

Fault Movement Potential of Marzanabad Area, North Alborz, Iran

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Abstract

The major Quaternary faults of Marzanabad in the north Alborz can be classified based on their strikes into two sets: northwest and eastwest. In this paper, we use a model to evaluate their movement potential. Their theoretical model is based on the relationship between fault geometrical characteristics and regional tectonic stress field. The results show that Taleqan, Kandovan Chitan-Dozbon and Makaroud-Dalir fault zones are of the highest movement potential in the area. Also, the region where the fault zones have been intersected (northeastern part of study area) is prone to high seismicity; however, these fault zones don't have high movement potentials.

Keywords

Marzanabad, Faults, Alborz, Movement, Potential, Iran

1. Introduction

Seismicity is closely related to active Quaternary faults. This attracts many researchers to investigate the quantitative relationships between them. As a new parameter, FMP is defined to quantify earthquake risk along active faults by [1]. Therefore, we use it for evaluation of earthquake risk Marzanabad in the north Alborz, Iran. The landforms in this area are mainly controlled by two sets of Quaternary faults, striking northwest and eastwest, respectively (Figure 1).

The questions to be addressed in this paper are: 1) what are the activity levels of these faults? And 2) will these faults cause destructive earthquakes? Previous work regarding these topics was mainly based on field works [2]. In this paper, we use a new method to evaluate fault activity by considering the mechanical relationships between fault geometry and regional tectonic stress field. This method has been used to evaluate the fault

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movement potentials of all the major Quaternary faults Marzanabad in the north Alborz.

2. Materials and Methods

2.1. Fault Sets

Quaternary faults are well developed in Marzanabad area. They were classified into two sets based on their strikes: northwest and east -west. The northwest striking fault set comprises Mosha-Fasham, Taleqan, Varangeh Roud, Kandovan, Dona-Siahbisheh and Valiabad fault zones in the southern part of study area (Figure 1).

The east-west striking fault set are well exposed and can be traced intermittently for a long distance in nearly east-west direction. In this set, there are Nessen, Chitan-Dozbon, Sama-Majlar, Makaroud-Dalir, Dashte Nazir-Nater, Marzanabad, Azadkooh, Siahbishehfault zone in the northern part of study area (Figure 1).

In summary, all of these fault zones are active in current tectonic regime (CTR) and characterized by seismic events (Table 1). These events have shown that, seismic layer is in 20 - 30 km depth. This area has low to moderate earthquakes with intermediate frequency and repeat time.

In the following sections, we will evaluate the earthquake risk along these faults, and discuss which fault is most favored to move under the influence of present-day tectonic stress field. We make this evaluation based on the relationships between tectonic stress orientation and fault geometric properties.

2.2. Theoretical Model for Analysis of Fault Movement Potential

The fault movement potential (FMP) is closely related to tectonic stress (σ), fault plane geometry (G) and the

Table 1. Seismic events of study area (1957-2015).								
No.	Date (yyyy/mm/dd)	Time (UTC)	Latitude	Longitude	Depth	Magnitude	Reference	
1	1957/05/06	15:06:51.0	36.40	51.50		Ms:4.8	MEA	
2	1959/05/01	08:24:02.0	36.38	51.16	33	M:5.3	NOW	
3	1970/10/03	06:57:03.0	36.01	51.31	78	mb:4.1	ISC	
4	1981/08/04	18:53:59.0	36.45	51.27		mb:4.7	ISC	
5	1998/12/03	21:07:18.0	36.30	51.11	100	mb:3.9	ISC	
6	2004/05/28	14:47:51.0	36.49	51.37	10	ML:2.8	IIEES	
7	2004/05/28	16:45:21.0	36.26	51.46	28	ML:2.8	IIEES	
8	2004/05/28	17:34:51.0	36.48	51.36	22	ML:3.3	IIEES	
9	2004/05/28	17:51:21.0	36.27	51.38	15	ML:3.2	IIEES	
10	2004/05/28	19:47:05.0	36.43	51.40	25	mb:4.5	EHB	
11	2004/05/28	20:07:26.0	36.40	51.38	28	ML:3.2	IIEES	
12	2004/05/28	20:09:18.0	36.41	51.37	28	ML:3.1	IIEES	
13	2004/05/28	20:25:31.0	36.17	51.47	28	ML:3.4	IIEES	
14	2004/05/28	20:31:20.0	36.41	51.40	28	ML:3.2	IIEES	
15	2004/05/28	20:32:57.0	36.46	51.42	28	ML:3.3	IIEES	
16	2004/05/28	20:49:25.0	36.45	51.37	28	ML:2.8	IIEES	
17	2004/05/28	21:03:35.0	36.32	51.43	28	ML:3.5	IIEES	
18	2004/05/28	21:22:48.0	36.37	51.42	28	ML:2.5	IIEES	
19	2004/05/28	22:32:34.0	36.36	51.38	28	ML:2.6	IIEES	
20	2004/05/28	23:07:27.0	36.39	51.40	28	ML:2.6	IIEES	
21	2004/05/29	00:25:48.0	36.36	51.43	28	ML:3.4	IIEES	
22	2004/05/29	01:57:58.0	36.36	51.37	28	ML:2.9	IIEES	
23	2004/05/29	03:54:27.0	36.47	51.40	28	ML:3.3	IIEES	
24	2004/05/29	04:53:04.0	36.37	51.42	28	ML:3.7	IIEES	
25	2004/05/29	05:34:50.0	36.44	51.43	28	ML:2.8	IIEES	
26	2004/05/29	06:14:20.0	36.47	51.41	28	ML:3.4	IIEES	
27	2004/05/29	06:34:14.0	36.43	51.43	28	ML:3.2	IIEES	
28	2004/05/29	09:23:49.0	36.49	51.40	14	mb:4.7	EHB	
29	2004/05/29	10:20:40.0	36.37	51.41	28	ML:3.6	IIEES	
30	2004/05/29	11:01:32.0	36.42	51.38	28	ML:4.2	IIEES	
31	2004/05/29	12:56:43.0	36.50	51.42	28	ML:3.6	IIEES	
32	2004/05/29	13:34:24.0	36.39	51.39	28	ML:3.3	IIEES	
33	2004/05/29	14:56:46.0	36.45	51.44	28	ML:3.7	IIEES	
34	2004/05/29	15:03:49.0	36.50	51.41	28	ML:3.2	IIEES	
35	2004/05/29	15:23:49.0	36.50	51.35	28	ML:3.3	IIEES	
36	2004/05/29	15:41:03.0	36.46	51.42	28	ML:3.9	IIEES	
37	2004/05/29	17:30:26.0	36.49	51.44	28	ML:3.9	IIEES	
38	2004/05/29	18:38:07.0	36.45	51.37	28	ML:4.6	IIEES	
39	2004/05/29	18:42:45.0	36.43	51.41	28	ML:3.8	IIEES	
40	2004/05/29	19:58:19.0	36.46	51.45	28	ML:3.5	IIEES	
41	2004/05/29	22:55:19.0	36.43	51.36	28	ML:3.9	IIEES	
42	2004/05/29	23:05:46.0	36.48	51.36	28	ML:2.6	IIEES	

ontinued							
43	2004/05/30	02:41:27.0	36.46	51.45	28	ML:3.6	IIEES
44	2004/05/30	07:48:53.0	36.46	51.38	28	ML:3.5	IIEES
45	2004/05/30	08:57:06.0	36.18	51.49	28	ML:2.9	IIEES
46	2004/05/30	11:49:02.0	36.36	51.36	28	ML:2.8	IIEES
47	2004/05/30	13:09:55.0	36.41	51.45	28	ML:4	IIEES
48	2004/05/30	15:48:46.0	36.44	51.46	28	ML:3.5	IIEES
49	2004/05/30	16:18:24.0	36.26	51.44	28	ML:3.2	IIEES
50	2004/05/30	18:03:19.0	36.44	51.45	28	ML:2.7	IIEES
51	2004/05/31	08:12:16.0	36.46	51.26	14	ML:3.7	IIEES
52	2004/06/01	00:35:32.0	36.45	51.49	28	ML:3.2	IIEES
53	2004/06/01	02:41:28.0	36.37	51.46	28	ML:3.3	IIEES
54	2004/06/01	06:22:01.0	36.39	51.45	28	ML:2.6	IIEES
55	2004/06/01	10:57:46.0	36.45	51.41	28	ML:2.7	IIEES
56	2004/06/02	11:12:25.0	36.50	51.41	28	ML:3	IIEES
57	2004/06/05	00:10:45.0	36.40	51.40	28	ML:2.8	IIEES
58	2004/06/05	20:58:14.0	36.42	51.43	28	ML:3.2	IIEES
59	2004/06/06	06:08:21.0	36.44	51.41	28	ML:2.8	IIEES
60	2004/06/08	04:34:36.0	36.47	51.38	28	ML:2.5	IIEES
61	2004/06/24	02:36:28.0	36.47	51.42	28	ML:2.8	IIEES
62	2004/06/27	04:51:35.0	36.49	51.40	12	ML:3.3	IIEES
63	2004/07/11	13:16:46.7	36.35	51.31	14	ML:3.3	IIEES
64	2004/07/18	13:44:18.9	36.43	51.42	31	ML:2.5	IIEES
65	2006/03/07	19:58:42.7	36.30	51.41	33	ML:2.9	HEES
66	2006/03/08	00:04:20.8	36.44	51.38	14	ML:2.5	HEES
67	2006/12/25	20:06:54 5	36.12	51.31	14	ML:2.7	HEES
68	2006/12/25	20:15:48.2	36.12	51.32	14	ML:3.1	IIEES
69	2006/12/26	15:53:29.9	36.37	51.39	20	ML:2.5	IIEES
70	2007/06/04	08:04:16.6	36.37	51.32	34	ML:3.8	IIEES
71	2007/07/04	09:25:54 9	36.38	51.32	15	ML:2.5	IIEES
72	2007/07/04	09:41:22.8	36.31	51.33	13	ML:2.5	IIEES
73	2007/12/01	21:51:37.1	36.14	51.09	14	ML:2.9	IIFES
74	2007/12/01	22:12:45.8	36.08	51.07	13	ML ·2 8	UEES
75	2008/10/17	21:01:21.9	36.46	51.01	13	ML.2.8	IIEES
76	2008/10/19	22:01:21:9	36.42	51.01	15	ML .2.8	HEES
70	2008/10/19	22.28.49.3	26.15	51.01	13	ML.2.0	HEES
78	2008/10/21	10.14.10 5	36.41	51.25	14	ML.3	HEES
70	2009/02/02	10.14.10.5	26.20	51.40	14	ML:2.6	HEES
80	2011/08/11	07:50:48.0	26.40	51.33	14	ML.2.0	HEES
0U 01	2013/04/12	07.50.48.0	26.44	51.00	14	ML.2.9	HEES
8J 8J	2013/09/12	14.55.11 5	30.44 36.44	51.09	14	ML:2.9	HEES
83	2013/09/12	06.33.26.2	36.44	51.12	14	ML:2.5	IIEES
84	2013/09/20	12:42:35.0	36.45	51.17	14	ML:2.9	IIEES
85	2013/12/12	06:18:05.9	36.23	51.43	14	ML:2.5	IIEES
86	2014/08/29	22:57:04.1	36.26	51.50	14	ML:2.6	IIEES

physical property of the medium within and on both sides of the fault (P). FMP is the function of these factors [1]:

$$FMP = f(\sigma, G, P) \tag{1}$$

Although a geological medium is generally heterogeneous and very complicated, however it can be taken as homogeneous and isotropic in statistical view of our case. Based on this consideration, and for the purpose of simplification in the theoretical derivation, they also take the geological medium containing the faults as a homogeneous, isotropic and elastic material. Therefore fault movement potential can be simplified as:

$$FMP = f(\sigma, G) \tag{2}$$

Finally, according to [3] and [4] researches, [1] defines FMP to quantify the relationship between fault movement potential as a normalized factor by the following equations:

$$FMP = \begin{cases} 0 & \theta \in (0^{\circ}, 30^{\circ}) \\ \frac{\theta - 60^{\circ}}{30^{\circ}} & \theta \in (30^{\circ}, 60^{\circ}) \\ 1 - \frac{\theta - 60^{\circ}}{30^{\circ}} & \theta \in (60^{\circ}, 90^{\circ}) \end{cases}$$
(3)

 θ is the angle between the regional maximum principal compressive stress orientation (σ_1) and the normal line of fault plane.

2.3. Regional Tectonic Stress Orientations

Tectonic stress means an additional stress to lithostatic stress state, in the other words, the part of stress deviated from lithostatic stress. Earthquake focal mechanism solution is one of the commonly used methods in the study of contemporary tectonic stress field. Therefore, we use results of [5]-[7] and our field study to estimate the regional maximum principal compressive stress orientation (σ_1). The statistical result shows that the average attitude of σ_1 is 15°, 040°.

This area belongs to West-Central Alborz and lesser Caucasus hinterland [8] [9] that formed on the inverted back arc intra-continental rift since Oligocene. Dominant structural trend in West-Central Alborz and lesser Caucasus province is NW-SE (Figure 1). From tectonics view, it contains deformed zone (fold and thrust belt) of Cimmerian miniplate that formed in northern active margin until late Triassic. Then it has rifted by tension in a back arc basin of Neotethyian subduction zone in the south margin of Cimmerian miniplate.

Development of that rift stopped in the late Cretaceous and then, renewed in the Eocene by spreading in submarine arc basin of Neotethyian subduction zone. In the other word, this hinterland is result of a magmatic arc system spreading in the evolutional back arc basin. After that, this region has converted to back arc regime again and West-Central Alborz and lesser Caucasus hinterland has formed by its deformation and regional uplift from SW part of Caspian Sea to Black sea. Recently, Damavand and Sebalan cones have formed by late volcanism that related to final subduction of oceanic slab (Neotethys) toward north and northeast [10]. This area has an active tectonics regime [11]-[25] in compared to the Central Iran [26]-[35] and Zagros in the southern Iran [36]-[43]. Also, some concepts of its metal mineralization, have been investigated by [44]-[48].

3. Results and Discussion

The fault movement potential of the major Quaternary faults of Marzanabad have calculated using the equations (3) and the regional stress orientation as well as the fault plane attitudes. The results are shown in **Table 2** and **Figure 2**. Also, seismic events have 25 Km depth as average of focal depth. The earthquakes are related to two sets of Quaternary faults with northwest-southeast and east-west strikes.

Taleqan, Kandovan, Chitan-Dozbon and Makaroud-Dalir fault zones have large angle between the normal to the fault planes and the compressive principal stress along these fault zones. The fault movement potential of this fault set ranges from medium to high, suggesting that this fault set has the sufficient potential for generating destructive earthquakes. This situation has been shown by contoured map, too (Figure 3).

Table 2. The calculation of fault movement potential in Marzanabad in the north Alborz.						
Name of Fault Zone	No.	Dominant Attitude of Fault (Dip Dir.)	Normal Line of Fault Plane	θ	FMP	
	01	017,80	10,197	39	0.3	
Masha Fasham	02	022,77	13,202	33	0.1	
Mosna-Fasham	03	025,67	23,205	41	0.4	
	04	009,62	28,189	52	0.7	
	05	210,58	32,030	19	0.0	
Talagan	06	191,50	40,011	36	0.2	
Taleqali	07	162,39	51,342	58	0.9	
	08	182,35	55,002	50	0.7	
Voronach Doud	09	028,70	20,208	37	0.2	
varangen Koud	10	025,80	10,205	29	0.0	
Azadkooh	11	181,77	13,001	38	0.2	
AZaukoon	12	180,78	10,000	39	0.3	
	13	002,53	37,182	63	0.9	
	14	027,63	27,207	44	0.5	
Kandovan	15	024,67	23,204	41	0.4	
	16	004,47	43,184	67	0.8	
	17	019,58	32,199	51	0.7	
Dona-Siahbisheh	18	020,72	18,200	38	0.3	
Siahbisheh	19	175,60	30,355	44	0.5	
Valiabad	20	016,82	08,196	33	0.1	
N	21	009,78	12,189	41	0.4	
INessen	22	004,80	10,184	44	0.5	
Chitan Dashan	23	183,55	35,003	39	0.3	
Chitan-Dozbon	24	155,61	29,335	61	0.9	
Course Mailer	25	358,55	35,178	64	0.8	
Sama-Majiar	26	008,72	18,188	46	0.5	
	27	154,58	32,334	62	0.9	
Makaroud-Dalir	28	138,63	27,318	76	0.5	
	29	192,62	28,012	29	0.0	
	30	197,45	45,017	36	0.2	
Dashte Nazir-Nater	31	131,60	30,311	82	0.3	
	32	166,42	48,346	55	0.8	
Marzapabad	33	182,58	32,002	38	0.3	
iviai zallaUau	34	166,70	20,346	52	0.7	







Figure 3. FMP contoured map of Marzanabad area.



Figure 4. Epicenter of earthquakes on FMP contoured map of Marzanabad area.

In Figure 4, Epicenter of earthquakes (Table 1) on FMP contoured map of Marzanabad area have shown and it implied that high FMP values are not always coincident to previous seismic history.

4. Conclusion

Based on this research, the contemporary movements potential along fault zones of various orientations are different under the action of present-day regional north-northeast compressive stress field in study region. Taleqan, Kandovan, Chitan-Dozbon and Makaroud-Dalir fault zones have medium to high movement potentials. Also, the region where the fault zones have been intersected (northeastern part of study area) is prone to high seismicity; however these fault zones don't have high movement potentials.

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