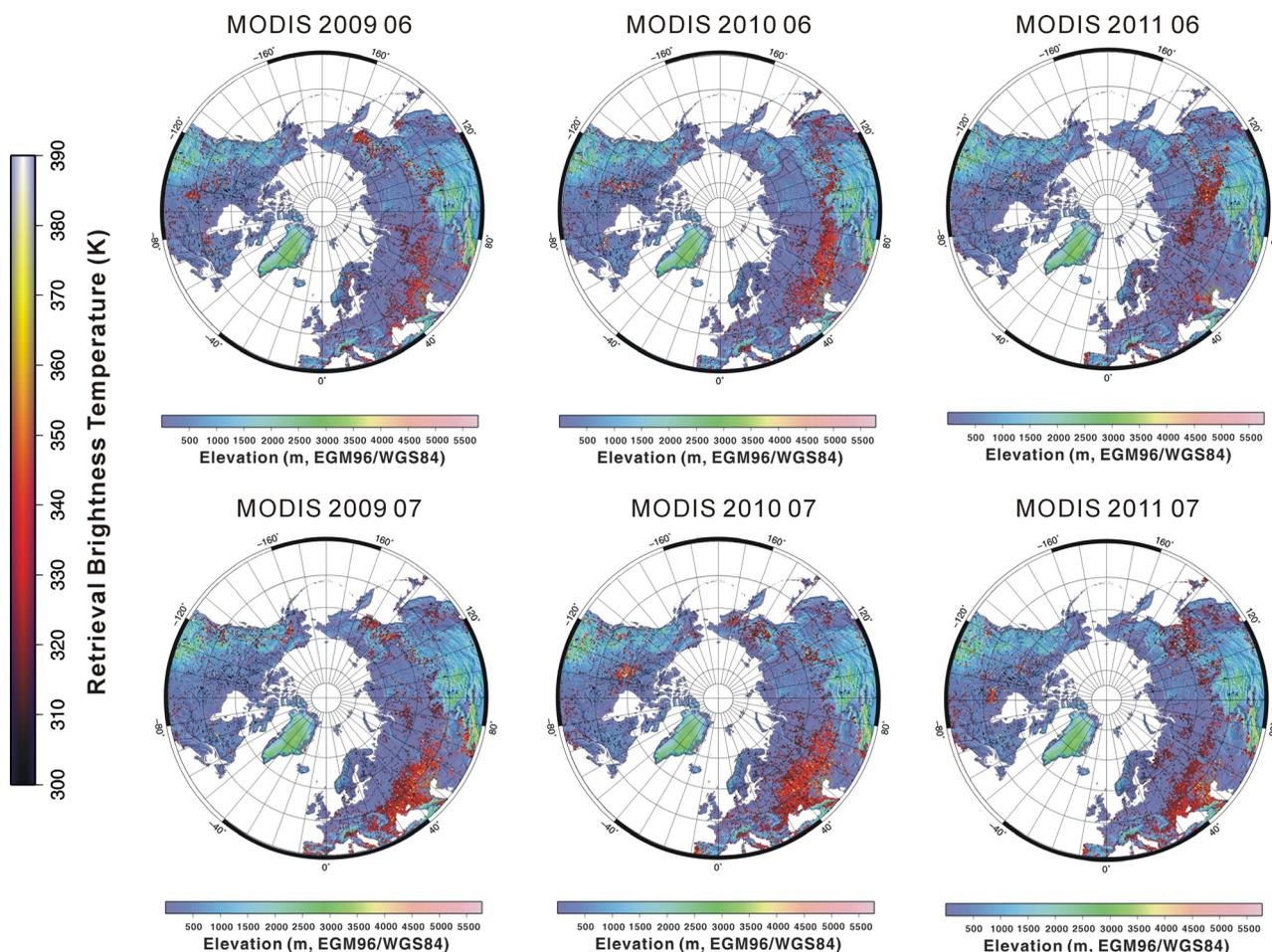
the MODIS MOD16 product Global Terrestrial Evapotranspiration for the NASA<sup>2</sup>. Files contain derived evapotranspiration (ET, mm/mo) and include latent heat flux (LE), potential ET and potential LE with quality control flags. The time period covered begins January 2000 and continues through the current year with about one to two month processing lag. Spatial resolution is 1 km with temporal resolution at 8-day, 1-month and 1-year intervals. Geolocation information is relative to the WGS-84 reference ellipsoid.

Since 2012 MOD16 ET is derived algorithmically based on the Penman-Monteith equation [12]. The algorithm derived ET includes surface evapotranspiration from vegetation stomata, evaporation from moist and wet soil, evaporation from open water bodies and evaporation of rain water on vegetation. During winter months,

<sup>2</sup><http://www.ntsug.umt.edu/project/mod16>.

evaporation from surface snow is included in ET. Cloud contamination estimates derive from albedo and fraction of photosynthetic available radiation/leaf area index (FPAR/LAI) MODIS inputs. Validations have been performed using available tower eddy covariance observations at FLUXNET stations. Correlation of tower measurements with the current version of the algorithm shows  $R = 0.86$  with the distribution centered on the 1:1 line. Errors in tower measured eddy covariance can themselves be as great as 30%. This arises from problems in the energy balance closure at tower sites and scaling issues when estimating landscape scale ET from specific and sparse site measurements.

Our processing of the monthly MOD16 files extracts ET and geolocation values into equal-area projection grids. We co-geolocate the GOSAT and MODIS grids on the WGS-84 reference ellipsoid. This is to ensure that for any geolocation there are paired ET rates, CH<sub>4</sub> and CO<sub>2</sub>



**Figure 2. MODIS FIRMS fire hotspot retrievals (K, 1 km resolution) during June and July 2009-2011 projected on the ACE2 DEM (1 km resolution).**

concentrations. The MODIS ET grids are shown in **Figure 3**.

**2.5. Regions of Interest in Eurasia**

Our investigation focuses on two regions of Eurasia: west-region 40° to 80°E by 50° to 58°N and east-region 100° to 140°E by 50° to 58°N. The west-region includes agricultural, industrial and urban area of Moscow and drained peatlands. The east-region is within the southern portions of the central Siberian Plateau and including heterogeneous terrains of the discontinuous permafrost zone, palsa and thermokarst lakes, and southern boreal and steppe ecosystems.

Wildfires are visually spectacular and physically dynamic events. We use these regions to investigate the wildfire events in June-July 2010 relative to 2009 and 2011 for affecting regional CH<sub>4</sub> and CO<sub>2</sub> levels, *i.e.* background concentration levels. We state this as a hypothesis to test: Did regional background levels of CH<sub>4</sub>, CO<sub>2</sub> and ET change in response to the June-July wild-

fires in Russia.

**3. Results**

Within the east- and west-regions we extract up to 299,800 co-geolocated CH<sub>4</sub> and CO<sub>2</sub> concentrations and ET rates in months June and July 2009, 2010 and 2011 and extract fire/wildfire hotspots. We then compute joint probability density functions of CH<sub>4</sub> with ET and CO<sub>2</sub> with ET. The joint probability density functions allow us to isolate the dominant mode, the most likely co-geolocated levels. We then compute least squares regression of the modes of CH<sub>4</sub> and ET with CO<sub>2</sub> with ET. We further plot the co-geolocation standard deviations of the dominant modes. This allows for evaluation of the test hypothesis.

**3.1. Fire/Wildfire Hotspots**

**Table 1** shows the counts of MODIS FIRMS fire/wildfire hotspots (1 km) in the east- and west-regions for June and July 2009, 2010 and 2011.

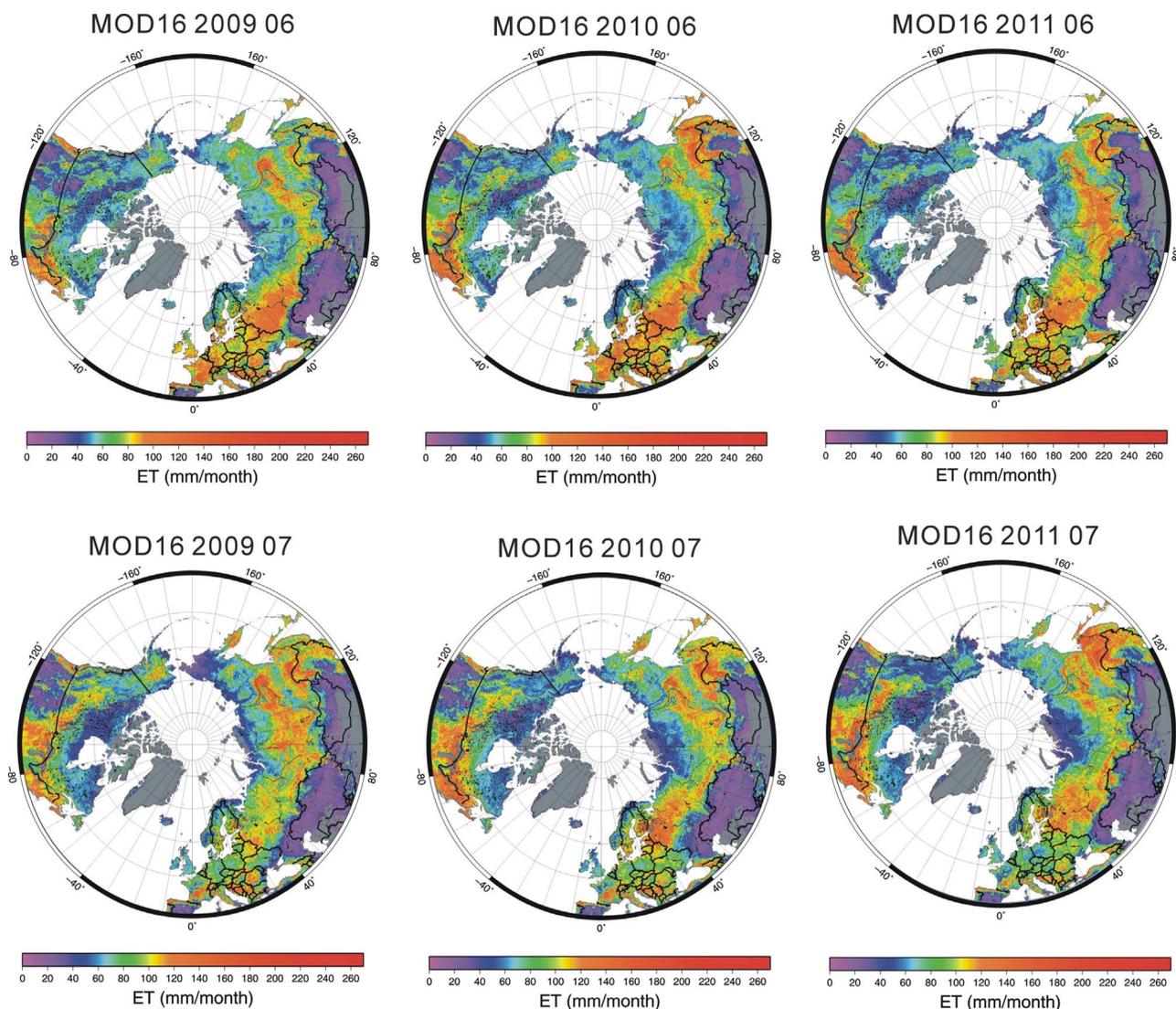


Figure 3. MODIS Evapotranspiration (mm/month) fields during June and July 2009-2011.

Table 1. Counts of fire/wildfire hotspots from MODIS FIRMS for regions 40° to 80°E by 50° to 58°N (West) and 100° to 140°E by 50° to 58°N (East).

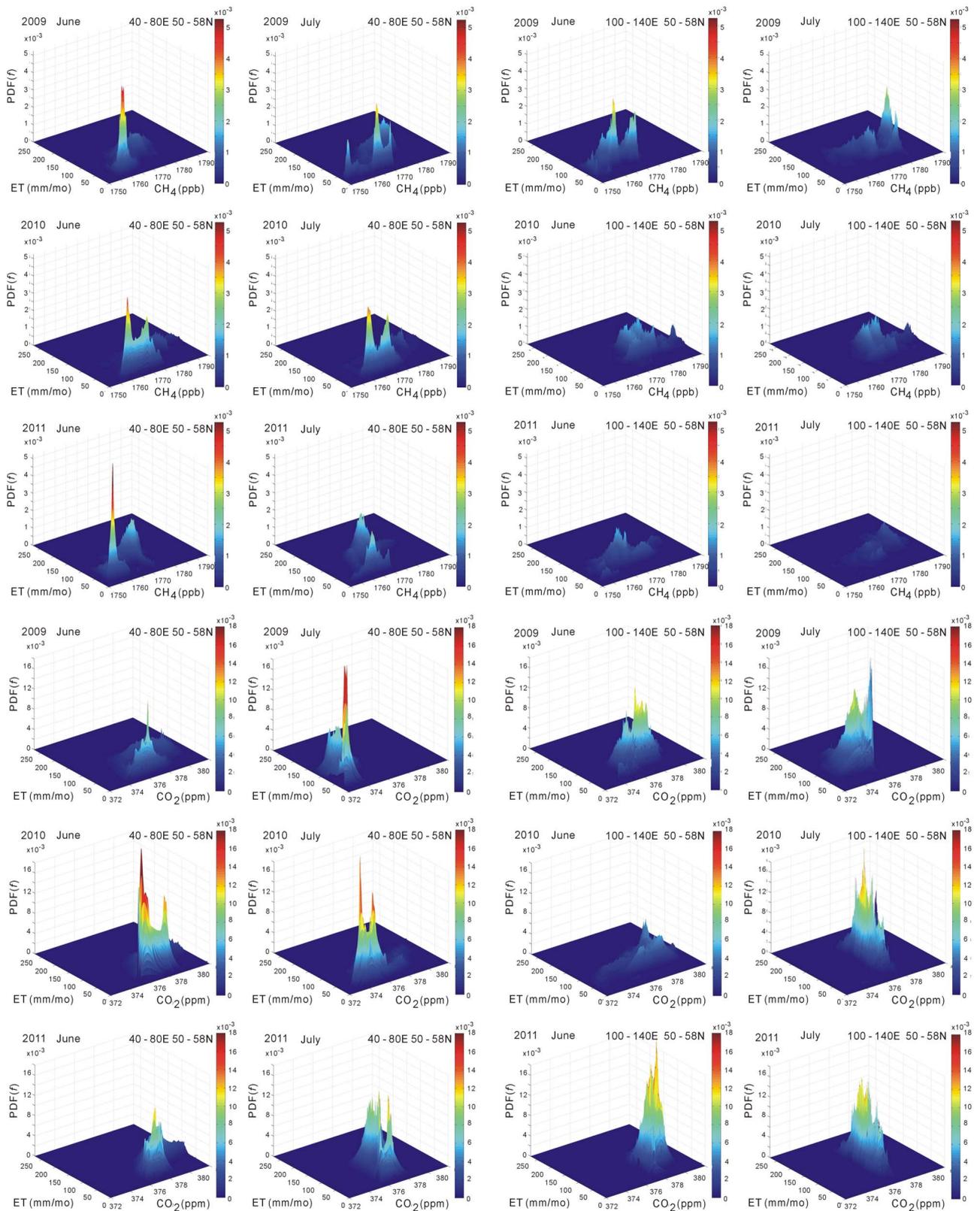
2009			
June		July	
West-Region	East-Region	West-Region	East-Region
2701	2705	5495	572
2010			
June		July	
West-Region	East-Region	West-Region	East-Region
8560	3915	18,022	1324
2011			
June		July	
West-Region	East-Region	West-Region	East-Region
992	12,679	3155	29,986

July counts in the west-region are greater by a factor of two above counts in June. In the east-region, July counts are lower by more than a factor of two than counts in June for years 2009 and 2010. In year 2011 the east-region July count is greater by a factor of two compared to the June count.

### 3.2. Joint Probability Density Functions

Figure 4 shows the joint probability density functions of CH<sub>4</sub> with ET and CO<sub>2</sub> with ET. The distribution change by month and year as well as the dominant modes, which represent values of co-geolocated CH<sub>4</sub> with ET and CO<sub>2</sub> with ET that are the most likely.

On inspection the dominant modes correspond to regional background levels of CH<sub>4</sub>, CO<sub>2</sub> and ET. Dominant modes do not correspond to any co-geolocated event or group of events such as a wildfire hotspot or industrial



**Figure 4. Joint probability density functions of MOD16 evapotranspiration (mm/month) and GOSAT CH<sub>4</sub> concentrations (ppb) and CO<sub>2</sub> concentrations (ppm), June 2009-2011 and July 2009-2011 in the region 40° to 80°E by 50° to 58°N and 100° to 140°E by 50° to 58°N.**

site (such as a coal-fire powered electrical generation facility, or a city in general). Events lend themselves to the spread of the probability distribution and have much lower probability, *i.e.* less significance.

### 3.3. Regressions of Dominant Modes

We next isolate the dominant modes in their values of CH<sub>4</sub> ET and CO<sub>2</sub> ET and plot these in **Figure 5**. **Table 2** also gives the values with their mode standard deviations. Modes occurring in the regions of interest are identified by symbol and the month is identified by color.

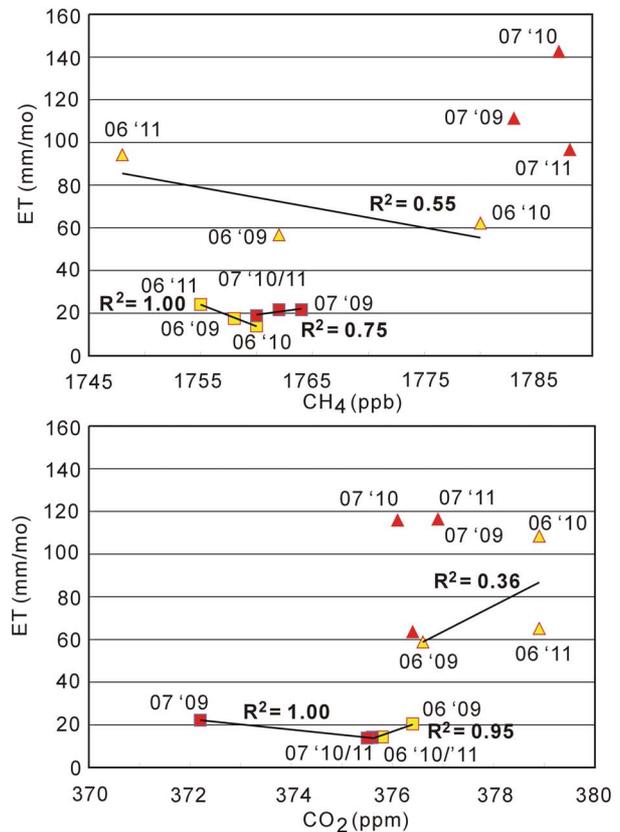
The west-region, which includes Moscow and locations of the 2010 fires show strong R<sup>2</sup> values. The regressions indicate inverse relationships of CH<sub>4</sub> to ET and CO<sub>2</sub> to ET in time of month and year. This equates low CH<sub>4</sub> and CO<sub>2</sub> concentrations to high ET rates.

The east-region, which includes much of Siberia and includes boreal and steppe ecosystems, locations wildfire activity and very numerous peatlands-wetlands and palsa and thermokarst lakes show very weak R<sup>2</sup> values. Significant forward or inverse relationships of CH<sub>4</sub> to ET and CO<sub>2</sub> to ET are not indicated.

The spread of the dominant modes of CH<sub>4</sub> ET and CO<sub>2</sub> ET are much greater for the east-region than for the west-region. This is an expression of the heterogeneity of eco-hydrologic systems, the discontinuous permafrost zone and the varied physical processes and terrain governing CH<sub>4</sub> and CO<sub>2</sub> concentrations and ET rates.

### 3.4. Standard Deviations of Dominant Modes

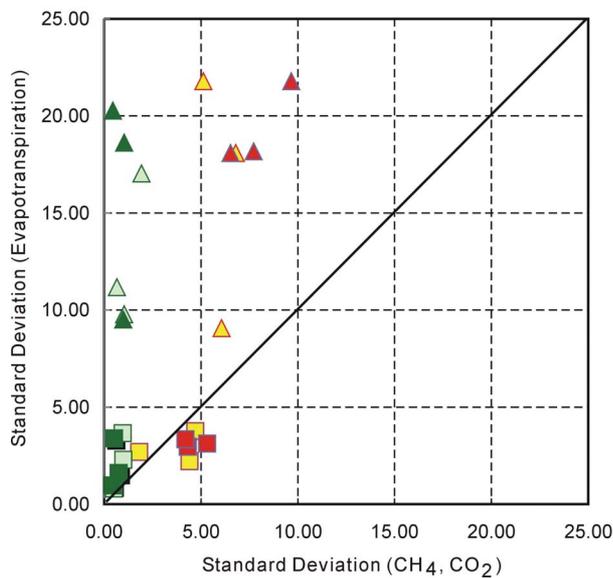
**Figure 6** shows the standard deviation comparison plot



**Figure 5. Regressions of dominant modes from joint probability density functions of MODIS Evapotranspiration (ET, mm/mo) with GOSAT CH<sub>4</sub> (ppb) and CO<sub>2</sub> (ppm) in regions 40° - 80°E by 50° - 58°N (Square) and 100° - 140°E by 50° - 58°N (Triangle) during June (yellow) and July (red), 2009, 2010 and 2011.**

**Table 2. Dominant modes and standard deviations of joint probability density functions of GOSAT CH<sub>4</sub> and CO<sub>2</sub> with MODIS Evapotranspiration (ET) during June and July 2009, 2010 and 2011 in west and east regions.**

	CH <sub>4</sub> (ppb)	ET (mm/mo)	CO <sub>2</sub> (ppm)	ET (mm/mo)
West Region				
2009 June	1758 ± 1.81	17.5 ± 2.69	376.4 ± 0.96	20.4 ± 3.65
July	1764 ± 5.31	21.6 ± 3.12	372.2 ± 0.51	22.2 ± 3.39
2010 June	1760 ± 4.41	14.0 ± 2.22	375.8 ± 0.98	14.3 ± 2.29
July	1760 ± 4.32	18.8 ± 2.96	375.6 ± 0.74	14.0 ± 1.61
2011 June	1755 ± 4.71	24.1 ± 3.76	375.6 ± 0.51	14.2 ± 0.81
July	1762 ± 4.20	21.6 ± 3.34	375.5 ± 0.44	13.8 ± 0.97
East Region				
2009 June	1762 ± 6.07	56.7 ± 9.07	376.6 ± 1.03	58.8 ± 9.78
July	1783 ± 7.72	111.3 ± 18.18	376.4 ± 0.99	63.6 ± 9.51
2010 June	1780 ± 6.79	62.2 ± 9.79	378.9 ± 1.93	108.4 ± 17.04
July	1788 ± 6.52	96.6 ± 18.09	376.1 ± 1.03	115.9 ± 18.63
2011 June	1748 ± 5.12	94.2 ± 16.0	378.9 ± 0.66	65.1 ± 11.17
July	1787 ± 9.66	142.5 ± 21.79	376.9 ± 0.45	116.4 ± 20.28



**Figure 6.** Standard deviation plot of dominant modes of GOSAT-derived  $\text{CH}_4$ ,  $\text{CO}_2$  to MODIS-derived Evapotranspiration in June (yellow, light green), July (red, dark green) of 2009, 2010 and 2011 in regions  $40^\circ - 80^\circ\text{E}$  by  $50^\circ - 58^\circ\text{N}$  (Square) and  $100^\circ - 140^\circ\text{E}$  by  $50^\circ - 58^\circ\text{N}$  (Triangle).

of the co-geolocated joint probability density function dominant modes. The standard deviations of the dominant modes in the west-region have restricted ranges and plot close to the 1:1 line near the origin. The east-region, while having a similar range in standard deviation  $\text{CH}_4$  and  $\text{CO}_2$  has a much larger range in standard deviation of ET.

This indicates that physical processes governing  $\text{CH}_4$ ,  $\text{CO}_2$  concentrations and ET rates in the west-region are different than those of the east-region. The background levels of  $\text{CH}_4$  and  $\text{CO}_2$  concentrations remain near constant, with some variation over June-July of 2009, 2010 and 2011. Evidence supports an inverse relationship of  $\text{CH}_4$  and  $\text{CO}_2$  concentrations to ET rates in the west-region.

#### 4. Discussion

In the summer of 2010 numerous media outlets including companies and governments in Europe and North America reported on the visually spectacular fires in southwest Russia particularly those southeast of Moscow<sup>3</sup>. There were many reported impacts to local air quality, and problems associated with the summer “heat wave” that accompanied the fires of that year [13].

Our investigation focus is on comparison and interrogation of GOSAT and MODIS datasets at both fine scale (landscape) and course scale (regional to hemispheric). Our analysis shows correspondence of increase-regions

of  $\text{CH}_4$ ,  $\text{CO}_2$  concentrations and ET rates during June and July 2009, 2010 and 2011. These increase-regions have correspondence to active-fire (wildfire, including forest and steppe and peatlands-wetlands including palsa and thermokarst lakes) during the same months and years. Fire/wildfire counts in June and July 2009, 2010 and 2011 show variations of a factor of two. The west-region in general has higher fire/wildfire counts in July 2010. The highest count of fire/wildfire occurs in July 2011 in the east-region.  $\text{CH}_4$ ,  $\text{CO}_2$  concentrations and ET rates are significantly higher in the east-region (Siberia) relative to the west-region (agricultural, industrial and urban) and during June and July in 2010, on average. Standard deviations of  $\text{CH}_4$  and  $\text{CO}_2$  concentrations occupy a narrow range in both regions of interest. Standard deviations of ET rates are much broader in the east-region relative to the west-region. This and the higher values of  $\text{CH}_4$ ,  $\text{CO}_2$  and ET point to heterogeneous ecosystems, hydrology including palsa and thermokarst lake processes, discontinuous permafrost zone, terrain of the east-region relative to the west-region. The east-region is well known for its long history of lightning ignited wildfires within larch-dominated communities of the central Siberian Plateau [14].

On average, background levels of  $\text{CH}_4$ ,  $\text{CO}_2$  concentrations and ET rates are not significantly perturbed during June-July 2010 relative to 2009 and 2011. This points to the buffering capacity of the lower atmosphere and terrestrial ecosystems [15-17]. Therefore, evidence does not at this time support the test hypothesis. In conclusion the fires and wildfires of summer 2010 did not significantly affect regional or hemispheric and multi-year background levels of  $\text{CH}_4$ ,  $\text{CO}_2$  concentrations and ET rates.

#### 5. Conclusion

We investigate the correspondence of GOSAT  $\text{CH}_4$  and  $\text{CO}_2$  near-surface concentrations and standard errors to MODIS ET rates and fire/wildfire hotspots on the northern hemisphere during June-July 2009, 2010 and 2011. Our month-year time frame is constrained by the sun-angle and cloud-free requirements of the Fourier Transform Spectrometer, spectral interferometer TANSO-FTS sensor onboard GOSAT. After processing our datasets to 5 km resolution we focus on two regions in Eurasia for analysis: west-region  $40^\circ$  to  $80^\circ\text{E}$  by  $50^\circ$  to  $58^\circ\text{N}$  and east-region  $100^\circ$  to  $140^\circ\text{E}$  by  $50^\circ$  to  $58^\circ\text{N}$ . Joint probability density functions identify significant dominant modes of  $\text{CH}_4$  and  $\text{CO}_2$  concentrations to ET rates. Evidence indicates that background levels of  $\text{CH}_4$  and  $\text{CO}_2$  are specific to each region and that these levels are near constant across the month-by-year observations. The west-region (agricultural, industrial and urban) shows strongly

<sup>3</sup><http://www.jpl.nasa.gov/news/news.php?release=2010-261>.

correlated inverse relationships of CH<sub>4</sub> and CO<sub>2</sub> concentrations to ET rates. CH<sub>4</sub>, CO<sub>2</sub> concentrations and ET rates in the east-region have significantly greater levels than those in the west-region. In the east-region (Siberia) evidence shows no relationships (forward or inverse) of CH<sub>4</sub> and CO<sub>2</sub> concentrations to ET rates. This is likely an expression of the heterogeneity of ecosystems, hydrology including numerous palsa and thermokarst lakes and varied physical processes and terrain of the discontinuous permafrost zone of the central Siberian Plateau.

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