

# Surface Rupture and Hazard of Wenchuan Ms 8.0 Earthquake, Sichuan, China

Yong Li<sup>1</sup>, Runqiu Huang<sup>2</sup>, Liang Yan<sup>1</sup>, Alexander L. Densmore<sup>3</sup>, Rongjun Zhou<sup>4</sup>

<sup>1</sup>State Key Laboratory of Oil and Gas Reservoir Geology and Exploitation,  
Chengdu University of Technology, Chengdu, China

<sup>2</sup>State Key Laboratory of Geo-hazard Prevention & Geo-environment Protection,  
Chengdu University of Technology, Chengdu, China

<sup>3</sup>Institute of Hazard and Risk Research and Department of Geography, Durham University, Durham, UK

<sup>4</sup>Institute of Earthquake Engineering, Seismological Bureau of Sichuan Province, Chengdu, China

E-mail: yanliang1003@163.com

Received March 25, 2010; revised April 19, 2010; accepted May 13, 2010

## Abstract

Longmen Shan is located the special joint between the Qinghai-Tibetan Plateau in the west and the Yangtze craton in the east. Consisting of a series of parallel imbricated thrust, it develops, from the west to the east, the Maoxian-Wenchuan, Yingxiu-Beichuan and Pengxian-Guanxian faults. Wenchuan Ms 8.0 earthquake is a thrust with strike-slip type, and surface ruptures are located in Yingxiu-Beichuan fault zone and Pengxian-Guanxian fault zone. Based on the geological background, tectonic setting, the active tectonics of Longmen Shan and surface ruptures of the Wenchuan earthquake, a dynamical model to illustrate possible links between surface processes and upward extrusion of lower crustal flow channel at the eastern margin of the Tibetan plateau have been studied, and the results is the material in lower crust in the Longmen Shan moving as nearly-vertical extrusion and uplift, resulting in the surface rate of tectonic movement differing according to depth rate as well as the occurrence of large shallow Wenchuan earthquake.

**Keywords:** Wenchuan Earthquake, Longmen Shan, Surface Rupture, Hazard, Sichuan, China

## 1. Introduction

The Ms 8.0 Wenchuan Earthquake in May 12th, 2008 is one of the most disastrous earthquakes since the foundation of P. R. China, which destroyed not only the epicenter of Sichuan province but also several closed provinces. It was felt in most regions of China, as well as nations outside China. This tragic event provides the opportunity to advance the subject of seismic sciences.

Based on our accumulated activity on active tectonics in Longmen Shan area, we have undertaken several new field surveys, including international collaborative efforts after the earthquake. This paper compiles 70 sets of data from accumulated past and new surveys, detailing surface rupture and seismic disasters since our work began in the 1990s [1-15].

## 2. Geological Structure in Longmen Shan Seismic Belt

Longmen Shan is located the special joint between the Qinghai-Tibetan Plateau in the west and the Yangtze craton in the east, which is a foreland basin and orogeny system with a complex structure and unique structural history. In the depth, it is located in structural transform belt of southwest China and the south part of Helan-shan-Longmen Shan. To the northwest, there is thickened crust and thickened mantle area; and to the southeast, there is thin crust and thin mantle area in Sichuan basin. From the Longmen Shan foreland basin to the plateau, the crust is thickening roughly and forms a belt with western-cline belt; the central line of the belt is located the position of Longmen Shan thrust under the surface. Comparing with its surface position, the line moves some distance to the west, which suggests that Longmen Shan thrust dips to the west and does not have a mountain root. It also suggests that Longmen Shan is an Intracontinental

---

This research was supported by National Natural Science Foundation of China grant 40841010 and CGS grant 1212010918010.

mountain chain and an independent tectonic unloading system.

The isostatic gravity anomaly in Longmen Shan and joint regions show that the crust in this area is still unstable. The figures in Longmen Shan are positive and the figures in Longquan Shan and the area to the east are negative while the figures in the area between them are around zero. Since the Cenozoic, at least 5-10 km strata have eroded in uplift rate of 0.6 mm/a. The survey data in recent years shows that there is an uplift rate in 0.3-0.4 mm/a in Jiuding Shan [1,16].

According to the regional geological setting, this area, from northwest to southeast, consists of Songpan-Ganzi Orogeny (main part), Longmen Shan thrust belt, foreland basin and forebulge; all of these four parts form a complete tectonic system. We would introduce the east margin of Qinghai-Tibetan Plateau as three geological units: Songpan-Ganzi orogenic belt (main part), Longmen Shan thrust belt and foreland basin [1-15].

### 3. Surface Rupture of the Wenchuan Earthquake

After the earthquake, a number of research institutions have published basic parameters and focal mechanisms. However, those data are different depending on different sources and methods. The USGS recorded that the  $M_w$  is 7.9, the epicenter is located in Yingxiu, Wenchuan (31.099, 103.279), the focal depth is 19 km, the rupture

length is 300 km and the duration is 120 seconds, trending  $229^\circ$  and fracture plane dipping  $33^\circ$  with strike-slip. According to the basic parameters and the focal mechanism solution, the Wenchuan earthquake had the following two characteristics: 1) it was a shallow hypocenter earthquake and took place in the brittle-ductility transform belt in 12 km-19 km depth with enormous destructivity; 2) it belongs to thrusting and dextral strike-slip one and from southwest to northeast, it propagated in one direction. The release process of strain was relatively slow and may have led to a greater intensity of aftershocks and longer duration. The intensity of the earthquake had an oval-shaped distribution and its long axis was in a northeast direction. In this direction, the casualties and property loss was much bigger than other directions.

According to the surface rupture survey (**Figure 1**) after the earthquake, the Wenchuan earthquake faults have characteristics of fragile, post-earthquake surface linear image is apparent, cutting a wide range of terrain units. They are located in Yingxiu-Beichuan fault zone (central zone), Pengxian-Guanxian fault zone (front fault) which only a small amount of surface rupture are found and Maoxian-Wenchuan fault zone (back fault) which surface rupture have not yet been found. Those surface rupture extend in northeast direction (in  $NE\ 30^\circ-70^\circ$ , most in  $NE\ 50^\circ-60^\circ$ ) and trend to northwest ( $30^\circ-80^\circ$ ). They distribute along the striking in a few meters to 200 meters range and were also distributed along the Yingxiu-Beichuan fault and Pengxian-Guanxian fault.

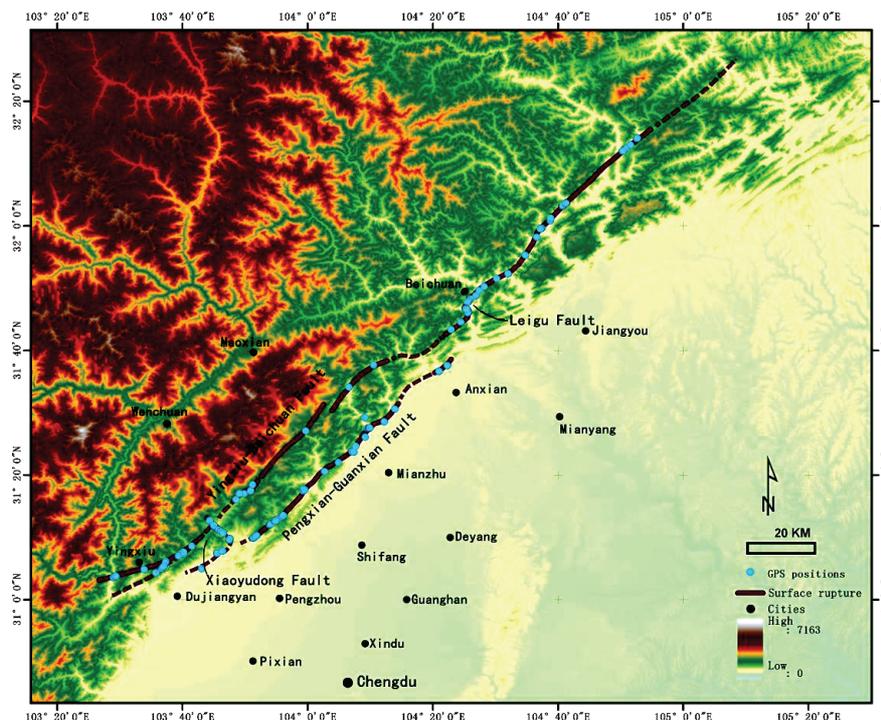


Figure 1. Distribution of surface rupture of Wenchuan Earthquake in Longmen Shan.

### 3.1. Surface Rupture of Yingxiu-Beichuan Fault

The Yingxiu-Beichuan fault (Figures 1-3) is the Wenchuan earthquake's main earthquake fault. The linear characteristics of the earthquake surface rupture belt are apparent. It starts from Yingxiu, Wenchuan in the southwest and then extends to Hongkou, Longmen Shan town (Baishuihe), Donglinsi, Hongbai, Qingping, Chaping, Leigu, Beichuan, Chenjiaba, Guixi, and then ends in Shikanzi, Pingwu, a distance of about 220 km. It was distributed along the strike of Yingxiu-Beichuan fault and disappears in the southwest 10 km of Yingxiu, belonging to the single-side and multi-point type rupture with thrust characteristics. Its plane had a steep dip angle and cut a wide range of terrain units, including mountain bedrocks, river terraces, alluvial fans, highways and bridges with bridge collapse or road displacement. Its vertical offsets were in the range of 1.60-6.20 m and its horizontal offsets were in the range 0.20-6.50, striking NE 30°-50° and trending northwest. The average vertical offset was 3.4 m and horizontal 2.9 m, the biggest offset being  $10.3 \pm 0.2$  m in Maoba, Beichuan (vertical fault) and  $6.8 \pm 0.2$  m in Leigu (horizontal offset). The ratio between thrust and dextral strike-slip components are 1:1, indicating the existence of thrust movement and dextral strike-slip movement, the amount of thrust movement being equal to the dextral strike-slip movement.

Basing this and according to the current surface data obtained from the survey, we calculated vertical and horizontal offsets in four parts of Yingxiu-Beichuan surface rupture. The preliminary analysis results are: 1) the epicenter located in the vicinity of Yingxiu (Figure 1) is

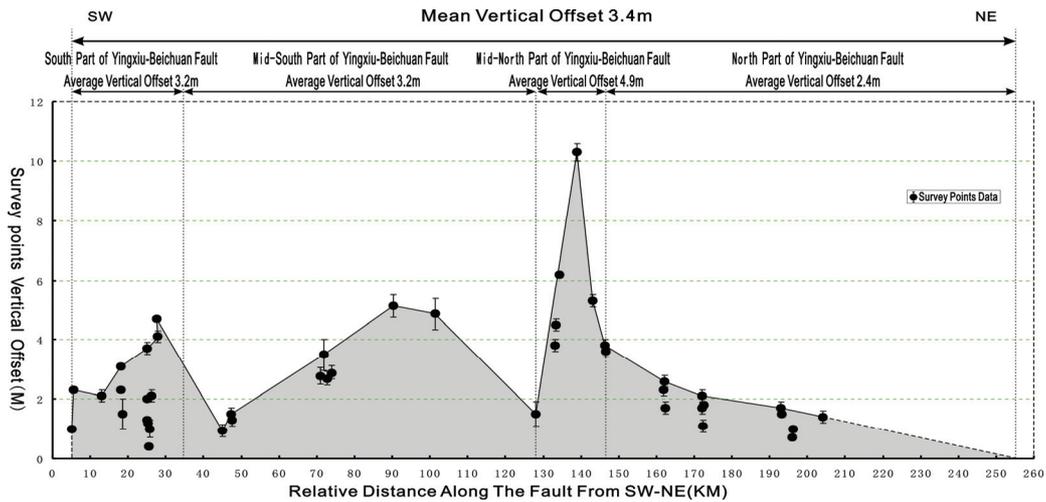


Figure 2. The vertical displacement along the Yingxiu-Beichuan Fault in Wenchuan Earthquake.

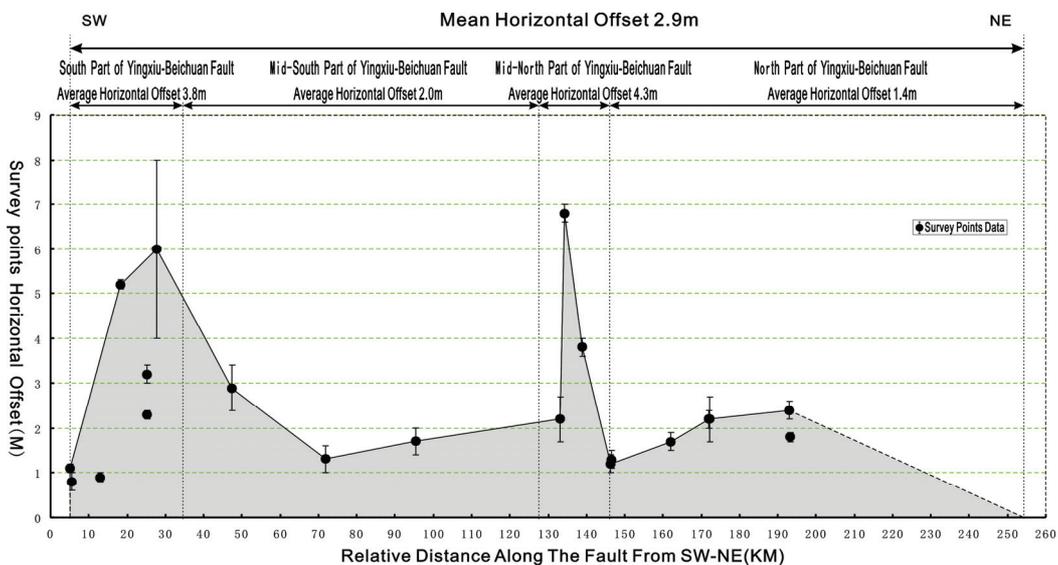


Figure 3. The horizontal displacement along the Yingxiu-Beichuan fault in Wenchuan Earthquake.

not the position where the biggest offset is located; the biggest vertical and horizontal offsets are  $10.3 \pm 0.2$  m (vertical offset) and  $6.8 \pm 0.2$  m (horizontal offset) in Beichuan. According to Wenchuan earthquake focal mechanisms solution, the greatest offset of the main shock is 9 m-10 m. As a result, the surface rupture is smaller than the biggest displacement under ground; the biggest offset are 60%-70% of the biggest offset under ground; 2) by surface offsets, the Yingxiu-Beichuan fault can be divided into two high-value and two low-value zones. The two high-value zones are located in the southern part of Hongkou, Yingxiu and the northern part of the Leigu, Beichuan-Dengjiaba area and the two low-value zones are located in mid-south part of the Baishuihe-Chaping area and Huangjiaba-Shikanzi, Pingwu area. These four sections are divided by Xiaoyudong fault, Leigu fault and Dengjiaba fault; the two high-value areas being caused by the Xiaoyudong and Leigu faults; 3) Chen [17] use the global seismic network data to confirm two largest static sliding displacement areas, which are located at area between epicenter and northeast 100 km to the epicenter and 150 km northeast of the epicenter; the two largest displacement of the slip surface locate in the two high-value areas, indicating that the largest displacement in the surface are the response to the displacement under ground; 4) compared the surface rupture of earthquake with before, the places in Yingxiu, Leigu, Baishuihe and Gaoyuan with historical earthquake rupture are also the places with new rupture[3-8,12,15]. In the places with historical earthquake since Quaternary, there will be new strong earthquake in the future.

### 3.2. Surface Rupture of Pengxian-Guanxian Fault

The Pengxian-Guanxian fault (Figures 1 and 4) experienced surface fractures in this earthquake. The surface

rupture started from Xiang'e, Dujiangyan and extended to Cifeng, Pengzhou, Bailu, Jinhua, Hanwang and Sangzao over a total distance of about 40-50 km, characterized by dextral strike-slips shortening with steep dipping. The hanging wall was in the northwest and its foot wall in the southeast. The vertical offset was 0.39-2.70 m and the horizontal offset was 0.20-0.70 m (the average vertical offset was 1.6 meters with an average horizontal offset of 0.6 meter), indicating the existence of thrust displacement and dextral strike-slip displacement and thrust displacement is greater than the dextral strike-slip displacement, showing thrust and shortening characteristics.

Compared with the surface rupture of the Yingxiu-Beichuan, the fault rupture is far less than the Beichuan-Yingxiu fault rupture and it has a relatively short length of the rupture and vertical offset; and on the other hand, the horizontal offset displacement is relatively small, the thrust displacement is main.

### 3.3. Surface Rupture of Xiaoyudong Fault

The fault (Figures 1, 5 and 6) is located between Yingxiu-Beichuan fault and Pengxian-Guanxian fault and is a new fault being discovered after the earthquake, showing that the fault is a transformed fault between the Yingxiu-Beichuan fault and the Pengxian-Guanxian fault. The fault, striking south-north, extends about 15 km through Xiaoyudong Bridge; the stretch of the surface rupture is stable and we have confirmed more than 10 rupture points. The southwest block is a hanging wall and the northeast block is foot wall with an average vertical offset of 1.0 meter and an average horizontal offset of 2.3 meters (vertical offset/horizontal offset ratio is 1:1), showing the vertical offset is equal to the horizontal offset and, from south to north, the thrust displacement is increasing.

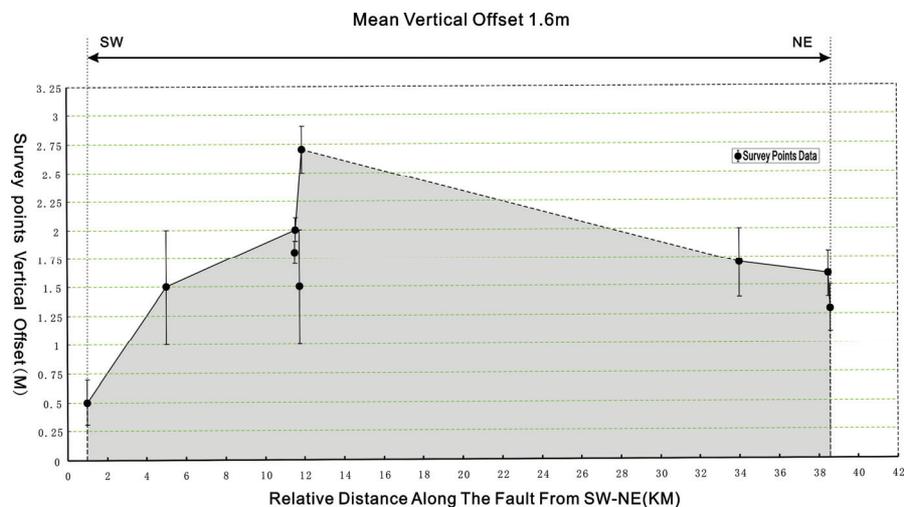


Figure 4. The vertical displacement along the Pengxian-Guanxian fault in Wenchuan Earthquake.

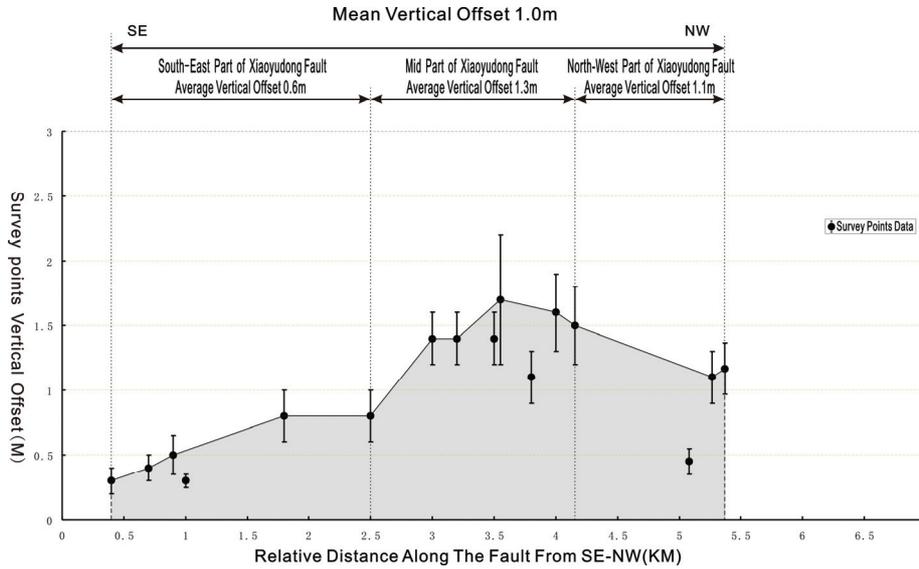


Figure 5. The vertical displacement along the Xiaoyudong fault in Wenchuan Earthquake.

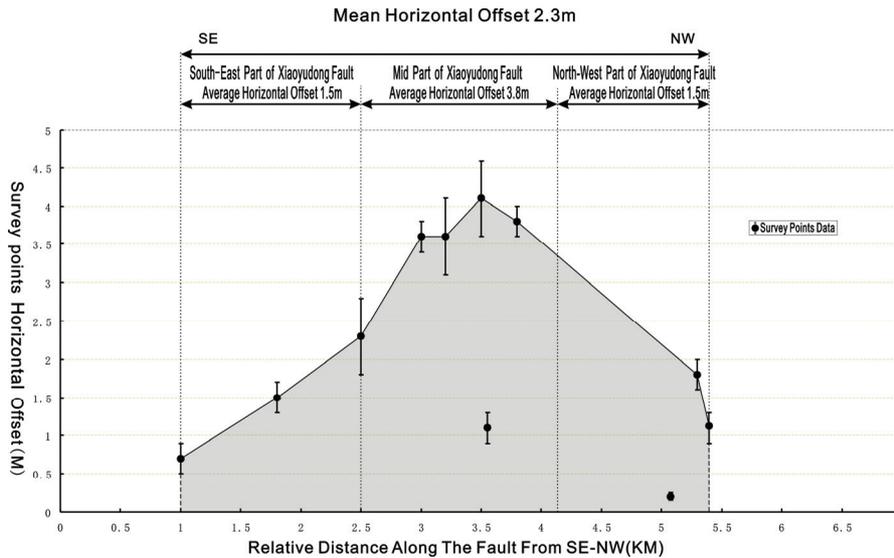


Figure 6. The horizontal displacement along the Xiaoyudong fault in Wenchuan Earthquake.

### 3.4. Surface Rupture of Leigu Fault

This fault (Figures 1 and 7) belongs to Yingxiu-Beichuan fault belt and between two parallel northeast striking faults. The fault's surface rupture is very complex and it is also a new fault discovered after the earthquake, showing that it is a transform fault between those two parallel Yingxiu-Beichuan faults, which has been named Leigu fault by us.

The fault, striking south-north, extends 3-5 km from south to north. The surface rupture pattern is stable and we have confirmed that more than 10 rupture points. Its southwest block is hanging wall and the northeast block is foot wall. The largest vertical offset is  $2.2 \pm 0.5$  m and

the largest horizontal offset is  $1.9 \pm 0.1$  m (the average vertical offset is 1.8 m and the average horizontal offset is 1.4 m).

In the surface rupture, there are obvious changes from the north to the south, changing from striking south-north in the north to striking NWW in the south, which shows vertical and sinistral horizontal offset in the north (a channel has been offset vertically 1.9 m and horizontally 1.3 m); and, in the south, the rupture zone shows vertical and dextral offset (a house foundation has been offset vertically 1.5 m and 0.5 m and dextral 1.3 m and 0.3 m).

The vertical offset with horizontal offset ratio is 1:1, showing the thrust displacement is equal to the strike-slip displacement and, from south to north, the thrust changes

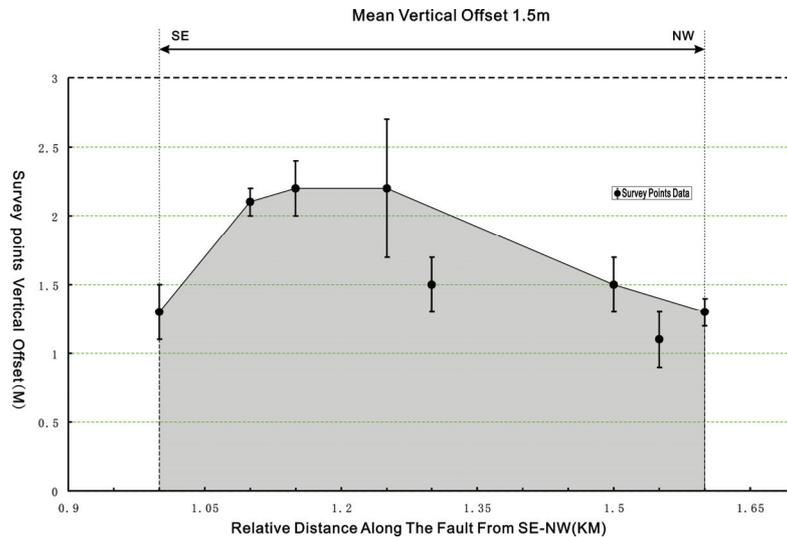


Figure 7. The vertical displacement along the Leigu fault in Wenchuan Earthquake.

bigger. It was noticed that in the north, the fault is sinistral strike-slip and, in the south, it appears to be a dextral strike-slip.

### 3.5. The landform of Surface Rupture

According to our field observation of the surface rupture, it had various forms, mainly showing as ridge offsets, alluvial fan offsets, terrace offset, slopes offsets, channel offsets, path offsets, highway offsets, cement road overlays, structural fractures, sag ponds, ground cracks, earthquake domes, pressure ridge and other types, of which road scarps are easy to identify.

In Yingxiu region, the surface rupture are showed in structural cracks and road deformation (asymmetrical fold ridge) and graben; in Shengxigou and Gaoyuan, the surface are showed road scarps, road deformation and pressure ridge; in Bajiaomiao, the surface rupture are showed in fault scarps. The width of surface rupture is different along the fault, but on the whole, it is generally less than 20 meters. Strong deformation is distributed in the hanging block of the thrust fault and the foot block has no obvious deformation, or only has deformation in the vicinity of the fault.

### 3.6. Dip of Surface Rupture

The Yingxiu-Beichuan fault's surface ruptures are generally steep scrapes; the northwest block is hanging wall and southeast block is foot wall, indicating that the fracture plane should trend northwest. But the plane is difficult to be discovered; in Bajiaomiao, Hongkou, the plane was directly exposed at the surface. It tends to northwest with dip angle  $80^{\circ}$ - $86^{\circ}$  and its hanging wall is composed of upper Triassic Xujiahe coal-bearing strata and its foot

block is loose gravels. The scratches can be distinct in the plane, indicating high-angle thrust characteristics. As a result, in the Yingxiu-Beichuan fault, the northwest trending fault in high-angle should be representative.

However, in Gaoyuan, Hongkou, Yingxiu-Beichuan fault's rupture surface trend to southeast. The plane here is also showed as a very steep scrapes, but the northwest block is foot side and southeast is hanging side, which indicates that the plane should trend to northwest. As a result, we think that Yingxiu-Beichuan fault's surface rupture should be northwest trending high-angle surface, but in some areas, the tendency could be reversed from northwest to southeast.

The Wenchuan earthquake's focal mechanisms solution reveals that the dip angle of plane is only  $33^{\circ}$ - $39^{\circ}$  [17]. It is obvious that the surface crack plane is different from one calculated by focal mechanisms solution: the dip angle is gentle when deeper. This characteristic is also similar to an imbricate thrust fault's structural features.

### 3.7. Surface Rupture and Shortening Rate

Although in the Yingxiu-Beichuan fault there are different structural shortening forms, including road offset and road overlay, the proper measurement position of surface deformation is less. In a cement road in Gaoyuan, we measured the pressure ridge and the thrust overlay, showing a clear structural shortening phenomenon. In the cement road, we measured surface deformation in detail, indicating that the structural shortening rate of this section is 7.61 percent and that the biggest structural shortening rate of the pressure ridge is 28.6 percent.

The Longmen Shan is the most typical klippe zone in China; it has a structural shortening rate of 42-43% [1,3-8]. The result is 7.61%-28.6% in Gaoyuan, showing the clear structural shortening rate in Wenchuan Earthquake.

### 3.8. Rupture Process of Wenchuan Earthquake

Although the surface rupture in Bajiaomiao is also steep scarp as the main form; its hanging wall is Upper Triassic Xujiahe coal-bearing strata and its foot wall is loose gravel. But the scrape marks have preserved the evidence for the process analysis of Wenchuan earthquake.

The scrapes plane strikes to N60°E, trending to N300°W with a dip angle of 85°. We discovered two scratches in the plane. The first one is nearly vertical, mainly distributing in upper and lower parts of the plane, while the second one is nearly horizontal, mainly locating in the lower part of the plane.

In the plane of the scrapes, we can observe three kinds of cross cutting relationships between nearly vertical and horizontal scratches: a nearly vertical one cutting the nearly horizontal one, a nearly horizontal one cutting the nearly vertical one and a nearly vertical one in the top gradually changing to the nearly horizontal one in the bottom.

In view of the Upper Triassic Xujiahe coal-bearing strata in the hanging block and the loose gravel in the foot block, the scratch preserved in the hanging wall bedrock should be the evidence of foot wall loose gravel in the movement process. As a result, the upper marks preserved in the hanging wall should be the early evidence of the rupture process of the earthquake and the lower marks preserved in the hanging wall should be the late evidence of the rupture process.

For this reason, we speculate that the nearly vertical scratches in the top recorded the thrust movement early, while the scratches in the bottom recorded the strike-slip movement lately. According to these three different kinds of cross cutting relationship of scratches, we think that the late movement is oblique.

According to the focal mechanisms solution, Chen [17] thought that the Wenchuan earthquake can be divided into 7 stages, and that, in the beginning stages, the movement showed thrust and then gradually shifted to strike-slip movement lately. Yuji Yagi [18] has a similar viewpoint on the movement process of the devastating earthquake which includes two phases: the first phase (0-50 seconds) with 6.4 m thrust and the second phase (60-120 seconds) with 4.6 m strike-slip.

As a result, the scratches of the surface rupture reveal the process confirmed by the focal mechanism solution, which is that the thrust movement occurred early and the strike-slip movement occurred lately.

### 4. Discussion on Kinematic of Active Tectonics in Longmen Shan and Wenchuan Earthquake

Indo-Asia collision is the most important events in Cenozoic, which results in the uplift, deformation and

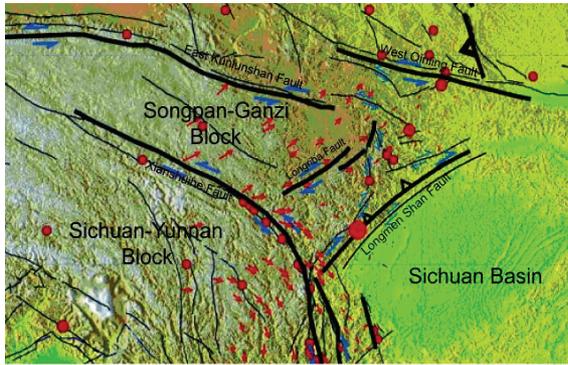
thickening of Qinghai-Tibetan Plateau. The event and its impact to Cenozoic geological structure has been noticed and discussed by geoscientists. Two well-known hypotheses have been presented; one is a crustal thickening model [19] and the other one is a lateral extrusion mode [20]. The former emphasizes on north-south crustal shortening and thickening, and the later emphasizes the east-extrusion along main faults. In the eastern margin of the Qinghai-Tibetan Plateau, there are two corresponding patterns which are: Avouac and Tapponnier [20] eastward thrusting mode, and England *et al.*'s dextral strike-slip mode.

However, the active tectonics and Wenchuan earthquake show the Longmen Shan fault is characterized by thrust and dextral strike-slip movement, which does not coincide with England and Molnar's large scale dextral shear movement in the eastern margin of Qinghai-Tibetan Plateau, and which also does not coincide with Avouac and Tapponnier [20] eastward thrusting mode in the eastern margin of the Qinghai-Tibetan Plateau. The performance of Longmen Shan fault zone has its uniqueness, which can not be explained in one single mode.

According to the direction of surface movement showed by GPS measurements [21,22], the northeast corner of Qinghai-Tibetan Plateau consisting of Longmen Shan fault and Kunlun Shan fault has the characteristics of northeastward (Qingling direction) extruding out, which means that when the Songpan-Ganzi block in eastern Qinghai-Tibet moved to the east, because of the blocking of Yangtze platform, it was forced to move to Qinling Mountains in the northeast and resulted in the characters with thrust and dextral strike-slip movement mode in Longmen Shan. The West Qinling fault is thrust with sinistral strike-slip movement and the blocks between Longriba fault and east Kunlun fault is also northeastern extruding out (**Figure 8**).

### 5. Discussing of Dynamics Mechanism of Active Tectonics in Longmen Shan and Wenchuan Earthquake

According to GPS measurements, India plate is moving to the north in 50 mm a<sup>-1</sup>, and different blocks in the eastern margin of Qinghai-Tibetan Plateau have significantly different movement rates. The Songpan-Ganzi block in the south has a faster horizontal movement rate than the Songpan-Ganzi block in the north. Active tectonics research [3-8,12,15] shows that the horizontal movement rate in Longmen Shan is small (1-3 mm a<sup>-1</sup>), so is its uplift rate (0.35-0.40 mm a<sup>-1</sup>) [1]. Its movement mode is characterized by thrust with dextral strike-slip. But Longmen Shan is the steepest mountain around Qinghai-Tibetan Plateau, in the range of more than 30 kilometers, it uplifts from 700 meters above sea level to more than 5000 meters. The small surface slipping rate



**Figure 8. Surface movement, epicenter and tectonic framework in Eastern Qinghai-Tibetan plateau.**

does not only coincide with the high and steep mountain, but also does not coincide with the fact of the strong earthquake of Ms 8.0. As a result, in the Songpan-Ganzi block and Longmen Shan in the northeast margin of Qinghai-Tibetan Plateau, the rate of surface movement does not coincide with the deep tectonic movement rate. The deep structural process must be thought about in the seismic hazard assessment.

The Wenchuan earthquake’s focal mechanism solution shows that Wenchuan Earthquake’s focal depth is 12-19 km, belonging to shallow earthquake. In addition, we count the historical earthquake in Longmen Shan as the longitude [3-8,9-11]. The results show that Longmen Shan tectonic earthquake in lower magnitude was in depth of 5-15 km in advantage and the strong earthquake was in depth of 15-20 km. Obviously, it is clear that the depth of

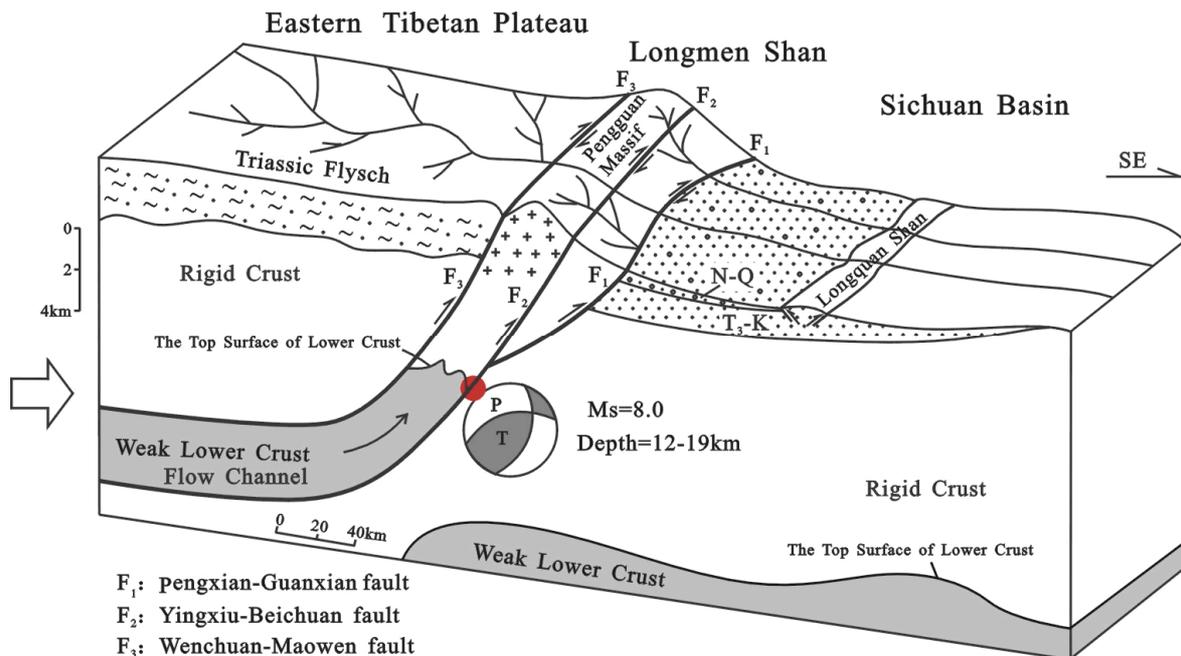
Wenchuan Earthquake coincide with the depth of historical strong earthquake and all of them are shallow earthquakes, which coincide with the depth of 20 km in low-velocity layer and also coincide with the uplift depth of the top of lower crust [9-11].

As a result, in the process of Qinghai-Tibetan Plateau’s movement to the east, blocked by the basement of the Sichuan Basin, the lithosphere is not synchronized as a whole flow to the east and the flow velocity is different in different layers. This results in lower crust material in the Longmen Shan moving as nearly-vertical extrusion and uplift [23], resulting in the surface rate of tectonic movement differing according to depth rate as well as the occurrence of large shallow Wenchuan earthquake (Figure 9).

### 6. Geological Hazard Caused by Wenchuan Earthquake

The Wenchuan Earthquake is the most destructive in intensity and the vastest of affected area in China since 1949. The area of the heavily destroyed region is more than 100,000 km<sup>2</sup>. It has the most serious casualties after the Cicheng Earthquake (Inner Mongolia, 1290), the Huaxian Earthquake (Shanxi, 1556), the Haiyuan Earthquake (Ningxia, 1920) and the Tangshan Earthquakes (Hebei, 1976).

The hazards caused by the Wenchuan Ms 8.0 Earthquake is known by next main characteristics: 1) seismic waves have propagated with strong ground motion to damage construction; 2) surface rupture zone directly



**Figure 9. A dynamical model to illustrate possible links between surface processes and upward extrusion of lower crustal flow channel at the eastern margin of the Tibetan plateau.**

destroyed and tear construction along Yingxiu- Beichuan fault for 220 km and Pengxian-Guanxian fault for 40-50 km; 3) collapse, landslide and other geological hazards caused by strong vibration of ground motion destroyed buildings. Because the earthquake occurred in the valleys of the Longmen Shan, especially in Yingxiu-Beichuan fault, there have been large-scale collapse, landslides and other geological hazards in Longmen Shan region. Those hazards have buried or destroyed a large amount of buildings, for example, half of the buildings in Chenjiaba, have directly been destroyed by landslides and a large number of barrier lakes have flood buildings and have given potential secondary flood threat to the basin.

According to the earthquake data collected by survey teams of Gansu, Shanxi and Chongqing seismological bureaus, InSAR imaging and seismography records, we studied some important regions with high seismic intensity and have drawn an isoline map of Wenchuan Earthquake and distinct the intensities in this region (**Figure 10**). District XI is located in Yingxiu-Hongkou, and Leigu-Beichuan, which are distribute along Yingxiu-Beichuan fault zone's hanging wall with a total area of about 680 km<sup>2</sup>. Construction in this region almost com-

pletely collapsed, especially in Yingxiu town and Beichuan town. Large-scale surface rupture can be followed. A number of large-scale collapses and landslides blocked rivers to form barrier lakes, such as the Qingping and Tangjiashan barrier lakes. District X starts from southwest Yingxiu and end in Shikanzi, Nanba in the northeast, including Yingxiu-Beichuan surface rupture in a narrow belt with the long axis N50°E and an area of about 2520 km<sup>2</sup>. In this region, most construction collapsed and landslides are common.

To summarize, District VI and above this of the Wenchuan Earthquake affected the area of about 333000 km<sup>2</sup>. The isoline map (**Figure 10**) has the following characteristics: 1) high-seismic intensity lines, especially in district IX, distributed along the Longmen Shan in N40-50°E direction and the long/short axis ratio are 8:1-10:1. District XI degrees are distributed in three isolated areas with the typical characteristics of multi-point instant cracks. 2) According to the results of seismic inversion, the earthquake rupture tear from the vicinity of Yingxiu to the northeast with characteristics of a one-direction rupture process. Isoseismal lines also show the direction of disappearing rapidly to the southwest and slowly to the northeast with the characteristics of the

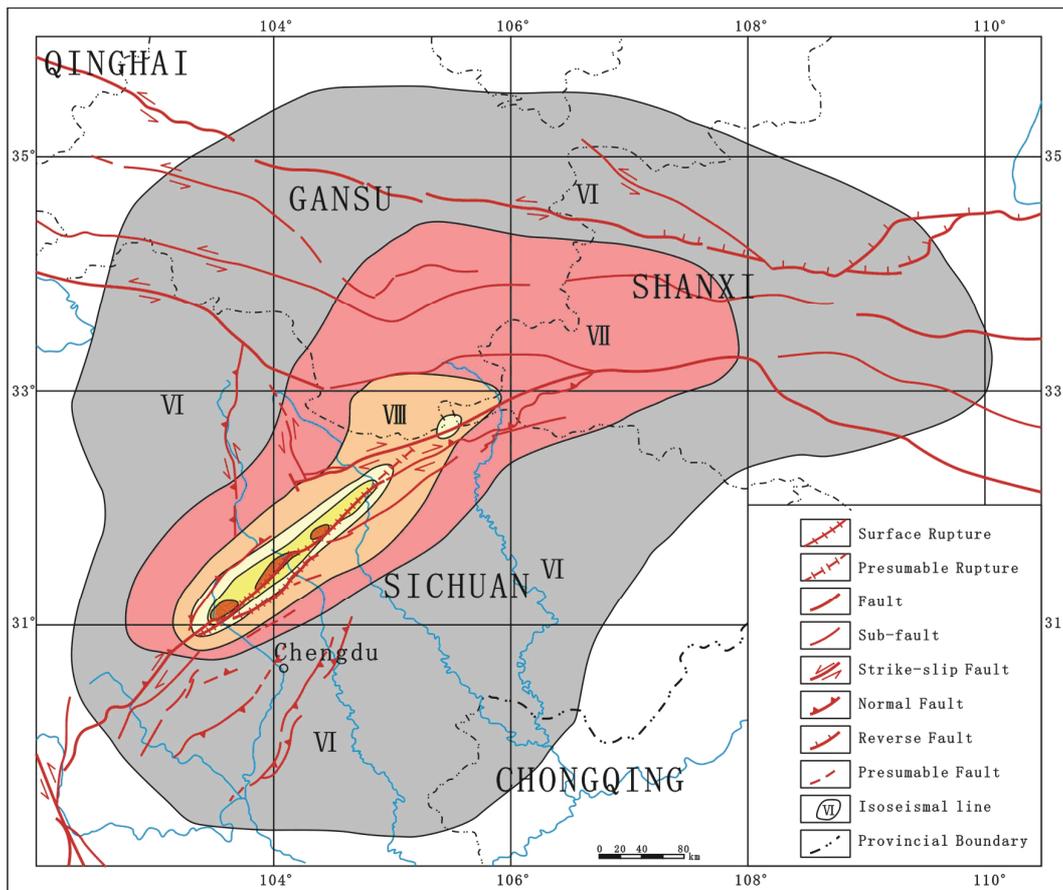


Figure 10. Isoseismal line of Wenchuan M8.0 earthquake on 12 May, 2008.

mode. 3) District VI disappear to Hongyuan and Ruergai slower than to the Basin, which may be related to that the Sichuan Basin with relative rigidity Yangtze plate is not easy for seismic wave absorbing as well as the soft soil in Hongyuan and Ruergai enlarge the damage. 4) District VI, VII and VIII in the area of north Sichuan, Gansu and Shanxi are gradually turning to east-west direction and this may be the result of controlling of the regional faults.

## 7. Preliminary Conclusions

In this paper, based on the Longmen Shan seismic zone's geological background, tectonic setting, stratigraphy and lithology, we have summarized the Longmen Shan surface rupture and progress and explore the tectonic kinematics and dynamics of Wenchuan Earthquake. According to the geological disasters caused by Wenchuan Earthquake, we have made several suggestions for reconstruction in the affected areas.

According to historical records and active tectonics, we believe that the Longmen Shan fault zone is a dangerous earthquake zone and the three main faults have the ability to induce earthquakes > Ms 7. The Yingxiu-Beichuan fault is the most important earthquake-inducing fault; the interval should be at least about 1000 a. It belongs to the low-frequency seismic activity zone, but it has a potential risk to induce a strong earthquake with characteristics of thrust and dextral strike-slip movements.

By comparing the initial analysis of Wenchuan Earthquake and historical earthquakes records, we believe that, the May 12, 2008 Wenchuan Earthquake belongs to thrust and strike-slip type earthquake. According to north-south striking fault (Xiaoyudong, Leigu and Dengjiaba faults) and surface offsets, Yingxiu-Beichuan fault surface rupture zone is divided into two zones of high-value and low-value. According to scratches on the scraps in Hongkou, the earthquake rupture process is divided into two phases: early thrust and lately-inclined strike-slip movement. As a result, in the earthquake ruptured zone, there are two components of thrust and strike-slip, thrust displacement is slightly larger than dextral strike-slip one. This does not coincide with the crustal thickening mode and the lateral extrusion mode, which means that Yingxiu-Beichuan fault is special one and can not be explained by one single model. In view of the surface moving rate of Longmen Shan does not coincide with its moving rate in the depth, we discussed the dynamic geological model between surface process and lower crust flow and think that the vertical pressure and vertical movement of lower crustal materials in Longmen Shan results in the eastern thrust movement, uplift of

Longmen Shan tectonic belt and Wenchuan Earthquake as well.

## 8. Acknowledgements

We are very grateful to the people who provided us supports and help. Finally, we thank the people of the Longmen Shan region for their unending curiosity, hospitality, and generosity.

## 9. References

- [1] Y. Li, Y. F. Zeng and H. S. Yi, "Sedimentary Record and Tectonic Evolution of Longmen Shan Foreland Basin," Chengdu University of Science and Technology Press[A], Chengdu, 1995 (in Chinese).
- [2] Y. Li, P. A. Allen, A. L. Densmore and X. Qiang, "Evolution of the Longmen Shan Foreland Basin (western Sichuan, China) during the Late Triassic Indosinian Orogeny," *Basin Research*, Vol. 15, 2003, pp. 117-138.
- [3] Y. Li, R. J. Zhou, A. L. Densmore, *et al.*, "Continental Dynamics and Geological Response of the Eastern Margin of Qinghai-Tibetan Plateau," Geological Publishing House, Beijing, 2006.
- [4] Y. Li, R. J. Zhou, A. L. Densmore, *et al.*, "The Geology of the Eastern Margin of the Qinghai-Tibet Plateau," Geological Publishing House, Beijing, 2006.
- [5] Y. Li, R. J. Zhou, A. L. Densmore, *et al.*, "The Geology of the Eastern Margin of the Qinghai-Tibet Plateau," Geological Publishing House, Beijing, 2006.
- [6] Y. Li, R. J. Zhou, A. L. Densmore and M. A. Ellis, "Strike-Slip Direction of Longmen Shan Fault and Sediments and Geomorphic Signs," *Minerals and Rocks*, Vol. 26, No. 4, 2006, pp. 26-34.
- [7] Y. Li, R. J. Zhou, A. L. Densmore and M. A. Ellis, "Strike-Slip of Eastern Margin in Qinghai-Tibetan Plateau since Cenozoic-Geomorphic Signs of Thrust Movement," *Quaternary Research*, Vol. 26, No. 1, 2006, pp. 40-51.
- [8] Y. Li, R. J. Zhou, A. L. Densmore, *et al.*, "Strike-Slip Process in Longmen Shan of Eastern Margin of Qinghai-Tibetan Plateau and Sedimentary Response," *Sedimentary Journal*, Vol. 24, No. 2, 2006, pp. 1-12.
- [9] Y. Li, R. J. Zhou, S. L. Dong, *et al.*, "Wenchuan Earthquake Thrust and Strike-Slip and Surface Rupture," *Journal of Chengdu University of Technology (natural science)*, Vol. 35, No. 4, 2008, pp. 404-413.
- [10] Y. Li, R. J. Zhou, A. L. Densmore, *et al.*, "Surface Rupture, Thrusting and Strike—Slipping in the Wenchuan Earthquake," The Gongwana 13 Program and Abstracts, 2008, pp. 114-115.
- [11] Y. Li, R. J. Zhou, A. L. Densmore, *et al.*, "Surface and Deformation of the Yingxiu-Beichuan Fault by the Wenchuan Earthquake," *ACTA Geologica Sinica*, Vol. 82, No. 12, 2008, pp. 1688-1706.

- [12] R. J. Zhou, Y. Li, A. L. Densmore, M. A. Ellis, Y. L. He, Y. Z. Li and X. G. Li, "Active Tectonics of the Longmen Shan Region on the Eastern Margin of the Tibetan Plateau[J]," *ACTA Geologica Sinica*, Vol. 81, No. 4, 2007, pp. 593-604.
- [13] R. J. Zhou, Y. Li, A. L. Densmore, M. A. Ellis, Y. L. He, F. L. Wang and X. G. Li, "Active Tectonics of the Eastern Margin of the Tibet Plateau," *Journal of Mineral Petrol*, Vol. 26, No. 2, 2006, pp. 40-51.
- [14] A. L. Densmore, Y. Li, M. A. Ellis, *et al.*, "Active Tectonics and Erosional Unloading of Eastern Margin," *Journal of Mountain Science*, Vol. 2, No. 2, 2005, pp. 146-154.
- [15] A. L. Densmore, M. A. Ellis, Y. Li, R. J. Zhou, G. S. Hancock and N. J. Richardson, "Active Tectonics of the Beichuan and Pengguan Faults at the Eastern Margin of the Tibetan Plateau," *Tectonics*, Vol. 26, 2007, p.17.
- [16] S. G. Liu, "Longmen Shan Thrust and Evolution of Foreland Basin in West Sichuan," Chengdu University of Technology Press, Chengdu, 1993.
- [17] Y. T. Chen, N. S. Xu and Y. Zhang, "Report of Focal Characteristics of Wenchuan Earthquake on May 12th, 2008," 2008. Internet Available: <http://www.csi.ac.cn>
- [18] Y. Yagi and Y. Fukahata, "Importance of Covariance Components in Inversion Analyses of Densely Sampled Observed Data: An Application to Waveform Data Inversion for Seismic Source Processes," *Geophysical Journal International*, Vol. 175, 2008, pp. 215-221.
- [19] P. C. England and P. Molnar, "Right-Lateral Shear and Rotation as the Explanation for Strike-Slip Faulting in Eastern Tibet," *Nature*, Vol. 344, 1990, pp. 140-142.
- [20] J. P. Avouac and P. Tapponnier, "Kinematic Model of Active Deformation in Central-Asia," *Geophysical Research Letters*, Vol. 20, No. 10, 1993, pp. 895-898.
- [21] Z. Chen, B. C. Buchfiel and Y. Liu, "Global Positioning System Measurements from Eastern Tibet and their Implication for India/Eurasia Intercontinental Deformation," *Journal of geophysical Research*, Vol. 105, No. B7, 2000, pp. 16215-16228.
- [22] P. Z. Zhang, Z. K. Sheng, M. Wang, W. J. Gan, R. Brüggmann, P. Monlar, Q. Wang, Z. J. Niu, J. Z. Sun, J. C. Wu, H. R. Sun and X. Z. You, "Continuous Deformation of the Tibetan Plateau from Global Positioning System Data," *Geology*, Vol. 32, 2004, pp. 809-812.
- [23] Y. Li, G. D. Xu, R. J. Zhou, *et al.*, "Isostatic Gravity Anomaly in Longmen Shan and its Constraint in Crustal Uplift of the Eastern Margin of the Qinghai-Tibetan Plateau," *Geological Bulletin of China*, Vol. 24, No. 12, 2005, pp. 1162-1169.