

Regional Analysis of NASA Satellite Greenness Trends for Ecosystems of Arctic Alaska

Christopher Potter

Mail Stop 232-21, Moffett Field, USA Email: <u>chris.potter@nasa.gov</u>

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Abstract

Trends in the growing season MODerate resolution Imaging Spectroradiometer (MODIS) Enhanced Vegetation Index (EVI) time-series were analyzed for the period from 2000 to 2010 to understand landscape-level patterns of vegetation change in ecosystems of arctic Alaska. We compared datasets for vegetation cover types, wetland cover classes, wildfire boundaries since the 1940s, permafrost type, and elevation to identify the most likely combination of factors driving regional changes in habitat quality and ecosystem productivity. Approximately 57% of all arctic ecosystem areas in Alaska were detected with significant (p < 0.05) positive or negative MODIS growing season EVI trends from 2000 to 2010. Nearly all (99%) of these ecosystem areas (covering 178,050 km²) were detected with significant positive growing season EVI trends. The vast majority of the arctic Alaska region detected with significant positive growing season EVI trends was classified as upland tundra cover, although non-forested wetlands (marshes, bogs, fens, and floodplains) were co-located on 8% of that area. Herbaceous wetlands were co-located on 55% of the total area detected with significant negative growing season EVI trends, mostly on the arctic coastal plain and foothills. This evidence supports the hypothesis that temperature (warming) has markedly enhanced the rates of upland tundra vegetation growth across most of arctic Alaska over recent years.

Keywords

MODIS EVI, Arctic, Tundra, Wetlands, Wildfire, Alaska

1. Introduction

Major changes have been observed in Alaskan ecosystems over the past several decades, including rapidly

melting ground ice (permafrost), ice wedge degradation, wetland drying, and increased shrub growth [1]-[4]. Along the arctic coastal plain of Alaska, permafrost temperatures have increased 5°F increase since the 1980s [5]. Many areas in the continuous permafrost zone have seen increases in temperature in the seasonally active layer and a decrease in re-freezing rates. Alaska saw a 10% increase in precipitation from 1949 to 2005, with the greatest increases recorded in winter snowfall [6]. The length of the growing season (photosynthetically active period) in the Arctic has increased by 9 days per decade since 1982 [7]. Snow cover variability may be an important cause of such phenological change in the arctic tundra. Zeng and Jia [8] found greening onset was related to final snowmelt date and the end of the growing season was strongly related to the date of first snow cover.

Satellite remote sensing has been used to effectively monitor changes in high-latitude vegetation cover and productivity, especially following disturbance [9]-[14]. Like many others, Goetz *et al.* [15] [16] analyzed the seasonal and inter-annual variations of post-fire forest cover by using the Advanced Very High Resolution Radiometer (AVHRR) normalized difference vegetation index (NDVI) time-series across boreal North America and reported vegetation compositional changes consistent with early successional plant species and susceptibility to drought. Kim *et al.* [17] examined changing soil freeze-thaw signal from satellite microwave remote sensing and vegetation greenness patterns for the 9-year (2000-2008) vegetation record from satellites over North America, and reported that the relationship between the non-frozen period (June-August) and mean summer greenness index anomalies was generally positive for tundra and boreal forests areas of Canada.

In this study, we analyzed trends in the NASA MODerate resolution Imaging Spectroradiometer (MODIS) Enhanced Vegetation Index (EVI) satellite data record over the time period from 2000 to 2010 to understand landscape-level patterns of vegetation change in ecosystems of arctic Alaska. MODIS offers significant advances over AVHRR in spatial resolution and spectral information, particularly with respect to correcting surface images for cloud contamination [18] in persistently cloudy regions like the Arctic. We compared datasets for Alaska vegetation cover types (from MODIS), wetland cover classes from Whitcomb *et al.* [19], wildfire boundaries since the 1940s, permafrost type, and elevation, all to improve interpretations of changes in EVI green vegetation cover. The main objective of this multilayer spatial analysis was to help identify the most likely combination of factors driving long-term changes in habitat quality and ecosystem productivity across the arctic Alaska region (Figure 1).

2. Methods and Data Sets Used

Collection 5 MODIS data sets beginning in the year 2000 were obtained from NASA's Land Processes Distributed Active Archive Center site (LP-DACC, 2007) [20]. As described by Li and Potter [13], MODIS EVI values were aggregated to 8-km resolution from MOD13C2 (MODIS/Terra Vegetation Indices) products. MOD13C2 data are cloud-free spatial composites of the gridded 16-day 1-kilometer MOD13A2 product, and were provided monthly as a level-3 product projected on a 0.05 degree (5600-meter) geographic Climate Modeling Grid (CMG). Quality assessment (QA) for a cloud-free global coverage at 8-km spatial resolution was achieved by replacing clouds with the historical MODIS time-series EVI record. MODIS EVI was calculated from red, blue and NIR bands as described by Huete *et al.* [21]. Monthly EVI values were summed across each growing season period (May through September) to represent the variability in vegetation productivity for these past 11 years.

MODIS EVI was scaled from its original values to a range of 0 to 1.0 for time-series regression analysis. Slope values indicated the change in summed growing season EVI per year over the period 2000 to 2010. Positive slope values indicated greening over time, whereas negative slope values indicated declining green vegetation cover from 2000 to 2010. According to the coefficient of determination (R^2) of EVI time-based regressions, areas with a 95% statistical level of significance for a two-tailed t-test ($R^2 \ge 0.37$) consistently showed a negative or positive slope of greater than or equal to 0.02 EVI units per year [13].

Vegetation cover classes from 2009 MODIS imagery (1-km resolution; Friedl *et al.* 2002 [22]) and the wetland map developed from satellite radar by Whitcomb *et al.* [19] were added to the EVI trend analysis (**Figure 2**). Wetland types (1-km resolution) were listed by Whitcomb *et al.* [19] as combinations of estuarine (tidal inlet), lacustrine (lakes), or palustrine (bogs and fens) with scrub/shrub, emergent, or moss/lichen vegetation cover. Permafrost types for Alaska were mapped originally by Ferrians [23] and regridded to 8-km spatial resolution for this analysis. Elevation at 1-km resolution was derived from the United States Geological Survey (USGS) National Elevation Dataset (NED).



Figure 1. Major Land Resource Area (MLRA) coverage of arctic Alaska shown in the dark shaded areas.





Figure 2. Regional maps of (a) land cover, (b) wetlands, (c) elevation, and (d) permafrost types of arctic Alaska.

Areas burned by wildfires in interior Alaska since 1940 were delineated from the polygon boundary files of the Alaska Interagency Coordination Center (AICC) (available at <u>http://afsmaps.blm.gov/imf_firehistory/</u>). Prior to 1987, only fires in excess of 1000 acres (405 ha) coverage were included in this fire history database, whereas after 1987, all fires in excess of 100 acres (40.5 ha) were included.

3. Study Region

The Arctic zone is characterized generally by average annual precipitation of less than 50 cm and an average annual temperature of -6° C or less [24]. The arctic region of Alaska was delineated for this study as a total area of 317,380 km² within the Major Land Resource Areas [25] covering the Arctic Coastal Plain, Arctic Foothills, and Brooks Range Mountains (**Figure 1**). According to MODIS 1-km land cover mapping (**Figure 2**), this Alaskan arctic region is predominantly (99%) open tundra cover, 0.1% evergreen (coniferous) forest cover, and 0.7% barren land cover. Wetlands, as identified by Whitcomb *et al.* [19], covered 28% (89,180 km²) of the region, located mainly on the coastal plain, with 82% of these ecosystems classified as herbaceous and 18% as shrub covered wetlands across the region (**Figure 2**). About 85% of the region was located on continuous permafrost. Elevation across the region ranges from sea level on the coastal plain to 2500 m in the highest portions of Brooks Range Mountains. The AICC wildfire boundary record showed that only 1.4% (4563 km²) of the Alaskan arctic region was burned by fires between 1940 and 2010. During the period between 2000 and 2010 alone, an area of 1707 km² of the region was burned by wildfire, more than any other decade since before 1940.

4. Results

Approximately 57% (nearly 180,740 km²) of all area in arctic Alaska was detected with significant (p < 0.05) positive or negative MODIS growing season EVI trends from 2000 to 2010 (Figure 3). Nearly 99% of these ecosystem areas were detected with significant positive growing season EVI trends.

4.1. Negative EVI Slope

Only 1% of arctic Alaska ecosystem areas (nearly 2670 km²) were detected with significant negative growing season EVI trends, largely clustered on the Arctic Coastal Plain, north and east of 69°N and 154°W, which included northern portions of the Arctic National Wildlife Refuge (**Figure 4**). Areas detected with the most significant negative growing season EVI trends (<-0.04 units·yr⁻¹) were detected within the perimeter of the Anaktuvuk River wildfire (centered at 69.13°N, 150.61°W), which burned 1040 km² of tundra shrubland in the central Arctic Foothills during October, 2007. The time series of monthly MODIS EVI sampled from within this wildfire perimeter showed the nearly complete loss of green vegetation cover within the fire perimeter in late 2007, followed by steady regrowth in the subsequent years (**Figure 5**(a)).

Non-forested wetlands (marshes, bogs, fens, and floodplains) were co-located on 1467 km² (55%) of the total area detected with significant negative growing season EVI trends in the Alaskan arctic region. The vast majority (99%) of these non-forested wetlands with significant negative EVI trends were classified as palustrine emergent vegetation (herbaceous cover of sedges and grasses; Cowardin *et al.*, [26]). Elevation co-varied with the areas of significant negative EVI slope, showing that 90% such areas were located at less than 500 m; another 9% located at higher than 1000 m. All of the arctic Alaska areas detected with significant negative growing season EVI trends were classified with coverage of continuous permafrost cover.

To more closely examine ecosystems of arctic Alaska that have declined significantly in growing season greenness over the past decade, representative areas of tundra cover that were unburned since the 1940s were identified in the northern foothills of the Arctic National Wildlife Refuge (near 69.70° N, -143.43° W). The MODIS bi-weekly time series from 2000 to 2010 showed a steady decline in green cover beginning in 2004 (**Figure 5(b)**). Because the vegetation had not burned over at least the previous 60 years (AICC historical fire data, circa 2010), such declines detected in growing season greenness for these areas could not be attributed to vegetation loss from recent wildfires. Approximately 1630 km² of arctic Alaska ecosystems fell into this category of declining green vegetation cover outside of recent wildfires boundaries.

4.2. Positive EVI Slope

Out of a total of 178,050 km² in the arctic Alaska region detected with significant positive growing season EVI



Figure 3. Regional map of MODIS EVI slope results from 2000 to 2010 growing seasons over the arctic Alaska region. Negative slope categories indicated a decline in vegetation cover, whereas positive slope categories indicated a growth of vegetation cover.





trends (**Figure 3**), only 1% (1678 km²) were located in areas that had been burned by wildfires since the 1940s. Areas burned during the 1970s represented 44% of these positive EVI trends, while areas burned between 1990 and 2010 represented 41% of the total positive growing season EVI trends in arctic Alaska.

In contrast to areas detected in the arctic Alaska region with negative EVI trends, the vast majority (>90%) of areas detected with significant positive growing season EVI trends were classified as upland (non-wetland) tundra cover. Non-forested wetlands were co-located on only 8% (14,000 km²) of the total area detected with significant positive growing season EVI trends. One-half of these non-forested wetlands with significant positive EVI trends were classified as palustrine emergent vegetation and the other half as palustrine wooded (shrub) vegetation.



Figure 5. MODIS bi-weekly EVI composite for (a) burned arctic tundra (69.0567°N, -150.5231°W) within the Anaktuvuk River wildfire boundary of 2007; (b) unburned tundra vegetation cover identified with a significant negative growing season EVI slope, located within the Arctic National Wildlife Refuge (69.7025°N, -143.4251°W); and (c) unburned tundra (66.5011°N, 144.6069°W) vegetation cover identified with a significant positive growing season EVI slope, located within the Kobuk Valley National Park in the western Brooks Range (67.5799°N, -160.1490°W).

As with areas detected in the arctic Alaska region with negative EVI trends, elevation co-varied with the areas of significant positive EVI slope, showing that 86% such areas were located at less than 500 m; another 4% located at higher than 1000 m. The majority (94%) of the arctic Alaska area detected with significant positive growing season EVI trends was classified with coverage of continuous permafrost.

To more closely examine ecosystems of arctic Alaska that have increased significantly in growing season greenness over the past decade, representative areas of upland tundra cover that were unburned since the 1940s were identified within Kobuk Valley National Park (NP) in the western Brooks Range. This protected area between the Selawik National Wildlife Refuge and the Noatak National Preserve is noted for its historic caribou migration routes. The MODIS bi-weekly time series from 2000 to 2010 showed a steady increase in green cover after 2001, with growing seasons extending further and further into September after 2004 (**Figure 5(c)**). These upland areas were in undulating terrain (between 400 and 700 m elevation), underlain by continuous permafrost. Similar locations showing significant increases in growing season greenness over the past decade were detected within the southern Cape Krusenstern National Monument (67.42°N, -163.66°W), the Noatak National Preserve (68.22°N, -158.72°W), the Gates of the Arctic NP (67.52°N, -153.54°W), and the south-central Arctic National Wildlife Refuge (near 68.41°N, -144.86°W). Because the vegetation had not burned over at least the previous 60 years (AICC historical fire data, circa 2010), these increases detected in growing season greenness for such locations could not attributed to rapid regrowth after recent wildfires. Approximately 176,370 km² of arctic Alaska ecosystems fell into this category of significantly increasing green vegetation cover outside of recent wildfires boundaries.

5. Discussion

The MODIS EVI time-series data analyzed in this study provided a consistent large-scale metric of vegetation growth or decline trends across the arctic Alaska region. Among the relatively small coverage areas of growing season EVI decline trends within the region over the years 2000 to 2010, less than 50% could be explained by the recent (since 2000) impacts of wildfire. A notable fraction of non-forested wetlands were detected with significant negative growing season EVI trends on the arctic coastal plain and foothills.

On a watershed scale, Jorgenson *et al.* [27] used photointerpretation of arctic Alaska to classify ice-wedge degradation processes. In advanced stages of thawing, water-filled pits can form following the loss of a thin protective soil layer in tussock tundra and sedge marshes. Watershed studies by Bowden *et al.* [28] in the foo-thills of the Brooks Range implied that, under a warming climate in the Arctic, it is likely that small disturbances to surface soils and vegetation may promote accelerated subsurface warming that could lead to thermokarsting (subsidence of the surface), and in hilly or mountainous terrain could have important impacts on the stream ecosystems and headwater areas they drain.

Remote sensing analysis in the river valleys of Interior Alaska by Belshe *et al.* [29] led to the conclusion that, despite the thermal protection provided by the extensive organic soil accumulations, this insulation property is no longer sufficient under present climate conditions to keep the underlying (near-surface) permafrost from thaving and thermokarst from forming and expanding. With climate warming, an unfrozen soil layer can form, with water draining laterally on slopes, or forming pits, ponds, and fens on flatter lowland terrain. Vegetation change from thawing and permafrost collapse in lowland stands may follow one of (at least) two different pathways [30]: 1) tree mortality by "drowning" in stands adjacent to groundwater fens, with development of forb-dominated floating mats, or 2) tree loss into collapsed water-filled thaw pits, with colonization by early successional *Sphagnum* bog species under acidic conditions. These thermokarst-driven pathways of woody vegetation loss best explain the EVI decline trends detected outside of recent wildfire boundaries in the results presented above across the arctic Alaska region.

Succession and vegetation regrowth after wildfire did not account for a significant portion of the area of positive EVI growing season trends in the arctic Alaska region. Rather, unburned upland tundra of the western Brooks Range showed a particularly high coverage of positive EVI slope areas. From the results presented above, which expand upon the findings of Potter *et al.* [14], we hypothesize that temperature (warming) has enhanced the rates of upland tundra vegetation growth across arctic Alaska over recent years. In support of this hypothesis, evidence of shrub (alder, birch, and willow) expansion over the last half century has been documented through repeat photography of the past 50 years in Alaska [31]. Beck and Goetz [32] also postulated that the proportion of North America tundra areas increasing in productivity has steadily grown since 1982, reaching 32% of non-barren areas by 2008.

Hope *et al.* [33] and Jia *et al.* [1] both showed that an increase in satellite VI for tundra ecosystems was strongly associated with an increase in aboveground plant biomass. Nonetheless, increasing tundra biomass does not always merely indicate increasing shrub density, as shown by Beck and Goetz [32], whose detection of a regional greening trend appeared to be unrelated to shrub density. This suggests that tundra productivity may be increasing across a range of functional vegetation types in Alaska, including herbaceous cover.

The causal mechanisms of tundra growth enhancement were suggested by Tape *et al.* [31] as warming temperatures and deeper snow packs in recent years, which can promote increased microbial activity, which in turn can increase the availability of nutrients for plant uptake, particularly by small shrubs. Consistent with our MODIS EVI trend results from 2000 to 2010, the most notable expansions of shrubs detected by the photo-interpretations of these authors in Alaska were in geomorphic units subjected to regular non-fire disturbances in relative moist upland tundra.

Several consequences of changing tundra vegetation were reviewed by Sturm *et al.* [34]. More relatively dark-colored shrubs emerging above winter snow cover may reduce surface albedo and increase solar energy absorbed at the surface. Because shrubs allocate carbon to woody stems that have long turnover times compared with annual roots and the leaves of graminoid plant forms, increasingly shrub-dominated tundra is likely to assimilate carbon in a different way and store it for a different length of time than graminoid-dominated tundra [35]. Expansion of shrubland cover in arctic Alaska may also be altering forage quantity and quality for caribou herds [36].

In closing, it is worth noting that Wang *et al.* [37] reported on sensor degradation having had an impact on trend detection in North America boreal and tundra zone NDVI with Collection 5 data from MODIS. The main impacts of gradual blue band (Band 3, 470 nm) degradation on simulated surface reflectance was most pronounced at near-nadir view angles, leading to a small decline $(0.001 - 0.004 \text{ yr}^{-1}; 5\%$ overall between 2002 and 2010) in NDVI under a range of simulated aerosol conditions and high-latitude surface types. Even if this same sensor degradation problem affected MODIS EVI trends over the period of our analysis from 2000 to 2010, the apparent rate of greening in warming ecosystems of interior Alaska was not negated by such small, progressive changes in MODIS data quality. The methodology developed for mapping and characterization of arctic ecosystem growth trends in this study can be readily extended to other regions of the arctic-boreal zone, and over the next decade of Collection 6 MODIS data.

6. Conclusion

Ecosystems of arctic Alaska that have increased significantly in growing season greenness over the past decade were identified within Kobuk Valley NP in the western Brooks Range between the Selawik National Wildlife Refuge and the Noatak National Preserve. These upland areas were in undulating terrain (between 400 and 700 m elevation), underlain by continuous permafrost. Similar locations showing significant increases in growing season greenness over the past decade were identified for future ground-based monitoring studies within the southern Cape Krusenstern National Monument, the Noatak National Preserve, the Gates of the Arctic NP, and the south-central Arctic National Wildlife Refuge.

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