



# Impact of Urbanization on the Hydrological Cycle of Migina Catchment, Rwanda

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## Abstract

In this study, we carry out an assessment of the impact of urbanization on Butare, southern Rwanda and evaluate its effect on the hydrological process at the Migina catchment. We used data from the meteorological stations of the Migina catchment, land-use maps of the region (obtained in 1974 and 2010) and GIS technology in order to analyse hydrological fluctuations in Migina catchment and the expansion of Butare town. We observed that, between the two hydrological years (1974 and 2010) water runoff increased by over 3.5% with increasing flood risks. Furthermore, evapotranspiration and groundwater level have decreased by about 3% and 0.5% leading to water scarcity. Research on urban planning is necessary in order to improve on water resources management and reduce the effects of urbanization on water resources such as in the Migina catchment.

## Keywords

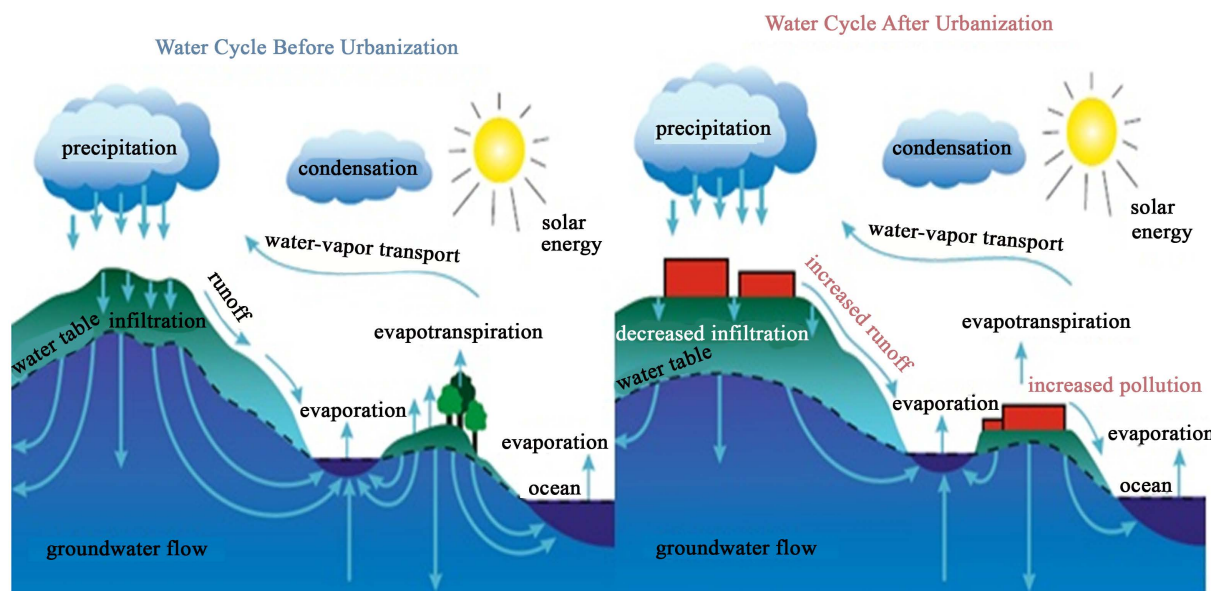
Water Resources, Urban Planning, Urbanization, Hydrological Cycle, Land Cover, Migina Catchment Rwanda

**Subject Areas:** Agricultural Science, Biodiversity, Conservation Biology, Ecology, Environmental Sciences, Hydrology, Natural Geography

## 1. Introduction

Rain water, either returns to the atmosphere through evapotranspiration or infiltrates the soil, runs off on the surface and likely ends up into a river, ocean or lake (Figure 1) [1]-[3]. The relative proportions depend on the nature of the surface, duration of rainfall and intensity [1] [3] [4].

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**Figure 1.** Water cycle diagram showing different processes involved before and after urbanization. Sourced from UNCE, 2015.

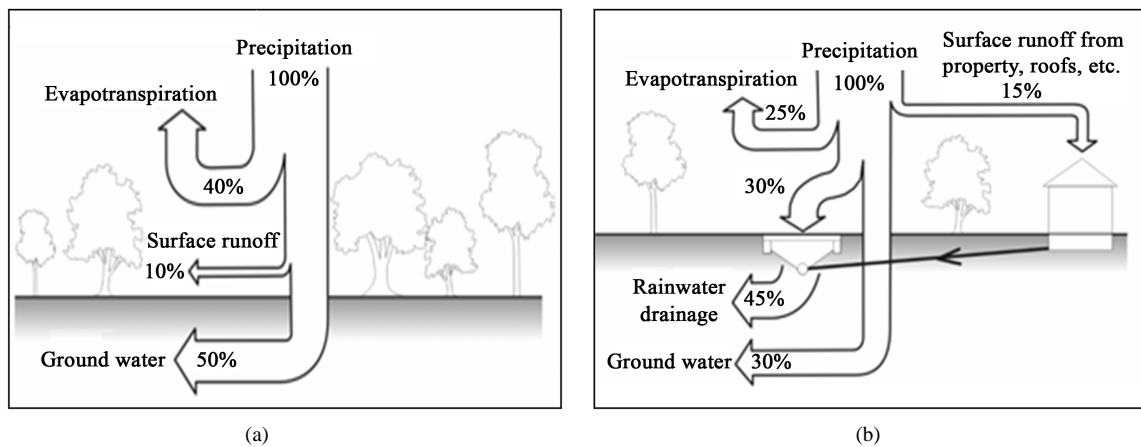
Development of an urban area involves covering the ground with artificial surfaces, and it significantly increases the amount of surface runoff in relation to infiltration and evapotranspiration (**Figure 2**). This therefore increases the total volume of water reaching the river when or after it rains. As a result, this will reduce water that would be stored within a catchment and contribute to water scarcity [5]. Furthermore, the construction of storm sewers and culvert of natural streams and channels which takes place during urbanization results in water being transmitted into the drainage network more rapidly; this increases outflow velocity [6] [7]. Since a larger runoff is discharged within a shorter interval, peak rate of water inevitably increases, giving rise to the danger of sudden flooding of rivers. These subsequent floods damage human and natural resources as well as properties [8].

When people develop within a watershed, the hydrologic cycle is affected. The increase in impervious or hard surfaces, including rooftops and pavement (roads, driveways, and parking lots) decreases the amount of water that soaks into the ground or infiltrates [2] [9]. This increases the amount of surface runoff (**Figure 3**). If infiltration is significantly decreased, groundwater levels may decline, affecting stream flows especially during the dry season. Meanwhile, lowered groundwater levels can result in subsequent well failures [1] [10]. While the effects of urbanization on the water cycle can be major, if wise choices are made during the development process, the impacts can be minimized and future water supply can be protected [10] [11].

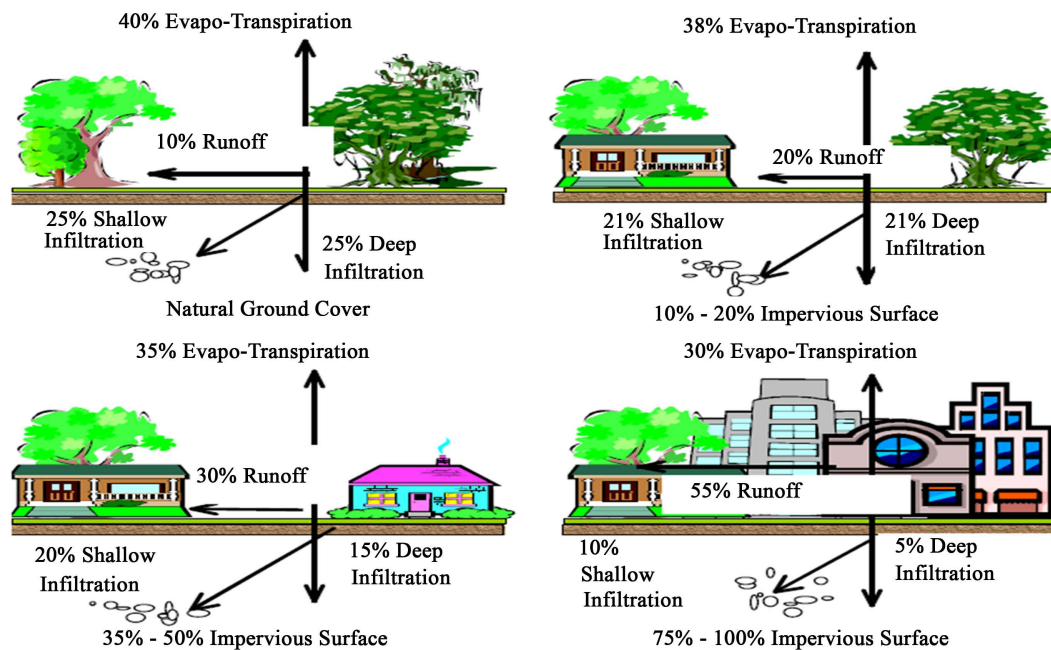
Although Rwanda is among the least urbanized countries with an urban population of 28.8% in 2015 [12], it had a rapid population growth of about 2.7% in 2010-2015 (UN, 2016) and a relatively high urbanization rate of 6.43% in 2015 [12] [13]. On the other hand in Rwanda, observations from existing data showed that over the last 30 years, some parts of Rwanda have experienced unusual irregularities in climate patterns including variability in rainfall frequencies and intensity, persistence of extreme weather conditions such as heavy rainfall in the northern regions and drought in the eastern and southern regions [14].

The meso-scale Migina catchment (257.57 km<sup>2</sup>) located in southern Rwanda experiences serious human activities in its northern zone. This is as a result of the rapid urbanization of Butare town as it progressively destroys the natural land cover surrounding the catchment area. The topography and land cover of drainage basins affect phenomena like erosion sediments that are transported from storm runoffs towards rivers. These produce changes in water balances of the concerned drainage basin [15]. Also, as new development projects are continuously being implemented within the area, their growing impact affects the hydrologic cycle in general [3] [10] [16].

The importance of the hydrological cycle has been in existence for billions of years and continuously renews and refreshes finite water supplies [10]. Therefore, lack of public awareness on implication of their daily activi-



**Figure 2.** Effect of urbanization on hydrological cycle. Sourced from Carlos, 2007. (a) Pre-urbanization; (b) Urbanization.



**Figure 3.** Influence of urbanization on different components of the water cycle. Sourced from CWP, 2015.

ties on water resources has contributed to its inefficient use and scarcity. This paper looks at urbanization of Butare town and studies its effects on the long term availability of water resources within the catchment. The study strives to espouse scientific evidence that would aid the local population, urban planners and decision makers, such as to promote cleaner and safer urban expansion while ensuring continuous water availability within the Migina catchment.

## 2. Methodology and Data Analysis

### 2.1. Study Area

Migina catchment is located between 2°32' and 2° 48' South, 29°42' and 29°48' East, in the Southern province of Rwanda. It is drained by perennial streams with the main flow following a north to south direction. The main stream is located in the eastern part of the catchment with most of the valleys draining from north-west to south-east towards the main stream. The Migina catchment drains into the Akanyaru River, which forms the border between Rwanda and Burundi. Downstream, the Akanyaru River joins the Nyabarongo River to form the

Kagera River, which is the largest tributary of Lake Victoria and a precursor of the Nile River, **Figure 4**.

The Migina catchment area is mountainous with elevation ranging from 1434 m in the south to 2251m around the Huye Mountains in the North West. Land use is dominated by pasture and arable farming such as rice, sorghum, maize, and sweet potato [17]. Butare town, with over 100,000 inhabitants is located in Huye district, stretching along the main axes of Akanyaru, Nyabarongo and Kagera Rivers.

The meteorological data recorded from different meteorological stations in the catchment (Butare airport, Butare mission, Kansi and Rubona) shows that the mean annual temperature is around 24°C and shows a progressive increase in the area's temperature, **Figure 5** [18].

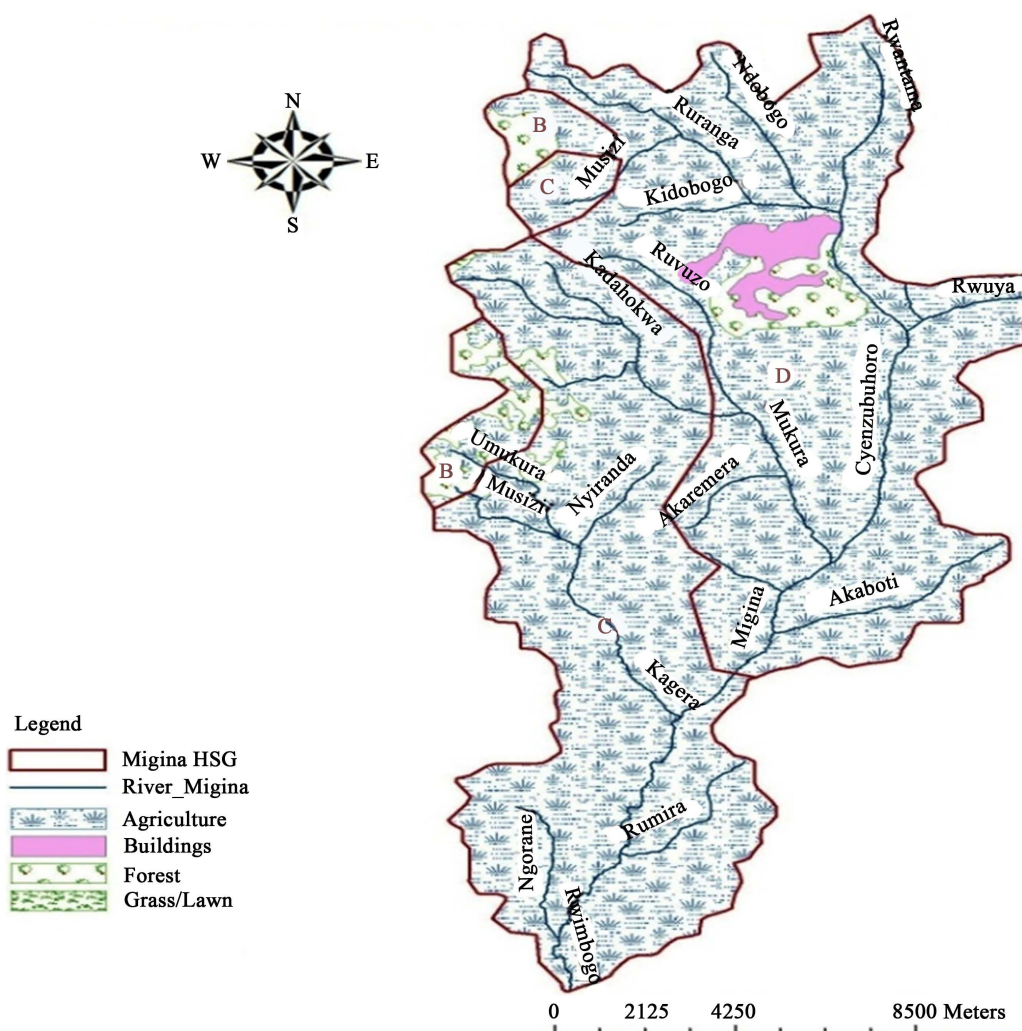
## 2.2. Determination of Two Annual Water Balances

The water balance is an account of the inputs and outputs of water, and it often applies to a watershed or a surface water reservoir. Any difference between the inflow ( $I$ ) and the outflow ( $O$ ) over a certain time period represents the water that has entered or left stored ( $S$ ) within the catchment [19].

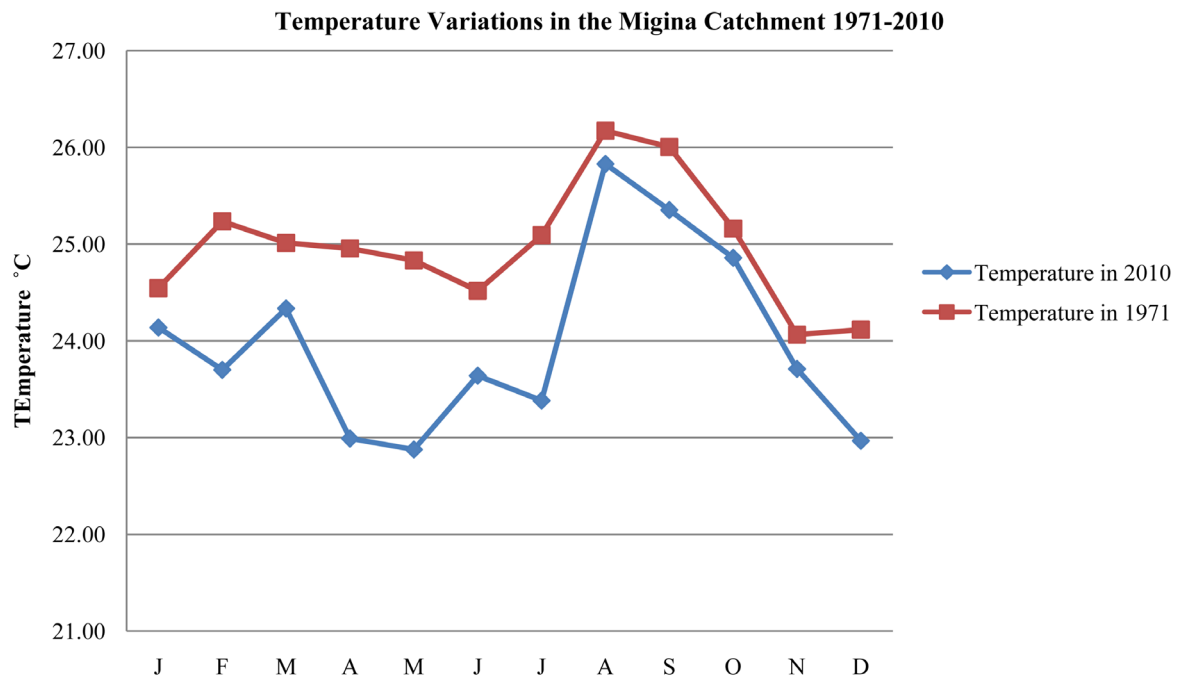
$$I - O = \pm S \quad (2.1)$$

According to [15], the water balance may also be written as;

$$A(P - E) - Q = \Delta S \quad (2.2)$$



**Figure 4.** Land cover of Migina catchment.



**Figure 5.** Mean annual temperature variation within the study area between 1971 and 2010.

The present study considers Inflow as Precipitation ( $P$ ), as Evapotranspiration ( $E$ ) + Outlet discharge ( $Q$ ), while Stock ( $S$ ) is the quantity of water stored in the catchment. The components of water balance were thus estimated by:

The measured evaporation ( $E_{pan}$ ) has to be compensated with a pan factor ( $f_{pan}$ ) to obtain the real open water evaporation [20].

$$ET_{ref} = f_{pan} E_{pan} \quad (2.3)$$

We used the most popular method for the calculation of the potential evapotranspiration ( $PET$ ), which computes the crop evapotranspiration according to [16]. Where  $k_c$  is the crop factor and  $ET_{ref}$  is the reference evapotranspiration.

$$ET_{pot} = k_c ET_{ref} \quad (2.4)$$

The actual evaporation ( $AET$ ) is then calculated by:

$$AET = PET \times K_{adj} \quad (2.5)$$

The Rational Formula for estimating peak runoff rate, introduced in the USA in 1889 has become widely used as a tool for drainage design, particularly for sizing water conveyance structures [21].

$$Q = KI \sum_{i=1}^n C_i A_i \quad (2.6)$$

where,  $Q$  is the peak runoff rate ( $m^3/s$ ),  $C$ , the runoff coefficient (unit less, ranging from 0 to 1) while  $I$  is the rainfall intensity ( $mm/hr$ ), and  $A$  is the watershed area ( $ha$ ). The conversion factor  $K$  is an estimated value 0.00278 (unit less).

### 2.3. Determination of Inflow

Rainfall data was collected from 12 meteorological stations located within the catchment area between June 2009 and May 2011 and Thiessen polygons were used to estimate the mean annual rainfall, [Table 1](#).

Similarly, the quantity of water that entered the catchment between 1973 and 1974 was obtained by collecting data from four rain gauges installed at Save, Butare airport, Nyakibanda and Kansi. Thiessen polygons were then used to calculate the mean (average) annual rainfall, [Table 2](#).



**Table 1.** Annual rainfall for 2009/2011 using Thiessen polygon method.

YEAR	Month	Rango P (mm)	Mubumbano P (mm)	Murama P (mm)	Vumbi P (mm)	Mpare P (mm)	Sovu P (mm)	Save B P (mm)	Muyira P (mm)	Kibilizi P (mm)	Gisunzu P (mm)	Rwasave P (mm)	Kansi A P (mm)
2009	June	6.73	7.57	4.46	5.72	10.94	12.79	4.04	18.68	7.41	21.60	4.00	20.20
	July	1.01	0.00	0.00	0.00	0.17	2.52	0.67	0.84	5.55	0.17	1.60	0.00
	August	22.55	48.22	50.16	26.93	38.71	4.38	2.78	40.06	25.00	28.61	3.50	55.63
	September	50.24	59.58	56.89	44.27	47.63	55.71	28.95	42.25	43.09	48.05	37.70	89.38
	October	134.99	94.51	91.99	66.49	78.02	151.82	121.44	55.38	64.47	132.82	129.30	113.00
	November	202.57	116.81	122.70	191.38	250.96	211.41	201.81	112.60	110.16	141.72	168.50	113.50
	December	123.55	120.01	131.96	119.34	132.38	134.65	136.42	128.76	63.62	127.74	51.70	111.10
2010	January	199.29	209.81	182.46	335.29	226.39	151.99	190.28	213.60	148.96	219.68	134.40	147.13
	February	208.38	260.64	238.93	237.66	183.44	180.52	183.63	280.25	167.98	253.48	288.70	288.20
	March	320.90	161.59	159.99	153.00	148.04	158.05	84.92	229.42	134.07	161.43	205.50	194.06
	April	390.92	168.23	144.92	226.39	245.67	158.81	123.55	196.17	142.57	181.87	112.60	138.88
	May	194.26	259.21	177.45	154.01	234.67	163.32	111.66	214.89	176.31	141.10	172.22	166.38
	June	32.08	33.41	53.81	46.79	27.50	23.38	11.92	11.00	7.70	27.50	17.88	37.58
	July	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
	August	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
	September	256.67	77.09	103.03	51.00	74.25	126.50	103.13	66.92	59.13	79.48	106.33	97.81
	October	231.00	51.84	114.31	47.30	90.75	107.71	53.17	62.33	45.83	45.93	92.13	52.34
	November	286.92	147.45	157.12	108.40	141.17	39.88	115.96	172.33	99.88	120.54	138.42	89.83
	December	151.25	87.36	108.63	109.49	154.00	77.46	185.17	70.13	103.22	119.63	107.25	93.59
2011	January	196.17	100.82	101.57	97.96	118.25	98.08	252.08	87.08	108.17	73.61	154.92	293.33
	February	239.25	144.67	89.83	74.40	132.92	192.04	150.88	166.83	132.55	125.03	197.54	179.76
	March	220.00	130.53	178.20	126.91	132.46	126.96	123.75	326.79	126.50	186.08	145.29	148.59
	April	420.75	147.19	120.54	127.25	155.83	109.08	141.17	546.33	333.30	178.75	142.54	136.68
	May	285.08	112.60	122.28	102.00	141.17	86.63	129.25	125.58	197.82	124.30	118.25	87.82
An. Rain. 1st Year (mm)		1855.39	1506.19	1361.90	1560.47	1597.01	1385.99	1190.16	1532.90	1089.18	1458.28	1309.72	1437.44
An. Rain. 2nd Year (mm)		2319.17	1032.97	1149.32	891.49	1168.29	987.71	1266.47	1635.33	1214.08	1080.84	1220.54	1217.33
Annual Rainfall (mm)		2087.28	1269.58	1255.61	1225.98	1382.65	1186.85	1228.31	1584.12	1151.63	1269.56	1265.13	1327.39
Thiessen Polygon (Km <sup>2</sup> )		18.18	33.81	34.35	23.99	6.36	34.72	11.82	11.45	11.27	28.72	17.45	25.45

## 2.4. Determination of Outflow

Data for pan evaporation ( $E_{pan}$ ) were obtained from two self-made replications of the US Weather Bureau “Class A” evaporation pans installed at the Rwasave fishpond and Gisunzu meteorological station. The reference potential evapotranspiration was established for data collected from June 2009 to May 2011. In order to obtain the real open water evaporation, the estimated  $E_{pan}$  was to be compensated for by a pan factor ( $f_{pan}$ ), where the pan factor depends on the type of pan used, pan environment and climate. For the “Class A” evaporation pan,

**Table 2.** Annual rainfall for 1973/1974 using Thiessen polygon method (mm).

Year	Month	Butare P(mm)	Kansi P(mm)	Nyakibanda P(mm)	Save P(mm)
1973	January	182.40	120.40	190.00	123.50
	February	110.50	176.30	67.80	100.10
	March	84.70	42.70	61.40	90.50
	April	141.00	223.20	189.10	177.60
	May	228.60	211.00	252.30	334.80
	June	54.50	20.70	0.00	19.50
	July	0.00	0.00	0.00	0.00
	August	3.60	17.70	0.00	16.20
	September	216.30	182.60	166.50	133.80
	October	117.50	107.60	146.50	66.90
	November	154.00	104.80	109.80	163.90
	December	72.20	69.50	104.80	36.50
1974	January	77.50	62.60	71.60	46.90
	February	58.90	49.90	100.80	85.00
	March	253.80	152.40	167.30	164.90
	April	172.00	172.50	183.40	211.40
	May	117.20	138.90	119.20	142.70
	June	71.90	16.80	30.10	51.50
	July	21.20	17.60	29.70	49.40
	August	12.90	0.00	0.00	4.20
	September	90.30	69.20	114.80	117.10
	October	55.00	47.80	45.00	17.70
	November	178.40	162.30	104.60	179.90
	December	86.40	125.90	91.80	77.00
Annual Rainfall 1973 (mm)		1365.30	1276.50	1288.20	1263.30
Annual Rainfall 1974 (mm)		1195.50	1015.90	1058.30	1147.70
Avg. Annual Rainfall (mm)		1280.40	1146.20	1173.25	1205.50
Thiessen Polygon (Km <sup>2</sup> )		60.30	105.02	58.77	33.47

$f_{pan}$  varies between 0.35 and 0.85 with an average value of 0.70 [20]. Based on a relative high humidity (>70%), moderate wind speed and green crops on a windward side distance in the order of 10 m [20], an empirically derived pan factor of 0.55 was used, **Table 3**.

The final, actual or adjusted crop factors ( $K_{adj}$ ) was calculated as shown in **Table 4**.

The dominant crops and their corresponding characteristics were described in [7] and the adjusted evaporation coefficient ( $K_{adj}$ ) is therefore equal to 0.74.

The Actual Evapotranspiration (Equation (2.5)) becomes:

$$AET = 854.88 \text{ mm} \times 0.74 = 632.61 \text{ mm}$$

**Table 3.** Reference evapotranspiration for 2009-2011 (mm).

Year	Month	Rwasave E (mm)	Gisunzu E (mm)	Avg. E (mm)	ET (mm)
		Column (1)	Column (2)	(3) = (1)*(2)/2	0.55*(3)
2009	June	140.00	125.60	132.80	66.40
	July	218.60	145.17	181.88	90.94
	August	145.50	113.61	129.56	64.78
	September	140.70	154.05	147.38	73.69
	October	151.30	141.80	146.55	73.28
	November	149.50	138.68	144.09	72.05
	December	118.70	133.75	126.23	63.11
2010	January	134.40	167.74	151.07	75.53
	February	100.90	159.07	129.98	64.99
	March	103.60	129.43	116.51	58.26
	April	124.90	133.87	129.38	64.69
	May	116.35	113.17	114.76	57.38
	June	135.58	130.50	133.04	66.52
	July	198.00	147.00	172.50	86.25
	August	194.00	170.00	182.00	91.00
	September	179.33	154.48	166.90	83.45
	October	170.13	153.93	162.03	81.01
	November	147.42	138.54	142.98	71.49
	December	151.25	129.63	140.44	70.22
2011	January	154.92	116.61	135.76	67.88
	February	183.54	127.03	155.29	77.64
	March	181.29	141.17	161.23	80.61
	April	98.04	157.75	127.89	63.95
	May	91.25	87.32	89.28	44.64
An. Et 1st Year (mm)		<b>1644.45</b>	<b>1655.94</b>	<b>1650.19</b>	<b>825.10</b>
An. Et 2nd Year (mm)		<b>1884.74</b>	<b>1653.94</b>	<b>1769.34</b>	<b>884.67</b>
Annual ET (mm)		<b>1764.60</b>	<b>1654.94</b>	<b>1709.77</b>	<b>854.88</b>

**Table 4.** Calculation of adjusted evaporation coefficient 2009-2011.

Type of Land Cover	Characteristics of Dominant Crops	Area (Km <sup>2</sup> )	Percentage (%)	Kc & Ke	Kadj.	Actual Kadj.
Column (1)	(2)	(3)	(4)	(5)	[Ks*(5)]	(4)*(5)
Agriculture	Banana trees	59.92	23.26	1.10	0.91	0.21
	Rain-fed herbaceous crops: coffee, maize ...	63.80	24.77	0.89	0.74	0.18
	Rain-fed shrub crops: sorghum, beans, cassava, potato	64.72	25.13	0.77	0.64	0.16
	Grasses and shrubs	0.14	0.05	0.70	0.58	0.00
Forest	Forests and scattered fields	14.59	5.66	0.70	0.58	0.03
Urban	Paved surfaces and roof tops	11.57	4.49	0.30	0.25	0.01
Wetland	Rice and papyrus	42.85	16.63	1.00	0.83	0.14
Entire Catchment		<b>257.57</b>	<b>100.00</b>			<b>0.74</b>



To estimate the final crop factors to be used for the calculation of the Actual Evapotranspiration (1973-1974), the land cover was assessed for the entire catchment during 1973-1974 (**Table 5**).

The adjusted evaporation coefficient ( $K_{adj}$ ) is therefore equal to 0.75.

The Actual Evapotranspiration ( $AET$ ) becomes:

$$AET = 768.55 \text{ mm} \times 0.75 = 576.41 \text{ mm}$$

#### a. Discharge in 2009-2010

To investigate surface water discharge in Migina catchment, five gauging stations were installed in 2009 **Table 6**.

Before the newly installed gauging stations in 2009, six old gauging stations were installed in 1993, but were destroyed during the Rwandan crisis of 1994. Therefore, Runoff coefficient was used to deduce the amount of discharge that flowed out of Migina catchment between 1973 and 1974. The rational method (Equation (2.6)) was used to know the runoff coefficient for the catchment.

**Table 5.** Calculation of adjusted evaporation coefficient for 1973-1974.

Type of Land Cover	Characteristics of Dominant Crops	Area (Km <sup>2</sup> )	Percentage (%)	Kc & Ke	K <sub>adj.</sub>	Actual K <sub>adj.</sub>
Column (1)	(2)	(3)	(4)	(5)	[Ks*(5)]	(4)*(5)
Agriculture	Banana trees	63.52	24.66	1.10	0.91	0.23
	Rain-fed herbaceous crops: coffee, maize	67.64	26.26	0.89	0.74	0.19
	Rain-fed shrub crops: sorghum, beans, cassava, potato	68.60	26.64	0.77	0.64	0.17
	Grasses and shrubs	0.14	0.06	0.70	0.58	0.00
Forest	Forests and scattered fields	19.78	7.68	0.70	0.58	0.04
Urban	Paved surfaces and roof tops	1.73	0.67	0.30	0.25	0.00
Wetland	Rice and papyrus	36.16	14.04	1.00	0.83	0.12
Entire Catchment		<b>257.57</b>	<b>100.00</b>			<b>0.75</b>

**Table 6.** Average discharge for 2009-2010.

Year	Month	Discharge (Q) [m <sup>3</sup> /s]
2009	August	0.31
	September	0.46
	October	0.62
	November	1.30
	December	0.56
2010	January	2.10
	February	2.05
	March	1.33
	April	1.42
	May	2.49
	June	0.54
	July	0.36
Avg.		<b>1.13</b>

For urban areas, the runoff coefficient is  $C_1 = 1$ .

The  $C_2$  (runoff coefficient) for the remaining land cover was therefore obtained as shown below;

$$1.13 \text{ m}^3/\text{s} = 0.00278 \times 0.15 \text{ mm/hr} \times (1 \times 1156.5198 \text{ ha} + C_2 \times 24600.65423 \text{ ha})$$

$$C_2 = \frac{(1.13 - 0.4823)}{10.26} = 0.06$$

The discharge

$$Q_p = 0.00278 \times 0.136 \text{ mm/hr} \times (1 \times 172.52 \text{ ha} + 0.06 \times 25584.65 \text{ ha}) = 0.65 \text{ m}^3/\text{s}$$

### 3. Results and Discussions

#### 3.1. Water Balance of Migina Catchment between 2009 and 2011

The water balance (Equation (2.1)) is given by:

$$\text{Inflows} - \text{Outflow} = \pm \text{Storage}$$

Inflows in this study are related to Precipitation ( $P$ ).

Outflow relates to Evapotranspiration ( $E$ ) + Outlet discharge ( $Q$ ).

Stock ( $S$ ) is the quantity of water stored in the catchment.

Therefore, the inflow in Migina catchment or the annual rainfall for June 2009-May 2011 is 1325.34 mm.

The outflow for Migina catchment is represented by evapotranspiration and outlet discharge where the annual reference evapotranspiration is 854.88 mm while the annual actual evapotranspiration for Migina catchment is equal to 597.80 mm. Also, the outlet discharge (the average) is equal to  $1.13 \text{ m}^3/\text{s}$ .

In estimating the water balance of Migina catchment for June 2009-May 2011, the following equation was used  $A(P - E) - Q = \Delta S$ .

From the results obtained:

$$A = 257.571740 \text{ km}^2 = 257571740 \text{ m}^2$$

$$P = 1325.34 \text{ mm} = 1.32534 \text{ m}$$

$$E = 597.8 \text{ mm} = 0.5978 \text{ m}$$

$$Q = 1.13 \text{ m}^3/\text{s} = 35635680 \text{ m}^3/\text{year}$$

The water balance of Migina catchment during June 2009 to May 2011 is calculated and found to be:

$$[(1.32534 - 0.597) \times 257571740] \text{ m}^3 - 35635680 \text{ m}^3 = 151964121 \text{ m}^3$$

And can be simplified as:

$$341370129.9 \text{ m}^3 - 153770328.8 \text{ m}^3 - 35635680 = 151964121 \text{ m}^3$$

The change in stock within Migina catchment during June 2009 to May 2011 is estimated to be 152 millions of  $\text{m}^3$ .

#### 3.2. Water Balance of Migina Catchment between 1973 and 1974

The inflow in Migina catchment (annual rainfall for 1973-1974) is 1191.5 mm. The outflow such as the annual reference evapotranspiration is equal to 768.55 mm. Therefore, the annual actual evapotranspiration for Migina catchment is equal to 576.41 mm. Meanwhile, the average discharge for 1973-1974 stands at  $0.65 \text{ m}^3/\text{s}$ .

The water balance of Migina catchment for 1973-1974 is obtained as follows;

From the results obtained:

$$A = 257.571740 \text{ km}^2 = 257571740 \text{ m}^2$$

$$P = 1191.5 \text{ mm} = 1.1915 \text{ m}$$

$$E = 576.41 = 0.576 \text{ m}$$

$$Q = 0.65 \text{ m}^3/\text{s} = 20498400 \text{ m}^3/\text{year}$$

The water balance of Migina catchment during 1973-1974 is calculated and found to be:

$$[(1.1915 - 0.576) \times 257571740] \text{ m}^3 - 20498400 \text{ m}^3 = 138037006 \text{ m}^3$$

And can be simplified as:

$$306896728.2 \text{ m}^3 - 148361322.2 \text{ m}^3 - 20498400 \text{ m}^3 = 138037006 \text{ m}^3$$

The change in stock within Migina catchment during 1973-1973 is estimated to be 138 million of  $\text{m}^3$ .

### 3.3. Impact of Urbanization on the Water Balance of Migina Catchment

#### a) During 2009 and 2011

Evapotranspiration represents 45% of the total inflow, while runoff represents 10.5% of the total inflow and change in storage represents 44.5% of the total inflow.

#### b) During 1973-1974

Evapotranspiration in 1973/1974 is 48% of the total inflow, runoff is 7% of the total inflow and change in storage is 45% of the total inflow. These findings tally and are in close range as that given in [22]. These studies revealed that, during the last 30 years the runoff has increased by 3.5%. This may be attributed to the infiltration of water into the soil [23] [24].

This may also lead to increase pollution due to surface water run-off which may also spread water borne diseases as a result of little or no water self-purification. This increase in runoff also reduces water storage hence leads to water scarcity within the surrounding area [25]-[27].

A decrease of 3% in evapotranspiration maybe due to the conversion of green spaces for urbanization purposes especially housing. This reduces surface albedo hence a drop in the contribution of evapotranspiration to rainfall. A progressive reduction in rainfall is gradually created with consequences for agriculture due to recurrent droughts [24] [26]. The change in storage has slightly decreased by 0.5% because increase runoff causes out-flow of catchment water that could have been stored as groundwater [26] [27].

These changes in the hydrological cycle could be attributed to the new residential plots of Butare town whose population increased from a mere 8400 in 1970 to over 100,000 inhabitants in 2009. This growth in population has been detrimental to the natural landscape cover hence fluctuations in the hydrological cycle.

## 4. Conclusions

The water balance of Migina catchment was estimated over a 30-year interval with more rainfall experienced during 1973/1974 in Migina catchment than in 2009-2011. These observations were also true for surface water runoff which was considerably lower in 1973/1974 and higher in 2009-2011. These changes are all due to an increase in demand for basic and social amenities as a result of an increase in population of Butare town.

The consequences of urbanization on the hydrological cycle could be curtailed by structural adjustment programmes including urban greening and provision of detention of runoff water in urban drainage systems.

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